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"Triassic of the Tethys Realm"  
A contribution to the Project

## Uppermost Triassic sequences of the Choč nappe (Hronic) in the West Carpathians of Slovakia and Poland

**ABSTRACT:** The paper presents lithological and stratigraphical analysis of the uppermost Triassic deposits of the Choč unit of the Strážovská hornatina (Slovakia) and West Tatra Mts (Poland). The unique lithological and paleontological characteristics of these deposits permit recognition of a new lithostratigraphical unit, the Norovica Formation, in the Triassic of the West Carpathians. The formation includes light-grey compact limestones overlying the Hauptdolomit and underlying the Lower Liassic crinoidal limestones. Three members are distinguished within the Norovica Formation, viz. Lower Limestone Member, Sáva Woda Limestone Member, and Mojtn Limestone Member. The Norovica Formation contains the conodonts *Misikella posthernsteini* Kozár & Mock and the foraminifers *Triasina hantkeni* Majzon indicative of the Rhaetian. Only the lowermost part of the formation may represent the Upper Norian (Sevatian).

### INTRODUCTION

The uppermost Triassic deposits of the Choč unit (Hronic) in the West Carpathians are only fragmentarily preserved, as they have underwent an erosion in several sections due to the Early Kimmerian epeirogenic movements. These are mostly light-grey to grey, compact, organo-detritic limestones resembling in lithology the Dachstein Limestone. They occur in the northwest part of the Strážovská hornatina (cf. Foetterle 1864; Kulcsar 1915, 1916; Mahel 1946, 1962; Kochanová 1959, 1962, 1967), the north Male Karpaty Mts (Mahel 1958, Kochanová 1964), northern slopes of the Nizke Tatry Mts (Biely 1962), and the Polish West Tatra Mts (Guzik 1959; Gaździcki & Zawidzka 1973; Gaździcki 1978a, b). Only in Hybe region the uppermost Triassic strata are re-

presented by dark limestones intercalated with marls and marly shales recognized for the Hybe Beds (cf. Stache 1868, Goetel 1917, Michalik 1973, Gaździcki & *al.* 1979).

The below presented stratigraphical and facies analysis of the uppermost Triassic of the Choč unit is based upon sections of the Strážovská hornatina in Slovakia, and of the Polish West Tatra Mts (Text-fig. 1).

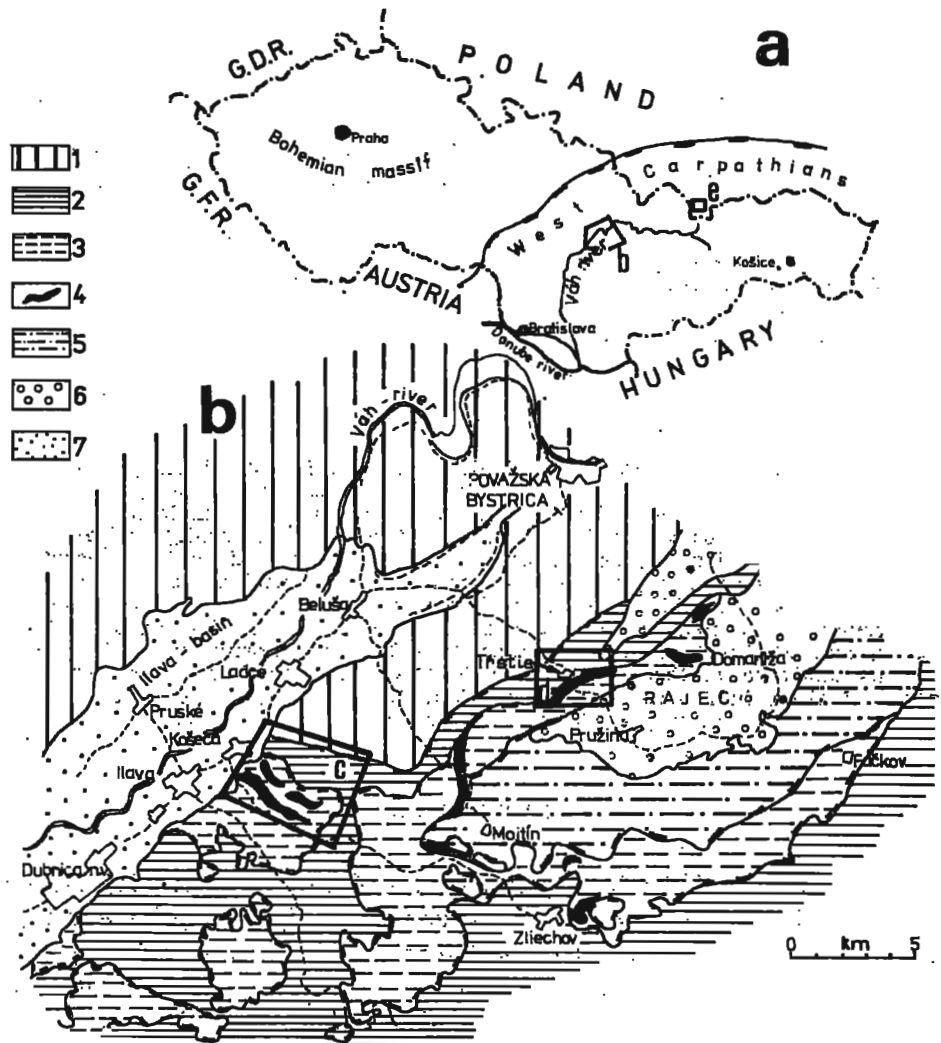


Fig. 1. Localization of the investigated areas in Czechoslovakian and Polish West Carpathians (a), and geological sketch-map of the middle Váh Valley (b) with

delimitation of the investigated areas (c and d — see Text-figs 5 and 7)

1 Klippen Belt (Jurassic-Cretaceous), 2 Kráľová nappes (Lower Jurassic-Middle Cretaceous), 3-4 Choč nappes (Middle Triassic-Lower Cretaceous), 5 uppermost Triassic limestones of the Choč unit (Norovica Formation), 6 Strážov nappes (Middle Triassic), 7 Paleogene sediments, 7 Neogene sediments

## CHARACTERISTICS OF THE DEPOSITS

The sedimentary sequence and microfacies of the uppermost Triassic of the Choč unit is presented after some selected sections from the Chochołowska and Lejowa valleys in the West Tatra Mts, and the sections of Mt. Norovica and Trstie in the Strážovská hornatina (cf. Text-figs 1—9 and Pls 1—10).

## WEST TATRA MTS

The uppermost Triassic strata of the Choč unit show considerable tectonic disturbances between the Chochołowska and Lejowa valleys. This may be due to post-orogenic gravitational tectonics of pre-Paleogene age: the uppermost Triassic limestones occur in form of isolated blocks scattered in the megabreccia of Middle to Upper Triassic dolomites (Text-fig. 2).

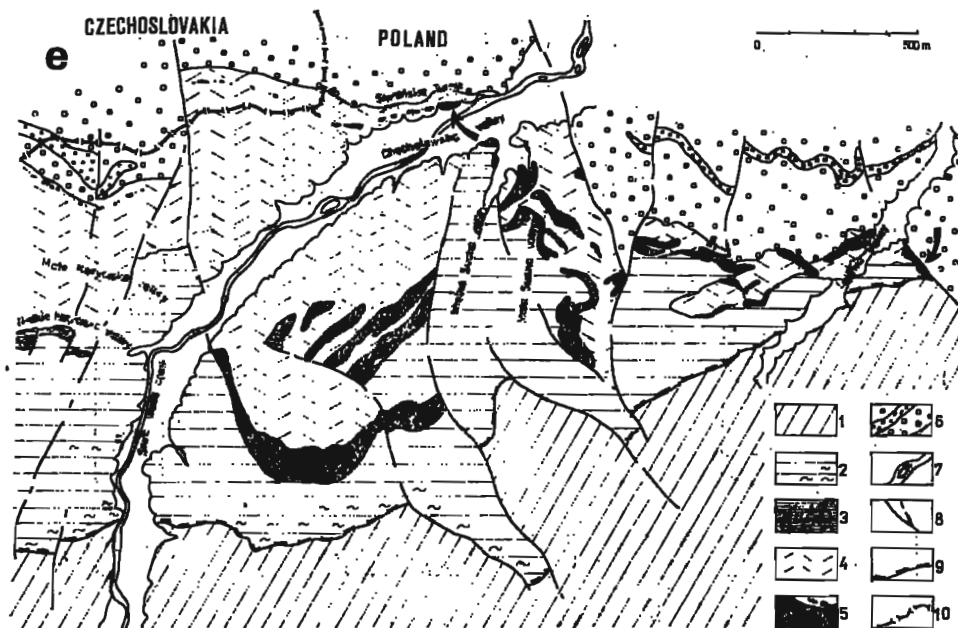


Fig. 2. Geological sketch-map of the area between the Lejowa and Chochołowska valleys, West Tatra Mts (for localization see area e in Text-fig. 1a)

1 Križna nappe (Triassic — Lower Cretaceous); 2—5 Choč nappe (Middle & Upper Triassic): 2 Middle Triassic dolomites, limestones and breccias (lower right corner); 3 Partnach Beds; 4 megabreccia of Middle and Upper Triassic dolomites; 5 uppermost Triassic limestones (Norovica Formation) preserved as blocks in the megabreccia; arrows indicate sections of the Norovica Formation

