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A middle Cambrian (Series 3, Stage 5) microfaunal assemblage from the Torgau-Doberlug Syncline (Central Germany) and its palaeogeographic implications for the configuration of West Gondwana

ABUBAKER ATNISHA

TU Bergakademie Freiberg, Geological Institute, Department of Palaeontology, Bernhard-Von-Cotta 2, Freiberg, D-09599, Germany. E-mail: aatnisha1977@gmail.com

ABSTRACT:

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A microfauna of small shelly fossils (SSF) is reported here for the first time from middle Cambrian (Series 3, Stage 5) subsurface strata of the Torgau-Doberlug Syncline (TDS), Central Germany. Considering that this microfauna is strongly limited and poorly preserved the material is quite abundant and diverse. The assemblage consists of molluscs (pelagiellids, bivalves), coeloscleritophorans (chancelloriids, halkieriids), poriferids, protoconodonts, cambroclaves, hyoliths, brachiopods, and disarticulated echinoderm remains. Additionally, a probable pterobranch hemichordate is noted. The assemblage is dominated by epifaunal suspension feeders from mid- to outer shelf depositional settings. Stratigraphically it represents (together with rare trilobites) the oldest middle Cambrian (Series 3, Stage 5) fauna known from Central Germany and the entire Saxothuringian Zone. Regardless the taphonomic problems related to the SSF occurrence, close palaeobiogeographic relations are indicated with the Mediterranean shelf of West Gondwana (especially with the areas of southwestern Europe and Morocco). The reported microfauna coupled with recent trilobite and palynomorph research supports assumptions that the Cambrian succession in the TDS is by far more complete than hitherto suggested, emphasizing its importance as a region yielding Cambrian rocks in Central Europe.

Key words: Small shelly fossils; Cambrian; Torgau-Doberlug Syncline; Germany; West Gondwana.

INTRODUCTION

During the early and middle Cambrian (Terreneuvian to Series 3), skeletal fossils appeared for the first time in great diversity and abundance (e.g., Bengtson 2004; Maloof *et al.* 2010; Kouchinsky *et al.* 2012; Skovsted *et al.* 2014). Many of these small skeletonised microfossils have been informally categorized as 'small shelly fossils' (SSF's; Matthews and Missarzhevsky 1975). Although most of the skeletons

were originally calcareous, others were aragonitic or phosphatic (Landing *et al.* 1989; Bengtson and Conway Morris 1992; Landing 1992; Porter 2010), or diagenetically phosphatised allowing for chemical extraction from the carbonate host rock.

SSF's containing a large number of mineralized metazoan microremains, often of uncertain systematic affinity (Yang 2014), occur from the latest Ediacaran to the Cambrian, and according to Dzik (1994) even until the Ordovician. SSF's are known from Cambrian

successions worldwide and represent important biostratigraphic tools especially for the lower and lower middle Cambrian (Terreneuvian to lower Series 3) (Qian and Bengtson 1989; Elicki 1994; Qian *et al.* 1999; Steiner *et al.* 2007; Parkhaev 2008; Elicki and Gürsu 2009; Vendrasco *et al.* 2009; Parkhaev and Demidenko 2010; Elicki 2011; Yang 2014; Betts *et al.* 2016), even though the viability of various SSF taxa for correlation has been challenged by some authors (e.g., Conway Morris 1988; Landing 1994).

Diverse assemblages have been documented from Cambrian successions worldwide, including: Siberia (e.g., Parkhaev 2004; Kouchinsky et al. 2017), Mongolia (Esakova and Zhegallo 1996), Australia (e.g., Bengtson et al. 1990; Betts et al. 2016), Antarctica (e.g., Wrona 2009), North America (e.g., Landing 1995; Landing et al. 2002; Skovsted and Peel 2010; Wotte and Sundberg 2017), Germany (Sdzuy 1962; Freyer and Suhr 1987, 1992; Elicki and Schneider 1992; Elicki and Debrenne 1993; Elicki 1994, 1996, 1998, 1999b, 2003, 2005, 2007), Spain (Geyer 1986; Fernández-Remolar 2005; Wotte 2009), France (Kerber 1988; Álvaro et al. 2002; Devaere et al. 2013), Sardinia (Elicki and Wotte 2003; Elicki and Pillola 2004), Turkey (Sarmiento et al. 2001; Elicki and Gürsu 2009), Morocco (Geyer 1986; Skovsted et al. 2014), China (e.g., Yang 2014; Chang et al. 2017), Iran (Hamdi et al. 1989), Jordan (Shinaq and Elicki 2007; Elicki 2011), and India (Brasier and Sigh 1987).