## CZOCZOLOWSKA VALLEY SECTION

The uppermost Triassic rocks are best exposed at the western slope of the Czocholowska Valley at the foot of Ślwińska Turnia (Text-fig. 2 and Pl. 1). The investigated, 9 m thick sequence comprises light-gray to grey limestones in normal position (Text-fig. 3 and Pl. 1). Three lithological sets are distinguished in the sequence:

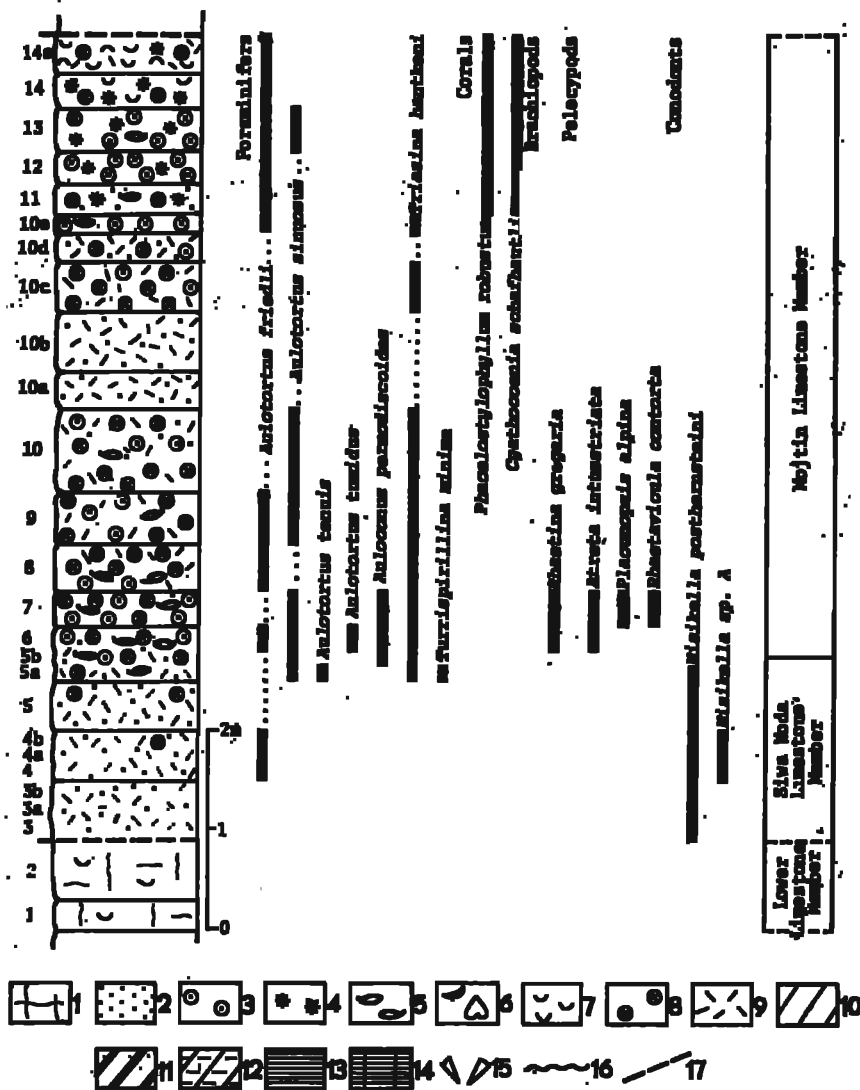
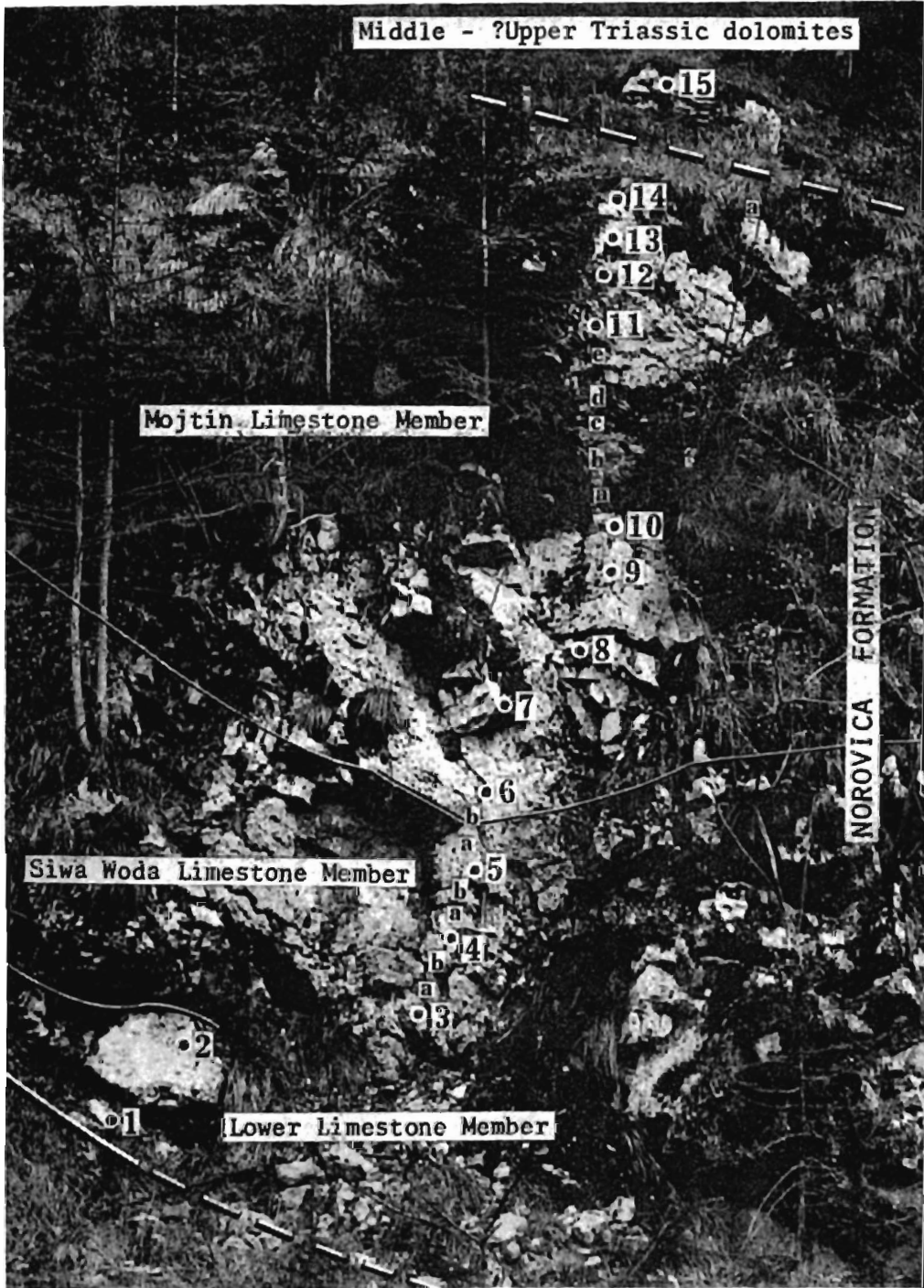
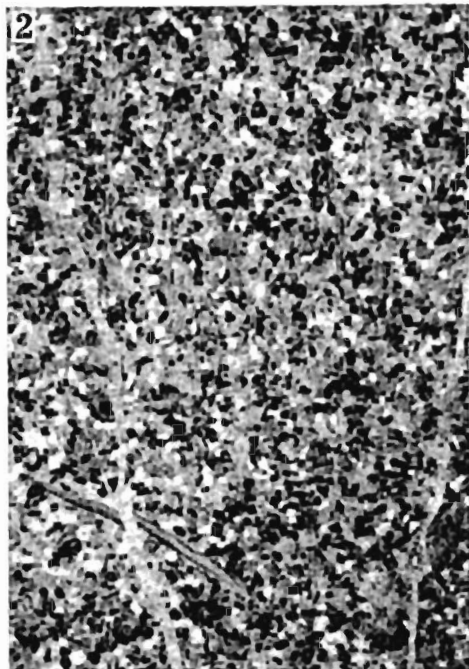


Fig. 3. Hypostrototype section of the Norovica Formation at the foot of Ślwińska Turnia in the Czocholowska Valley (s. in Text-fig. 1a, see also Text-fig. 2 and Pl. 1); the section comprises lithology and distribution of some important organic components

1 limestone, 2 sandy limestone, 3 white limestone, 4 coral limestone, 5 brachiopod limestone, 6 *Megastrophia* limestone, 7 *Rhynchonella*, 8 *Strophostrophia* limestone, 9 organostratal limestone, 10 *Strophostrophia*, 11 *Megastrophia*, 12 *Rhynchonella*, 13 *Megastrophia*, 14 *Rhynchonella*, 15 *Megastrophia*, 16 *Rhynchonella*, 17 *Megastrophia*



Exposure of the uppermost Triassic limestones of the Choč unit (hypostratotype section of the Norovica Formation) at the foot of Siwiańskie Turnie in the Chochołowska Valley; sampling sites (1—15) are indicated; photo taken in May 1978



Microfacies from Chochołowska Valley section; taken  $\times 10$

- 1 — Laminated micrite; Lower Limestone Member, layer 1
- 2 — Sandy biopelsparite containing the youngest Triassic conodonts of the genus *Misikella*; Siwa Woda Limestone Member, layer 4
- 3 — Crinoid-brachiopod bioosparite; Mojtn Limestone Member, layer 5b
- 4 — Biosparrudite composed of crinoid and brachiopod debris as well as of subordinate ooids and foraminifers *Aulotortus sinuosus*; Mojtn Limestone Member, layer 7

LAYERS 1—2: Laminated micrites (Pl. 2, Fig. 1), some 80 cm thick, with occasional remains of pelecypods, gastropods, and ostracodes associated with some foraminifers *Glomospira*.

LAYERS 3—5a: Sandy biopelsparites (Pl. 2, Fig. 3), some 180 cm thick, containing small fragments of crinoids and the conodonts *Misikella posthernsteini* Kozur & Mock and *Misikella* sp. A (*sensu* Gaździcki 1978b). The foraminifers are represented by *Glomospira*, *Trochammina*, *Agathammina*, *Nodosaria*, and very rare *Aulotortus friedli*; *Triasina hantkeni* appears at the top of the set. There are also the algae *Aciculella*. It is to be noted that the stratigraphically important conodonts *Misikella posthernsteini* co-occur in the uppermost part of the set (layer 5a) with the large benthic foraminifers *Triasina hantkeni*.

LAYERS 5b—14a: Crinoid-brachiopod biocosparites with corals, pelecypods, and intraclasts (Pl. 2, Fig. 4 and Pl. 6, Figs 1—3). There is also a rich foraminifer assemblage representative of the family Involutinidae Bütschli, 1880, *sensu* Piller (1978), including *Aulotortus friedli*, *A. sinuosus*, *A. tenuis*, *A. tumidus*, *Auloconus permodiscoides*, and especially *Triasina hantkeni*. The species *Turrispirillina minima* has also been recorded in the set. The brachiopod assemblage is dominated by *Rhaetina gregaria*, while the pelecypods are represented by *Atreta intusstriata*, *Placunopsis alpina*, and a few specimens of *Rhaetavicula contorta*. Coral colonies assigned to *Cyathocoenia schafhautli* and *Phacelostylophyllum robustum* in life position occur in the uppermost part of the set (layers 11—14a).

#### LEJOWA VALLEY SECTION

The uppermost Triassic rocks of the Choč unit are exposed at the eastern slope of the valley close to its mouth (Text-fig. 2 and Pl. 3). These are light-grey to grey organodetrital limestones intercalated here and there with loferitic and limy dolomites (Text-fig. 4). The exposed sequence is some 13 m thick but both the lower and upper boundaries of the unit are covered with waste deposits. The brachiopods *Rhaetina gregaria* and the pelecypods *Atreta intusstriata* occur rather commonly. The pelecypods *Rhaetavicula contorta*, *Placunopsis alpina*, and *Modiolus schafhautli*, and the corals *Phacelostylophyllum robustum* and *Retiophyllia* sp. have been found in the waste deposits. The rocks are represented mostly by biopelmicrites (Pl. 4, Fig. 1) and biotrasparites (Pl. 5, Fig. 3) containing brachiopod, crinoid, and gastropods debris associated commonly with pellets, intraclasts, and/or ooids as well as onkolitic crusts. The abundant foraminifer assemblage (cf. Pl. 5, Figs 1—2) includes *Aulotortus friedli*, *A. sinuosus*, *A. tenuis*, *Auloconus permodiscoides*, *Trocholina crassa*, and *Triasina hantkeni*.

The sequence resembles the uppermost set of the above described section from the Chochołowska Valley.

#### STRAŽOVSKÁ HORNATINA

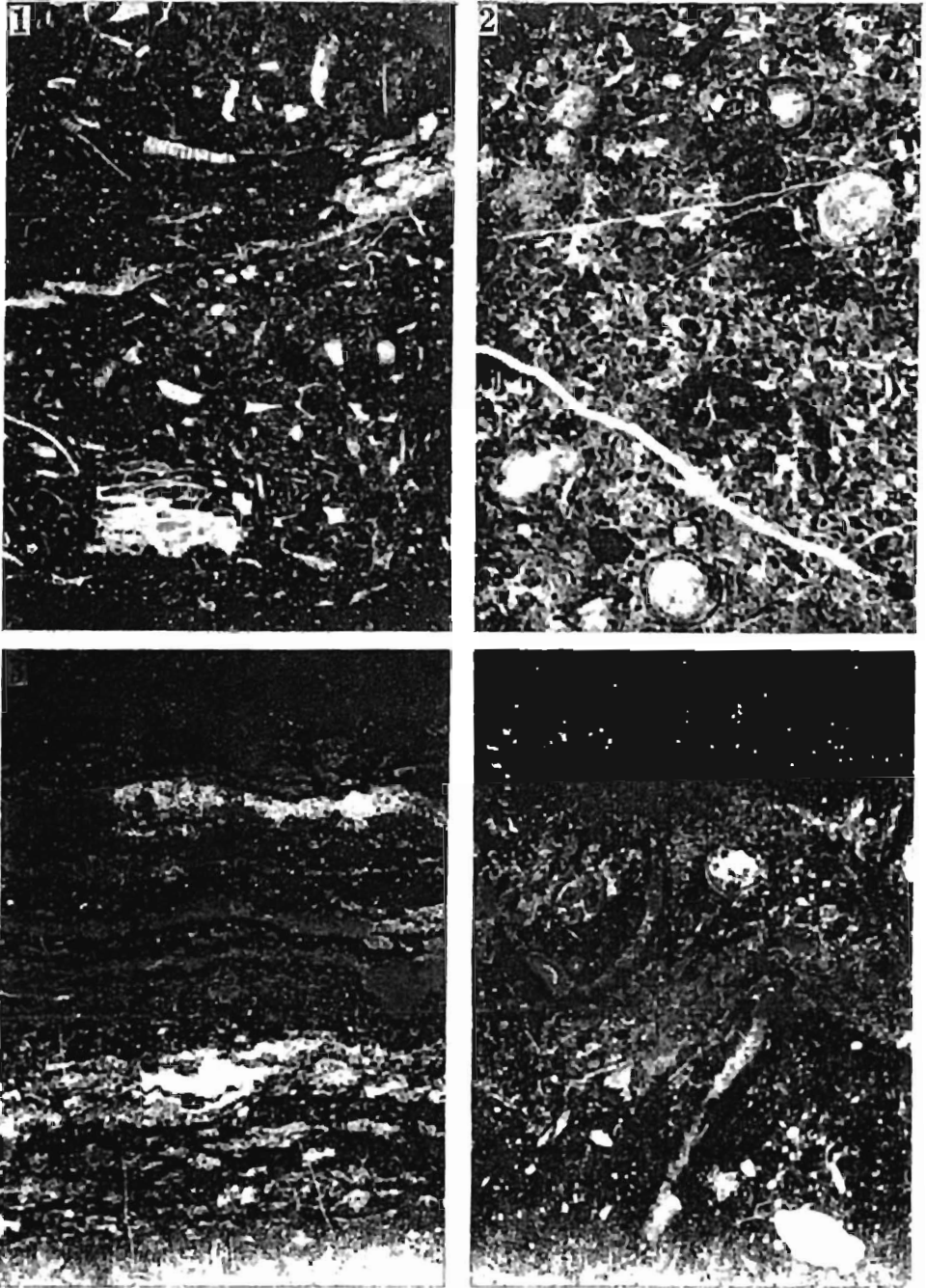
In the northwest part of the Strážovská hornatina, the uppermost Triassic limestones of the Choč unit overlie the Hauptdolomit of a considerable thickness (Text-fig. 5). They are overlain in turn by transgressive Lower Liassic crinoidal limestones yielding pelecypods of Hettangian age (Kochanová 1962). The Upper Triassic strata are cut by