The SSF's documented so far from Cambrian Series 2 strata in the Torgau-Doberlug Syncline (Elicki 1994, 2005, 2007) show some relation to the other faunas from Cambrian successions in West Gondwana, which palaeogeographically assign the region to the European shelf of Gondwana (Geyer and Peel 2011; Elicki and Geyer 2013; Geyer *et al.* 2014; Table 1). This paper presents observations on the skeletonised microfossil record from dissolved limestone interbeds and lenses of micropalaeontologically hitherto non-investigated middle Cambrian drill cores of the Torgau-Doberlug Syncline, and discusses their biostratigraphic and palaeogeographic implications.

GEOLOGICAL AND STRATIGRAPHIC BACKGROUND

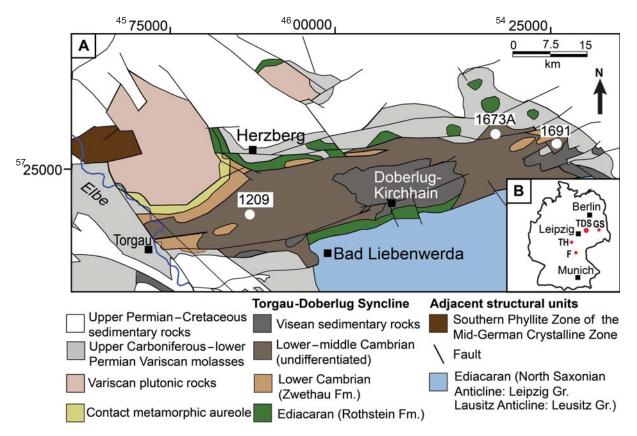
Fossiliferous Cambrian rocks are rare in Germany. They are known from the Görlitz area, the Franconian Forest, Thuringia, and from the Torgau-Doberlug Syncline (TDS) in the vicinity of Leipzig (see Elicki 2007; Geyer *et al.* 2008 and references therein; Textfig. 1). The latter region exhibits the most complete

Cambrian succession (Text-fig. 2). Here, fossiliferous Cambrian Series 2 and 3 sequences containing skeletal fauna and archaeocyathan-calcimicrobial reef mounds have been recognized in drill cores since the 1950s (see Elicki 1997; Geyer *et al.* 2014 for the history of research).

Tectonostratigraphically, the TDS is part of the Saxothuringian Zone of Central Europe (Text-fig. 1), and represents one of the Variscan zones defined by Kossmat (1927). In the TDS, Proterozoic and Palaeozoic strata represent subsurface units known only from boreholes, except for a very localized outcrop of Proterozoic rocks. The Proterozoic to Palaeozoic succession is covered by up to 200 m of Cenozoic sediments (Elicki 2007; Geyer *et al.* 2014). The Ediacaran rocks (Rothstein Formation) are unconformably overlain by Cambrian Series 2 (Zwethau Formation) and Series 3 strata (Tröbitz and Delitzsch formations, known as the Arenzhain Group; Freyer and Suhr 1987; Brause and Elicki 1997; Elicki 1999a; Geyer *et al.* 2014; Text-fig. 2).

Palaeogeographically, the TDS was part of the Mediterranean facies realm of the European shelf of West Gondwana during Epochs 2 and 3 (Text-fig. 3). This Mediterranean facies is characterized by a transition of marine environments from siliciclastic (deeper and shallow shelf areas) to carbonate (archaeocyath-bearing shallow ramps and shelfs) under equatorial to subequatorial conditions in Cambrian Series 2 (e.g., Elicki 1999b; Álvaro et al. 2000; Elicki and Wotte 2003). In Cambrian Series 3, the Mediterranean facies were predominantly siliciclastic with few carbonates deposited at the beginning of the drowning of shelfs and ramps (e.g., Elicki 2001; Elicki and Wotte 2003). Facies changes are complex and not fully understood (Elicki and Wotte 2003). This general sedimentological evolution at the transition between Cambrian Series 2 and 3 produced a rather uniform stratigraphic pattern over most of the Mediterranean region (e.g., Geyer and Landing 1995; Elicki 1997; Elicki and Wotte 2003). According to Elicki and Wotte (2003), the TDS represents the northernmost occurrence of this palaeogeographic facies belt.