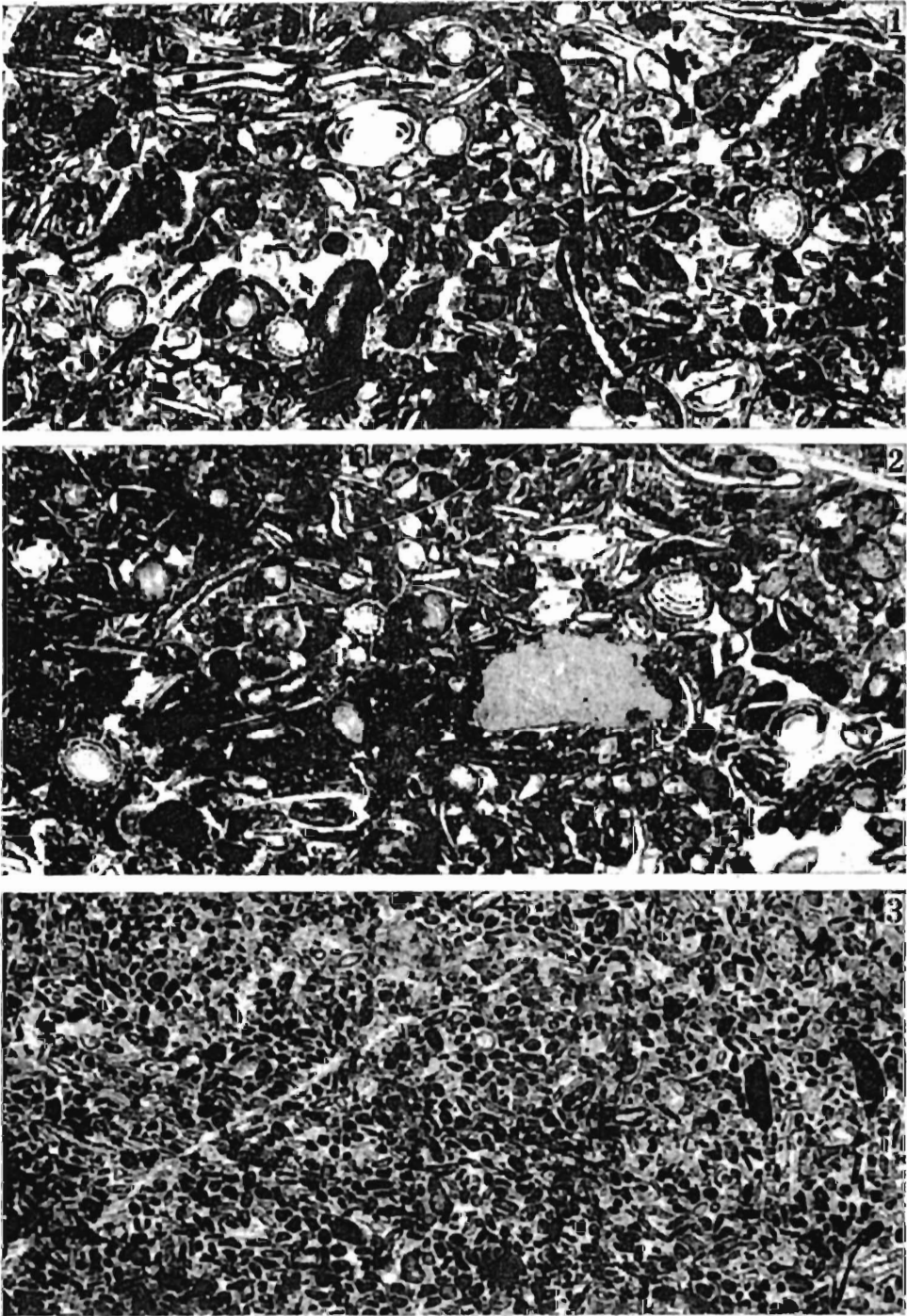


Exposure of the uppermost Triassic limestones of the Choć unit (Norovica Formation) in the Lejowa Valley; lower and middle parts of the section are exposed (sampling sites 1—11); photo taken in August 1978

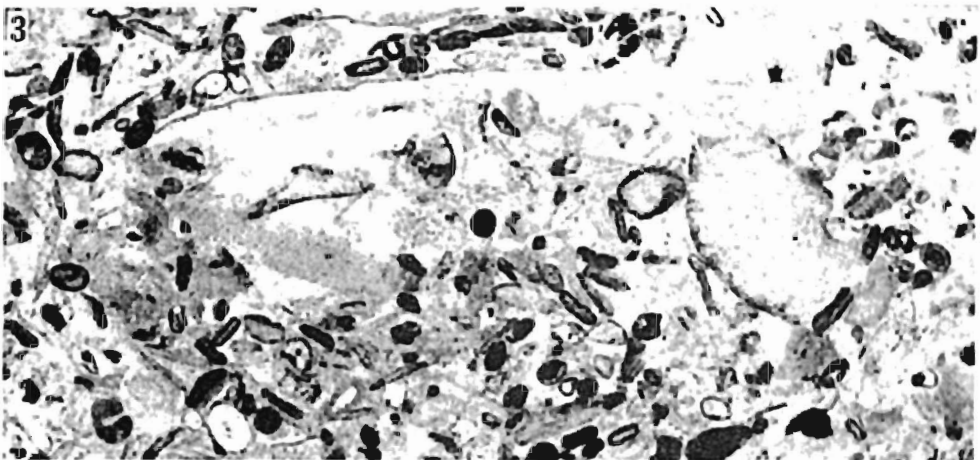
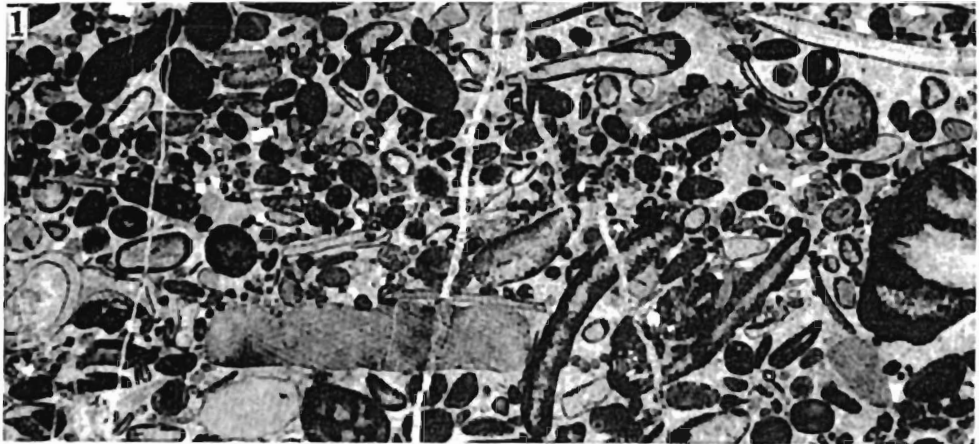


Lejowa Valley section, Mojtin Limestone Member; taken  $\times 10$

- 1 — Biopelmicrite composed of brachiopod and corals debris as well as of pellets; layer 3
- 2 — Biopelmparenite composed of strongly crushed bioclasts and some foraminifers *Triasina hantkeni*; layer 9
- 3 — Loferitic limy dolomite with algal lamination; layer 10
- 4 — Foraminifers *Triasina hantkeni* in biopelmicrite; layer 12a



Lejowa Valley section (cnt'd), Mojtin Limestone Member; taken X10  
 1-2 — Microfacies with foraminifers *Triasina hantkeni*, *Autotortus friedli*, *Autotortus tumidus* and *Auloconus permodiscoides* in brachiopod-crinoid biosparite; layer 13  
 3 — Biointrasparite composed of brachiopod, crinoid and gastropod debris, as well as of intraclasts and ooids; layer 17



Chochołowska Valley section (cnt'd from Pl. 2), Mojtin Limestone Member; taken  
 ×10

- 1 — Biosparrrudite containing crinoid and brachiopod debris with onkolitic crusts as well as of ooids and foraminifers *Triasina hantkeni*; layer 9
- 2 — Crinoid bioosparite with *Triasina hantkeni*; layer 10c
- 3 — Crinoid-brachiopod biosparite containing bioclasts with onkolitic crusts and some ooids; layer 13

covered with waste deposits. The rocks contain abundant brachiopods *Rhaetina gregaria* (Pl. 9, Fig. 4), pelecypods *Placunopsis alpina*, *Atrreta intusstrata*, *Mysidioptera* sp., and unidentifiable megalodontids. Fish teeth occur sporadically at the base of the section. The dominant microfacies are biocoenocrites with brachiopod, crinoid, and pelecypod debris and onkolitic crusts as well as ooids (Pl. 9, Fig. 4 and Pl. 10, Figs 3—4). Biopelmicrites (Pl. 9, Fig. 2) and biotrasparites (Pl. 10, Fig. 2) occur subordinately. Foraminifers occur abundantly throughout the sequence, including *Triasina hantkeni* (Pl. 9, Figs 3—5) associated with *Aulotortus friedli* (Pl. 10, Fig. 3). The species *Miliolipora cuvillieri* has been recorded in the lower part of the sequence.

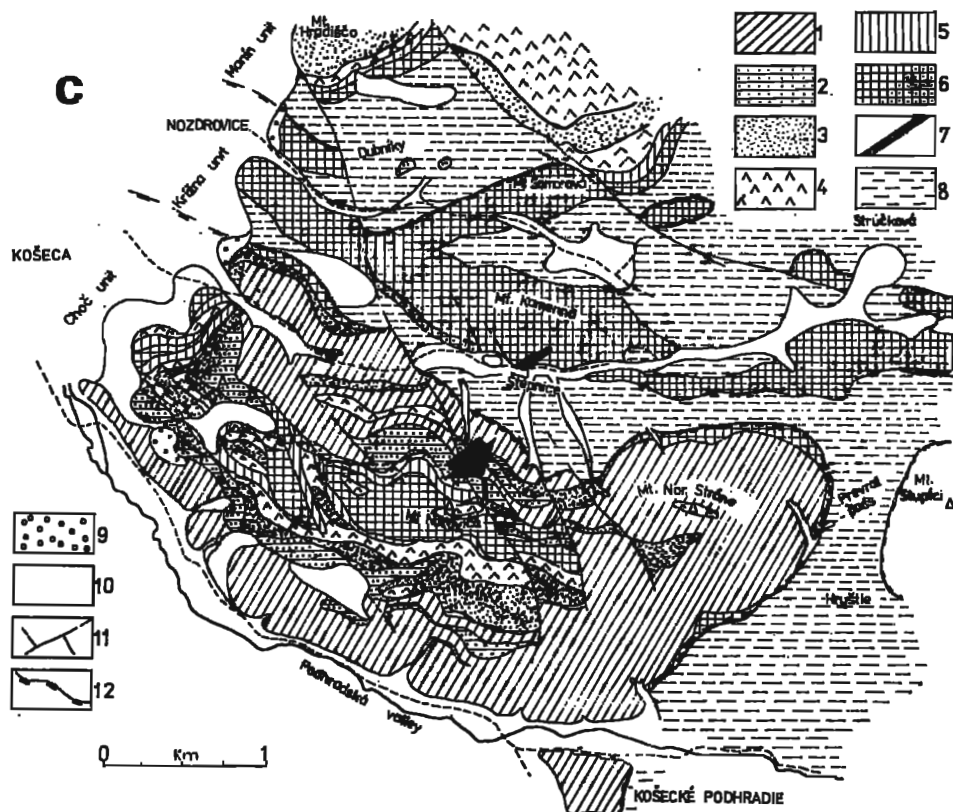


Fig. 5. Geological map of the area between Košeca and Košecké Podhradie (for localization see area c in Text-fig. 1b); arrow indicates stratotype section of the Norovica Formation