The Rothstein Formation represents the oldest unit and consists of very low-grade metasedimentary rocks, dominated by dark-grey to black distal turbidites, intercalating with greywacke and mudstone beds and thick layers of massive chert (Buschmann 1995). The formation is interpreted as a fragment of a Cadomian back-arc basin developed on thinned continental crust, slightly deformed during the Cadomian Orogeny (Buschmann 1995; Linnemann *et al.* 2000, 2007). Fossils are generally lacking except for some



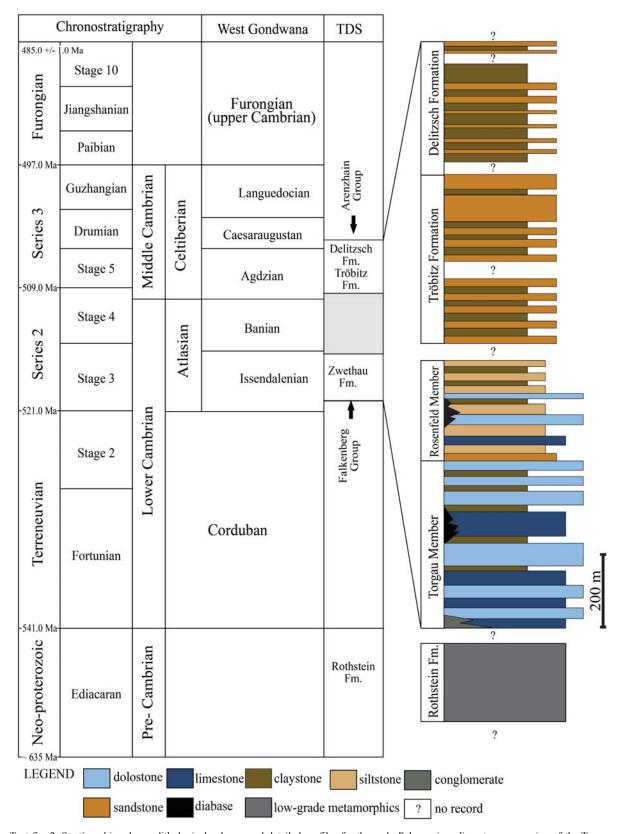
Text-fig. 1. A – Geological map of the Torgau-Doberlug Syncline (TDS) with core localities indicated by white circles (modified from Linnemann *et al.* 2010). B – Sketch map of Germany indicating regions with Cambrian successions (red circles: TDS – Torgau-Doberlug Syncline; GS – Görlitz Syncline; F – Franconian Forest; TH – Thuringia)

non-significant pyritized spherical and filamentous microforms (Buschmann *et al.* 2006). Radiometric dating of an ash layer from the middle part of the formation yielded ages between 570 and 565 Ma, which assigns the formation to the late Ediacaran (Buschmann 1995; Linnemann *et al.* 2000, 2007).

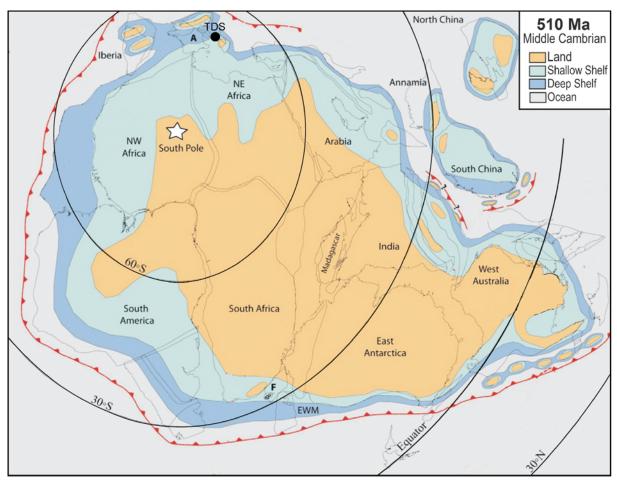
Cambrian deposition started after a stratigraphic and structural gap termed the Cadomian unconformity (Buschmann *et al.* 2006; Linnemann *et al.* 2008; Geyer *et al.* 2014). The succession commences with local conglomeratic debris flow deposits, followed by approximately 700 m of shallow marine carbonates and minor siliciclastics with common calcimicrobial biogenic carbonates that contain archaeocyaths (Zwethau Formation, Cambrian Series 2, Stage 3; Brause 1969; Freyer and Suhr 1987, 1992; Elicki and Debrenne 1993; Buschmann *et al.* 1995; Elicki 1997; Geyer *et al.* 2014; Atnisha *et al.* 2017a; Text-fig. 2).

The Zwethau Formation comprises the Torgau and Rosenfeld members. The Torgau Member consists of approximately 500 m of fossiliferous limestones and

dolostones. Calcimicrobial, oolitic, and intraclastic limestones are common. Cyanobacteria (Epiphyton, Renalcis, Girvanella, Proaulopora, Kordephyton, Botomaella, Subtifloria), archaeocyaths [Retecoscinus aff. guadalquivirensis Perejon, 1976, Neoloculicyathus magnus Debrenne, 1978, Afiacyathus paracompositus (Perejon, 1989) and others], rare trilobites (Dolerolichia pretiosa Sdzuy, 1962), and small shelly fossils (Archiasterella hirundo Bengston in Bengston et al., 1990, A. pentactina Sdzuy, 1962, Allonia tripodophora Doré and Reid, 1965, A. tetrathallis Qian and Bengston, 1989, Cambroclavus sp., Chancelloria sp., Halkieria sp. and Hipponicharion elickii Gozalo and Hinz-Schallreuter, 2002) have been described (Sdzuy 1957a, b, 1962, 1970; Elicki 1994, 1999b, 2007; Geyer et al. 2014). The sedimentary setting of this member is interpreted as a carbonate-dominated subtidal ramp with calcimicrobial-archaeocyathan buildups succeeded by a shallow subtidal to intertidal mixed ramp regime with oolitic shoal complexes and restricted areas (Elicki 1994, 1999b; Buschmann et al. 2006;



Text-fig. 2. Stratigraphic scheme, lithological column, and detailed profiles for the early Palaeozoic sedimentary succession of the Torgau-Doberlug Syncline (TDS) (modified from Elicki 2015)



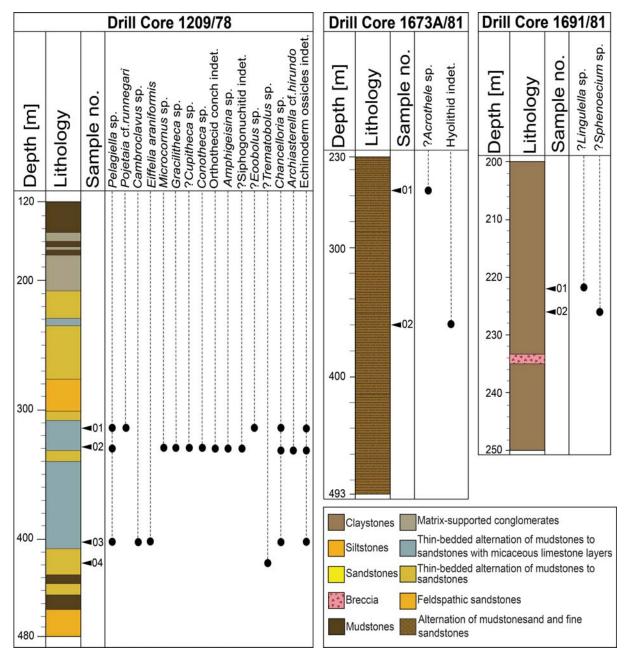
Text-fig. 3. Palaeogeographic map of Gondwana during middle Cambrian times (modified after Torsvik and Cocks 2017). Solid red lines are subduction zones; TDS: Torgau-Doberlug Syncline; A: Armorica; F: Falkland Islands; EWM: Ellsworth-Whitmore Mountains

Atnisha *et al.* 2017a). The biostratigraphic position of the Torgau Member is Cambrian Series 2 (Stage 3) as indicated by archaeocyaths (corresponding to the lower Ovetian of Iberia, which is correlated with the middle Issendalenian of Morocco; see Elicki 1997; Geyer and Landing 2004).

The overlying Rosenfeld Member (up to 280 m thick) is dominated by siliciclastics with occasional carbonate intercalations. The fossil content consists of re-deposited non-determinable archaeocyaths, algae and shelly remains (Elicki 1999a). The depositional environment is poorly constrained and was assigned to a deeper basinal area by Freyer and Suhr (1987), although the occurrence of coarser siliciclastic sediments might rather reflect climatically induced run-off changes under neritic conditions and a palaeogeographic dislocation of the area (Elicki 2003). Moreover, the TDS succession of Cambrian Series 2 (Stage 3) carbonate-dominated strata,

abruptly overlain by Cambrian Series 3 (Stage 5) siliciclastics is identical to other coeval successions in West Gondwana (Landing *et al.* 2015).