- 1 Middle and Upper Triassic dolomites, 2 uppermost Triassic limestones of the Norovica Formation, 3 Lias crinoidal limestones, 4 Dogger cherty limestones, 5 Malm fine-grained limestones, 6 Lower Cretaceous marly limestones and Aptian black organodetrital limestones (lower right corner), 7 Aptian? tuffites, 8 Albian-Cenomanian shales with intercalations of limestones (in the lowermost part) and sandstones, 9 Neogene conglomerates and sandstones, 10 Quaternary deposits, 11 faults, 12 main overthrusts



## TRISTIE SECTION

The uppermost Triassic rocks are exposed at the southwest slope of Mt. Trudovac on the Pružinka creek (Text-figs 7—8). The investigated sequence is located at the road cut Trstie-Pružina, 1 km southwest off Trstie (Text-fig. 9); it approximates 46 m in thickness. Some samples (M1-M28) were also taken for comparative purposes from a 14 m thick sequence located 50 m above the creek bed (Text-fig. 8); the latter section was studied by Račková (1979). The

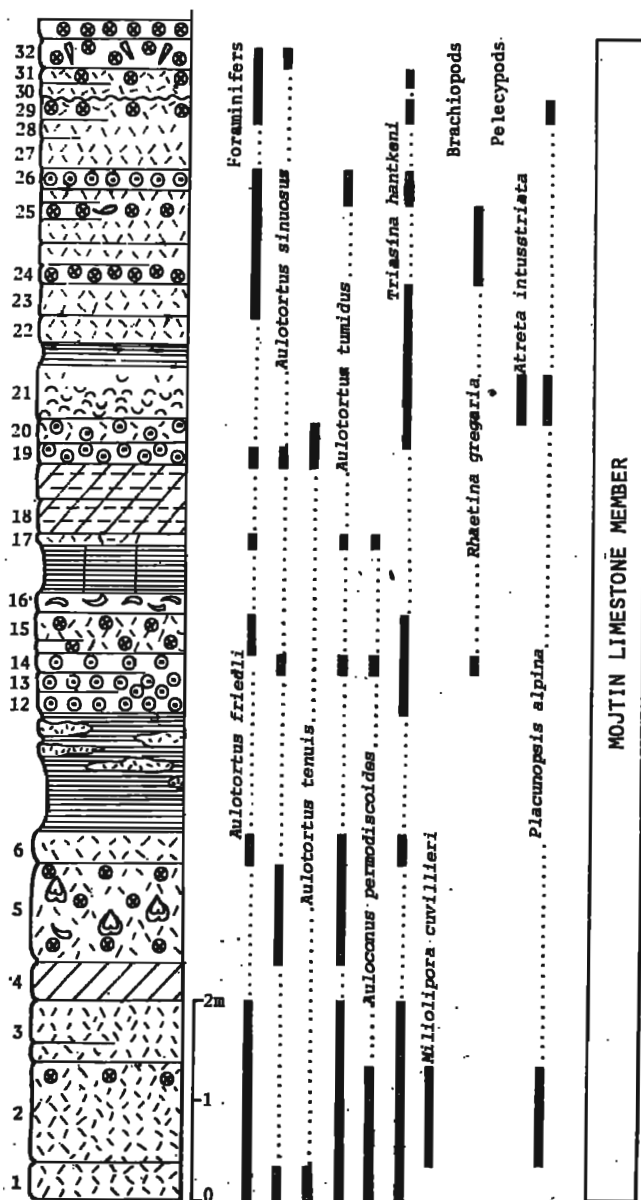


Fig. 6. Stratotype section of the Norovica Formation at Mt. Norovica (c. in Text-fig. 1b see also Text-fig. 5); explanations the same as for Text-fig. 3

boundary with the Hauptdolomit is covered with waste deposits. In turn, the contact with the Lower Liassic crinoidal limestones appears clearly in the additional, higher-located section (Text-fig. 8). The main investigated sequence is split into 6 parts by dislocations but nonetheless, the subunits can be easily correlated with one another. The uppermost Triassic rocks include light-grey, compact, organodetrital limestones intercalated commonly with oolitic limestones (Text-fig. 9). There are also some intercalations of brachiopod and coral limestones, dolomites, and marls. The recorded macrofaunal assemblage includes the corals *Retiophyllia clathrata* and *R. paraclathrata*, abundant brachiopods *Rhaetina gregaria* and *Zugmayerella uncinata*, pelecypods *Atreta intusstriata* and *Placunopsis alpina*, and unidentified gastropods. The rocks are represented mostly by brachiopod-crinoid biopelmicrites (Pl. 7, Figs 2 and 4) and biocosparenites with brachiopod, pelecypod and crinoid debris and onkolitic crusts and oolitic coatings (Pl. 8, Figs 4-6). Foraminifers occur especially abundantly in biopelsparites and biopelmicrites; they include mostly *Triasina hantkeni* (Pl. 7, Fig. 3 and Pl. 8, Figs 1-2) associated with *Aulotortus friedli*, *A. sinuosus* (Pl. 10, Fig. 1), *A. tenuis*, *A. tumidus*, *Auloconus permodiscoides* (Pl. 8, Figs 1-2), *Trocholina crassa*, and *Seminvoluta clari*. The higher-located section yielded also the spores *Globochaete alpina*, algae *Cylindroporella* sp., and holothurian sclerites *Theelia seniradiata*, *Th. stellifera*, and *Th. variabilis* (see Račková 1979).

#### SEDIMENTARY SEQUENCE AND IMPORTANT FAUNAL ASSEMBLAGES

The presented characteristics of the uppermost Triassic of the Choč nappe (Hronic) demonstrates clearly that these rocks are largely different from their time equivalents of the high-tatric (Tatric) and sub-tatric (Križna = Patric) units.

First of all, the investigated rocks are close in lithology to the Dachstein Limestone and show a unique position in the Triassic of the West Carpathians. They display a continuous sedimentary transition from the underlying Hauptdolomit, while they are transgressively overlain by the Lower Liassic crinoidal limestones. There is no such Upper Triassic to Lower Jurassic sequence in any other tectonic unit of the West Carpathians (cf. Michalik 1977).

The investigated rocks contain very rich and abundant associations of large benthic involutinid foraminifers dominated by the species *Triasina hantkeni* (see Text-figs 3-4, 6, and 9). This is the area with maximum frequency of the involutinids in the uppermost Triassic of the West Carpathians.

Furthermore, the investigated rocks yielded the latest known conodonts *Misikella posthernsteini* and *Misikella* sp. A *sensu* Gaździcki (1978b). The conodont frequency is here the highest among those few thus far known localities with *Misikella posthernsteini* (cf. Gaździcki 1978a,b).

The investigated rocks contain also the corals *Phacelostylophyllum robustum*, *Pinacophyllum lejowae*, and *Cyathocoenia schafhautli* (see Roniewicz 1974).

These unique characteristics of the uppermost Triassic of the Choč unit (Hronic) makes the basis for recognition of a new lithostratigraphic unit in the Triassic of the West Carpathians, the Norovica Formation.

### NOROVICA FORMATION

**NAME:** After Mt. Norovica, Strážovská hornatina, Slovakia (cf. Text-fig. 5).

**GENERAL LITHEOLOGY:** Light-gray to grey, compact, organodebtitic limestones close to the Dachstein Limestone. There are intercalations of corallitic, crinoidal, brachiopod, coral, and *Megalodon* limestones (Text-figs 2-4, 6, and 9), with loteritic dolomites and marls in minor amounts.

**SYNONYMY:** Thus far, these rocks were called as the Rhaetic or Rhaetic grey limestones (Mahel 1964), or the grey limestones of the Hronic (Michalik 1977).

**TYPE LOCALITY:** Northern slope of Mt. Norovica, northwest part of the Strážovská hornatina, Slovakia (see Text-figs 5-6).

**HYPOSTRATIOTYPE LOCALITIES:** (1) at the foot of Sławiański Turnie in the Chochołowska Valley, West Tatra Mts, Poland (see Text-figs 2-3 and Pl. 1); and (2) near Trstie, Slovakia (see Text-figs 7-8).

**BOUNDARIES:** The lower boundary is at the top of the Hauptdolomit. The upper boundary is marked by the transgressive Lower Liassic crinoidal limestones (see Text-fig. 9).

**THICKNESS:** Total thickness approximates 50 m.

**FOSSILS:** As those listed in the text (Text-figs 2-4, 6, and 9).

**AGE:** ?Late Norian (Sevastian) to Rhaetic.

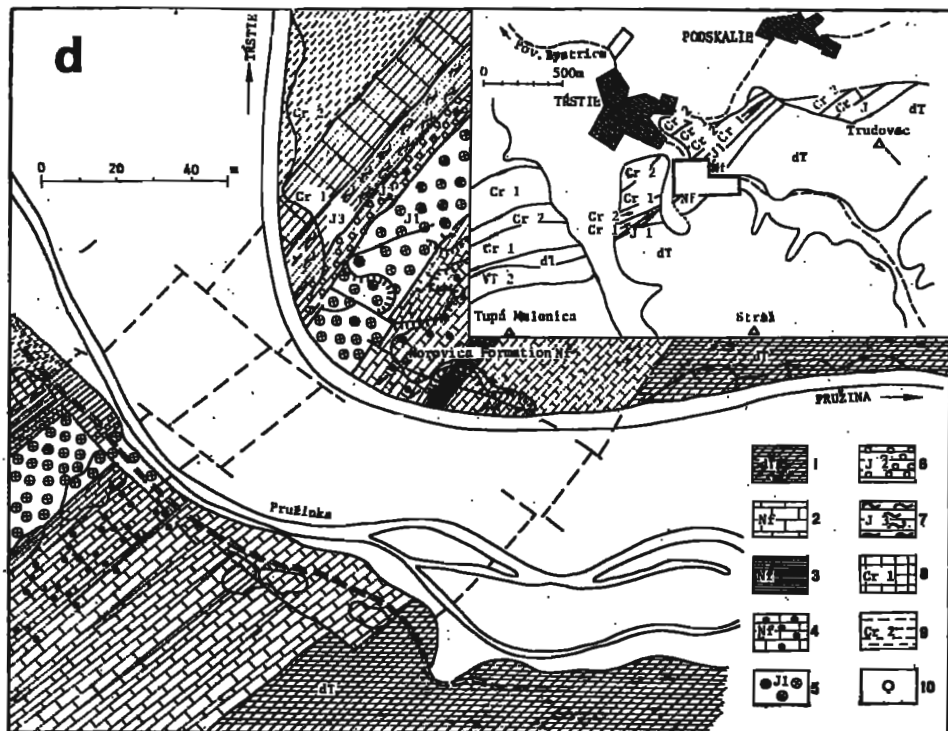


Fig. 7. Geological sketch-map of the Trstie area (for localization see area d in Text-fig. 1b)

1 Upper Triassic dolomites (dt), 2-4 uppermost Triassic sequence (Norovica Formation, Nf: 2 gray and light-gray limestones, 3 marls, 4 corallitic limestones), 5 Lower Liassic crinoidal limestones (J1), 6 Dogger cherty limestones (J3), 7 Malm gray marly limestones (J3), 8 Neocomian marly limestones (Cr1), 9 Albian marls (Cr2), 10 Quaternary deposits (Q)



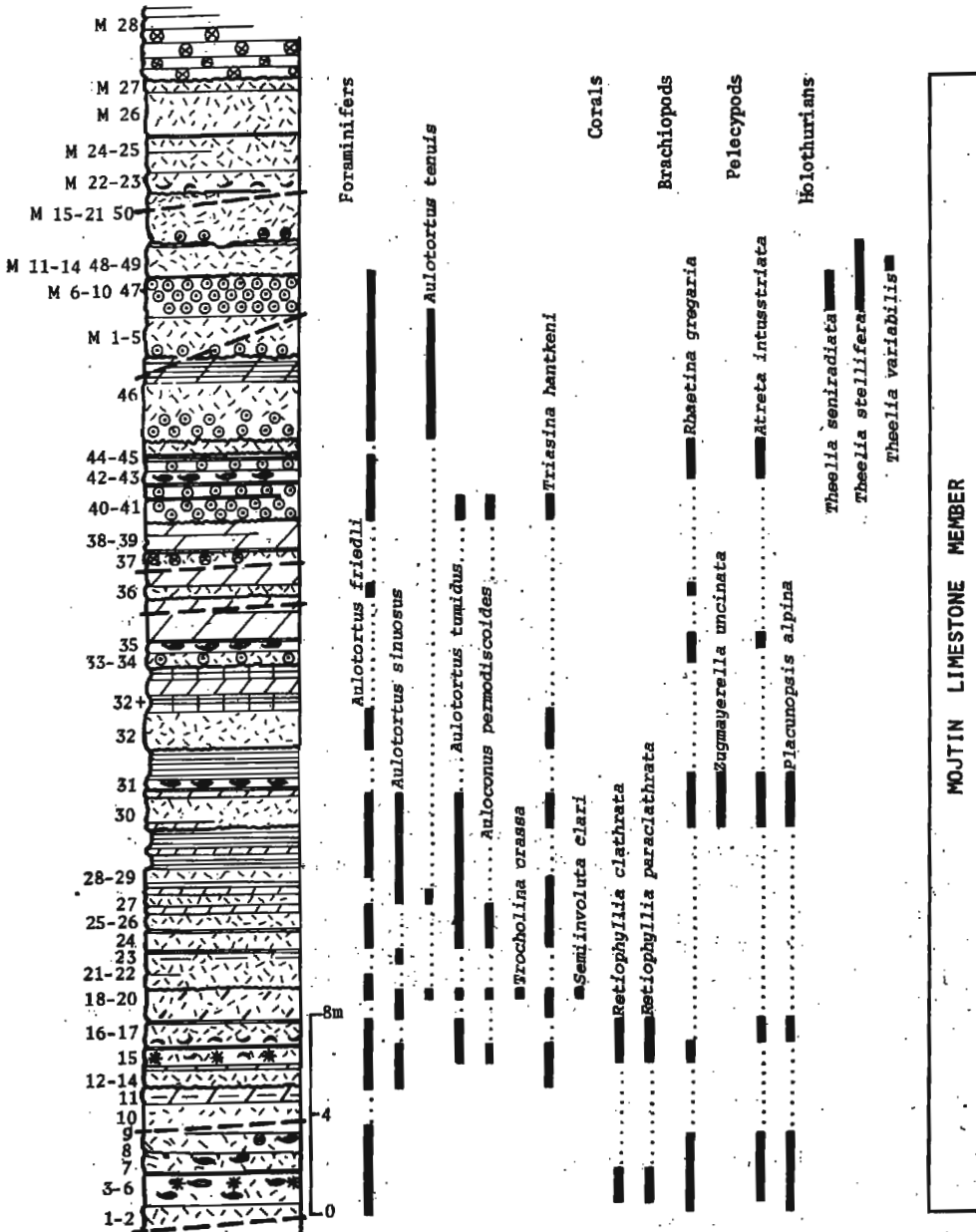
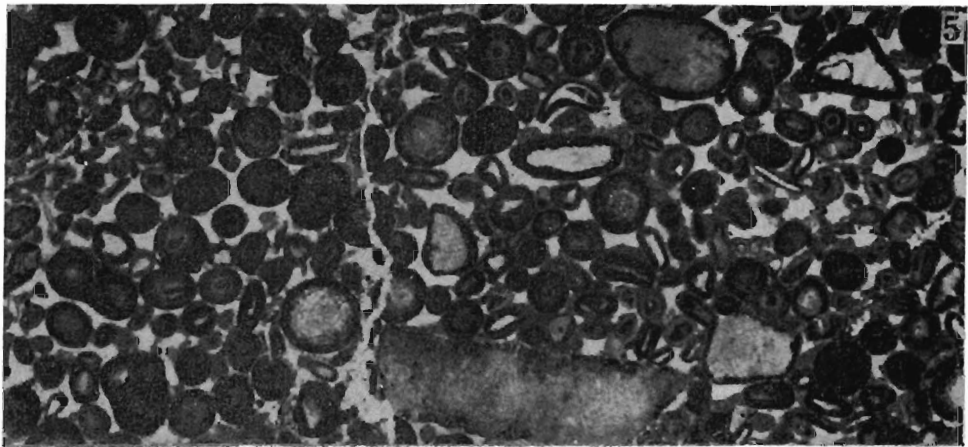
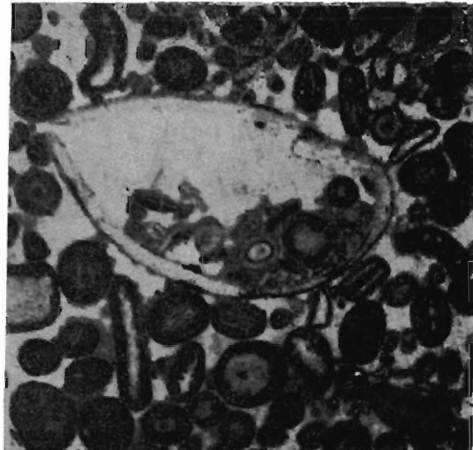
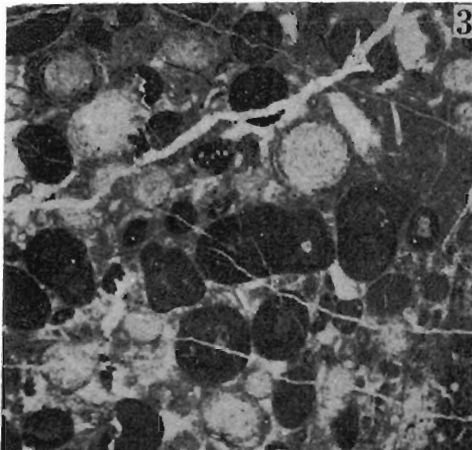
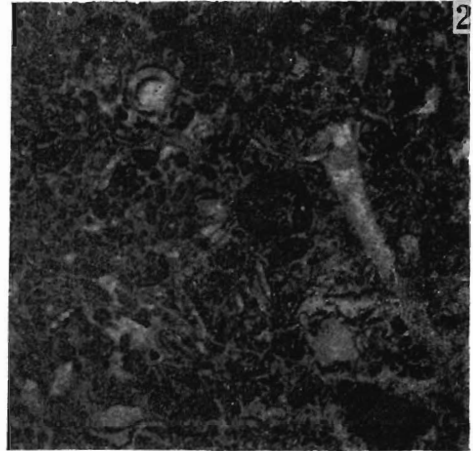
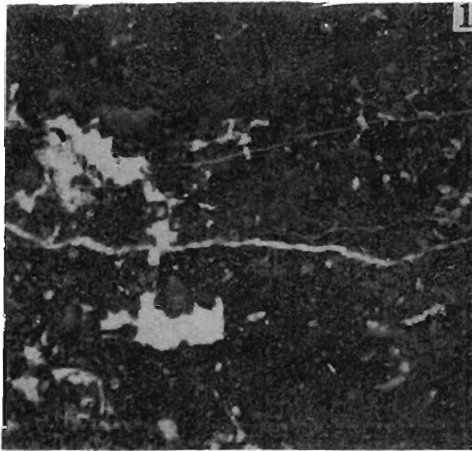
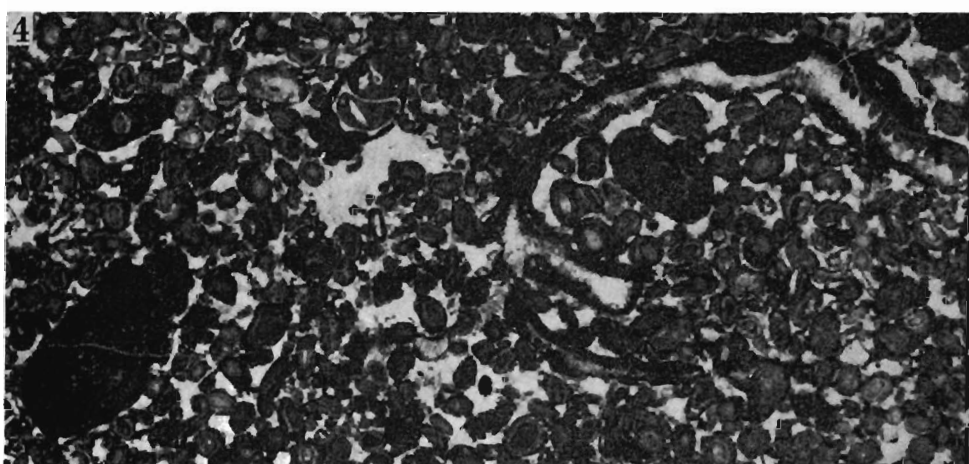
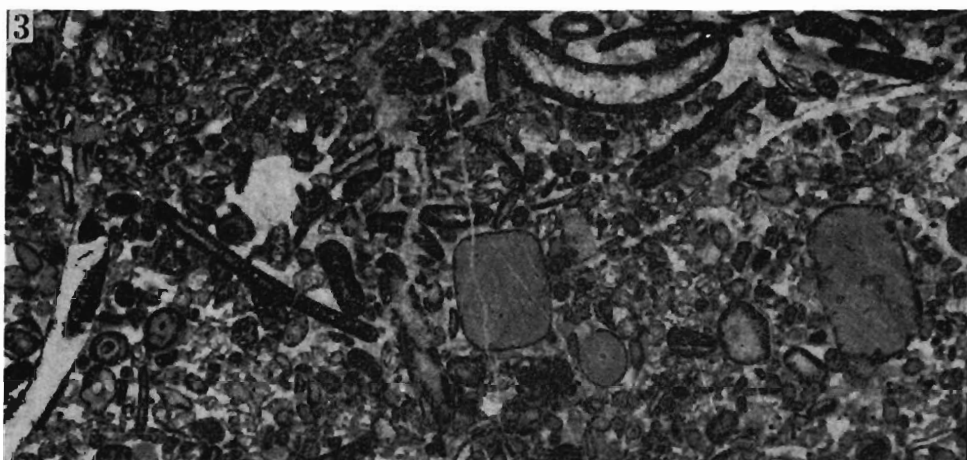
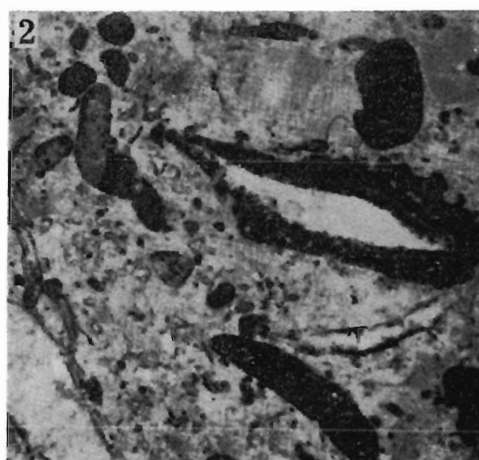
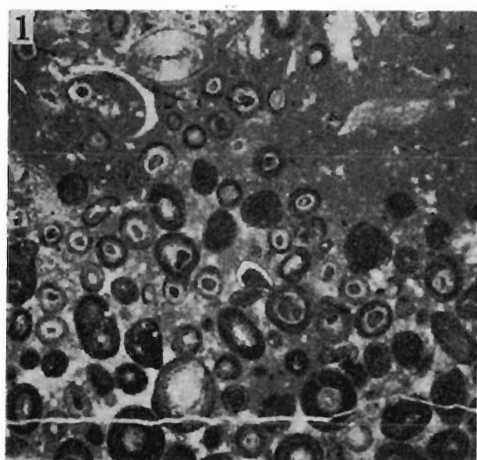


Fig. 9. Hypostratotype section of the Norovica Formation at Trstie (d in Text-fig. 1b see also Text-figs 7 and 8); explanations the same as for Text-fig. 3



Mt. Norovica section, Mojtin Limestone Member; taken  $\times 10$

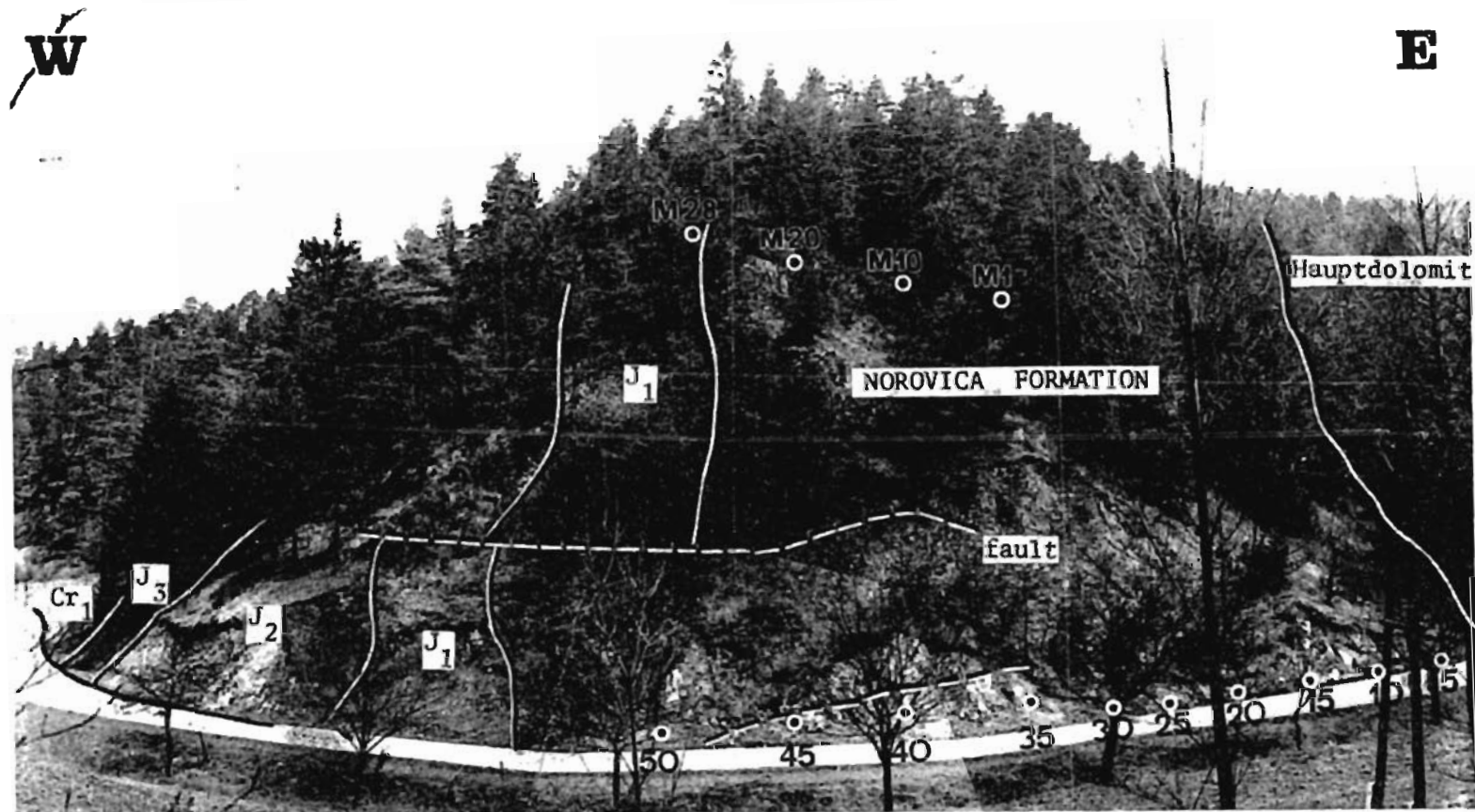
- 1 — Dolomiticrite with organic fragments (mostly pelecypod and crinoid debris); layer 4
- 2 — Biopelmicrite composed of strongly crushed bioclasts of pelecypod, crinoid and brachiopod debris as well as of foraminifers *Autotortus* and pellets; layer 6
- 3 — Association of foraminifers *Triasina hantkeni* in bioosparite; layer 14
- 4 — Brachiopod (*Rhaetina*) bioosparite; layer 15
- 5 — Bioosparite containing crinoid debris and foraminifers *Triasina hantkeni*; layer 15



Mt. Norovica section (cnt'd), Mojtin Limestone Member; taken  $\times 10$

- 1 — Bloomicrite with foraminifers *Triastina hantkeni* and *Aulotortus sinuosus*; layer 19
- 2 — Bioltrasparite composed of pelecypod and crinoid debris with onkolitic crusts as well as of intraclasts; layer 24
- 3 — Association of foraminifers *Aulotortus friedli* in crinoid-pelecypod-brachiopod biomicrite; layer 25
- 4 — Bloosparite containing larger pelecypod, gastropod and crinoid debris with onkolitic crusts as well as ooids; layer 28

Exposure of the Upper Triassic, Jurassic and Cretaceous sequence in the roadcut Třstie-Pružina (c 1 km SW off Třstie)



Sampling sites (5—50 and M1—M20) from the hypostratotype section of the Norovica Formation (uppermost Triassic of the Choč unit) are indicated; photo taken in May 1979  
 J1 Lias, J2 Dogger, J3 Malm, Cr1 Neocomian

## LITHOSTRATIGRAPHIC SUBDIVISION OF THE NOROVICA FORMATION

Three members are distinguished in the Norovica Formation after the nature of the sedimentary sequence, the microfacies, and the fauna. These are: the Lower Limestone Member, the Siwa Woda Limestone Member, and the Mojtin Limestone Member (see Text-figs 3 and 10, and Pl. 1).

	Crinoidal limestones of Lower Lias	Hettangian
NOROVICA FORMATION	Mojtin Limestone Member	Rhaetian
	Siwa Woda Limestone Member	
	Lower Limestone Member	?Upper Norian
	Hauptdolomit	Norian Carnian

Fig. 10. Lithostratigraphical subdivision of the Norovica Formation

## LOWER LIMESTONE MEMBER

This informal unit, most probably the lowermost part of the Norovica Formation, has been recorded only at the base of the hypostratotype section in the Chochołowska Valley (Text-fig. 3 and Pl. 1). It comprises grey to dark-grey laminated micrites with single fossils (Pl. 2, Fig. 1), ranging up to some 80 cm in thickness. The relationship to the overlying members of the Norovica Formation is unclear because of the tectonic nature of the contact. This unit is probably to be assigned to the Upper Norian (Sevatian), as it is overlain by the well documented Lower Rhaetian Siwa Woda Limestone Member (cf. Text-figs 3 and 10).

## SIWA WODA LIMESTONE MEMBER

This unit includes grey, compact, sandy, biopelsparite limestones (Pl. 2, Fig. 2) with abundant conodonts *Misikella posthernsteini* Kozur & Mock and *Misikella* sp. A *sensu* Gaździcki (1978b), and foraminifers *Aulotortus friedli* (Kristan-Tollmann). The name of the unit is after the Siwa Woda creek in the Chochołowska Valley (Text-fig. 2). The type section of the Siwa Woda Limestone Member is in the hypostratotype section of the Norovica Formation at the foot of Siwiańskie Turnie in the Chochołowska Valley (Text-fig. 3 and Pl. 1). The member attains some 180 cm in thickness. As judged from the occurrence of the conodonts *Misikella posthernsteini* and the foraminifers *Aulotortus friedli*,

the member is to be assigned to the Lower Rhaetian (cf. Gaździcki 1978b). The Siwa Woda Limestone Member has thus far not been recorded in the Strážovská hornatina.

#### MOJTIN LIMESTONE MEMBER

This is the basic subunit of the Norovica Formation; it comprises light-grey to brownish, compact, oolitic and organodetrital limestones of the "Dachstein facies" (Pl. 5, Figs 1—2; Pl. 6, Figs 1—3; and Pls 7—10). The following fossils have been recorded (cf. Kochanová 1962, Mahel & al. 1968) in the Mojtin Limestone Member in the Mojtiňská Valley, Strážovská hornatina:

Brachiopods: *Rhaetina gregaria* (Suess), *R. pyriformis* (Suess), *Zelleria austriaca* (Zugmayer), *Z. elliptica* (Zugmayer), *Zugmayerella uncinata* (Schafhäütl);  
 Pelecypods: *Atreta intusstriata* (Emmrich), *Cardita austriaca* (Heuer), *Gerpillia inflata* Schafhäütl, *G. praecursor* (Quenstedt), *Lima* cf. *discus* Stoppani, *Liostraea irregularis* (Münster), *Protocardia rhaetica* (Merlam), *Piacunopsis alpina* (Winkler), *Rhaeticula contorta* (Portlock);  
 Corals (reported for the first time): *Phacelostylophyllum robustum* Roniewicz, *Ph. medium* Roniewicz, *Retiophylla paraclathrata* Roniewicz, *Astracomorpha crassisepta* Reuss.

There are also abundant and rich involutinid foraminiferal associations dominated by *Triasina hantkeni* Majzón (cf. Pl. 4, Figs 2, 4; Pl. 5, Figs 1—2; Pl. 7, Fig. 3; and Pl. 9, Figs 3, 5). The name of the member is after the Mojtiňská Valley, Strážovská hornatina, where the type section is designated. Hypostratotype sections are in the Chochołowska Valley and in the Trstie region (cf. Text-figs 3 and 9). The Mojtin Limestone Member overlies the Siwa Woda Limestone Member in the Chochołowska Valley section (see Text-fig. 3 and Pl. 1). The upper boundary of the Mojtin Limestone Member coincides with the upper boundary of the whole Norovica Formation. This member attains up to 50 m in thickness and occurs in all the investigated geological sections of the West Tatra Mts and Strážovská hornatina (cf. Text-figs 3—4, 6, and 9). As judged from the distribution of the foraminifers *Triasina hantkeni*, the Mojtin Limestone Member is to be assigned to the uppermost Lower to Upper Rhaetian (cf. Gaździcki & al. 1979).

#### CONCLUSIONS

The new Norovica Formation is recognized for the unique uppermost Triassic sequence of the Choč nappe (Hronic) in the West Carpathians. The formation overlies the Hauptdolomit and underlies the Lower Liassic crinoidal limestones. It approximates 50 m in thickness, and ranges in age since the ?Late Norian (Sevatian) through the Late Rhaetian.

The Norovica Formation largely differs from its time equivalents of the high-tatric and sub-tatric (Križna) units, i.e. the Tomanova Formation and the Fatra Formation, respectively (cf. Michalik & al. 1976, Michalik 1977, Michalik & al. 1979).

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**LITOLOGIA I STRATYGRAFIA NAJWYŻSZEGO TRIASU JEDNOSTKI  
CHOCZAŃSKIEJ GÓR STRAŻOWSKICH I TATR ZACHODNICH**

(Streszczenie)

W oparciu o profile z Gór Strażowskich na Słowacji oraz dolin Chochołowskiej i Lejowej w Polskich Tatrach Zachodnich, przedstawiono następstwo osadów oraz charakterystykę mikrofacjalną utworów najwyższego triasu jednostki choczańskiej (*Hronicum*) Karpat Zachodnich (por. fig. 1—9 oraz pl. 1 i 3). Swoiste cechy tych osadów, jak również stwierdzony zespół elementów florystycznych i faunistycznych (por. pl. 2 oraz 4—10) stały się podstawą do wyróżnienia w obrębie utworów triasu Karpat Zachodnich nowej jednostki litostratygraficznej, a mianowicie formacji norowickiej. Formację norowicką stanowi około 50 m miąższości sekwencja jasno-szarych zwięzłych wapieni z wkładkami wapieni dolomitycznych i rzadko margli, spoczywającej na dolomicie głównym (*Hauptdolomit*), a przykrytej przez wapienie krynoidowe liasu dolnego (hetang). Stratotypem wyróżnionej formacji jest profil odsłaniający się na stokach Norowicy w Górach Strażowskich (por. fig. 5—6), zaś hipostatotypami są profile z Doliny Chochołowskiej i okolic Tfstie (por. fig. 3 oraz 9). Nazwa formacji jest nowa i została zaproponowana pierwszy raz w niniejszej pracy.

W obrębie formacji norowickiej wydzielono trzy ogniwa: ogniwo wapieni dolnych, ogniwo wapieni Siwej Wody, oraz ogniwo wapieni mojtińskich. W wapieniach Siwej Wody i wapieniach mojtińskich stwierdzono obecność konodontów oraz szeregu otwornic o znaczeniu stratygraficznym; są to przede wszystkim konodonty *Mistikella posthernsteini* Kozur & Mock, oraz bardzo liczne otwornice *Triasina hantkeni* Majzon. Wymienione mikroskamieniałości określają wiek zawierających je osadów na retyk (por. fig. 10). W tej sytuacji jedynie ogniwo wapieni dolnych może obejmować wiekowo górny noryk (sewat).

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