The Cambrian Series 3 strata (Arenzhain Group) are represented by siliciclastics with extremely rare intercalations of very thin limestone beds (Brause 1969, 1970; Elicki 1997; Geyer et al. 2014; Atnisha et al. 2017a; Text-fig. 2). The Tröbitz Formation (lower Arenzhain Group) is dominated by alternating quartzitic sandstones and minor dark-grey micaceous claystones, including thin limestone layers near the transition to the overlying Delitzsch Formation (Brause 1970; Geyer et al. 2014; Atnisha et al. 2017a). The fauna is composed of trilobites [Acadoparadoxides brausei Sdzuy, 1970, A. saxonicus Sdzuy, 1970, Cambrunicornia saxonica Geyer, Buschmann and Elicki, 2014, Condylopyge regia Sdzuy, 1970, Paradoxides aff. enormis Sdzuy, 1968, Parasolenopleura lusatica Sdzuy, 1970, Protolenus



Text-fig. 4. Generalized lithostratigraphic columns with the lithology and occurrence of the fauna in the studied drill cores from the TDS.

Location of selected samples indicated by black triangles

(Hupeolenus) bergstroemi Geyer, Buschmann and Elicki, 2014, and others], inarticulate brachiopods, and few hyoliths (Schmidt 1942; Sdzuy 1957a, b, 1970; Elicki 2007; Geyer et al. 2014). The trilobites indicate a stratigraphic level of Cambrian Stage 5 corresponding to the middle to upper Agdzian (lower part of Celtiberian sensu Geyer and Landing 2004 and Buschmann et al. 2006; see Geyer et al. 2014).

The overlying Delitzsch Formation is dominated by quartzitic sandstones alternating with micaceous claystones and by a decrease of the sandstone contribution towards the top. The fauna is similar to that of the Tröbitz Formation and is assigned to Cambrian Series 3. The fauna is composed of trilobites (*Acadoparadoxides brausei* Sdzuy, 1970, *A. cf. insularis* Westergård, 1936, *Acanthomicmacca ano-*

mocaroides Schmidt, 1942, Badulesia tenera Hartt in Dawson, 1868, Bailiella cf. emarginata Linnarsson, 1877, Condylopyge regia Geyer, Buschmann and Elicki, 2014, C. rex Barrande, 1846, Jincella? cf. sulcata Sdzuy, 1968, Paradoxides aff. asturianus Sdzuy, 1968, P. aff. enormis Sdzuy, 1968, Parasolenopleura lusatica Sdzuy, 1970, Peronopsella inaequalis Schmidt, 1942, Solenopleura picardi Schmidt, 1942 and others), helcionelloid molluscs, brachiopods, and few hyoliths (Schmidt 1942; Sdzuy 1957a, b, 1970; Elicki 2007; Geyer et al. 2014). Stratigraphically, two levels can be distinguished. The older level has an early middle Cambrian age, equivalent to the younger fauna of the Tröbitz Formation. The younger level characterizes a relatively narrow stratigraphic interval corresponding to the upper part of Cambrian Series 3 (Stage 5), respectively to the late Agdzian to early Caesaraugustan (Celtiberian) of West Gondwana (Buschmann et al. 2006; Geyer et al. 2014; Atnisha et al. 2017a).

MATERIAL AND METHODS

The material described herein was collected from drill cores nos. 1209/78, 1673A/81, and 1691/81 (Textfigs 1, 4). The drill cores are archived in the central repository of the Brandenburg Geological Survey in Wünsdorf (Landesamt für Bergbau, Geologie und Rohstoffe - LBGR). The location of the studied samples and the reported fauna are shown in Textfig. 4. Samples from core 1209/78 were dissolved in 10% acetic acid. Residues were sieved and dried, and the microfossil specimens were hand-picked. Fossil specimens were mounted, sputter-coated with gold and examined under a scanning electron microscope (SEM). The macro specimens from drill cores 1673A/81 and 1691/81 were coated with magnesium oxide for palaeontological studies. All figured specimens and residues of samples are housed in the Institute of Geology, TU Akademie Freiberg, Germany, under collection number FG 682.

SYSTEMATIC PALAEONTOLOGY

Phylum Mollusca Cuvier, 1797 Class Helcionelloida Peel, 1991 Order Pelagielliformes Parkhaev, 2017 Family Pelagiellidae Parkhaev, 2017 Genus *Pelagiella* Matthew, 1895

TYPE SPECIES: *Cyrtolites atlantoides* Matthew, 1894 from the Terreneuvian, New Brunswick, Canada.

Pelagiella sp. (Pl. 1, Figs 1–6)

MATERIAL: Six internal moulds: FG 682-1209/01, 02, 03, 04, 05 and 06, drill core 1209/78.

REMARKS: The small, dextrally coiled univalve internal moulds have 1-1.5 whorls. The spiral side is nearly flat or slightly depressed, and the umbilical side is narrow. The last whorl is wide. The cross-section is oval to sub-triangular and the protoconch is hookshaped. The aperture is often broken. Pelagiella is characterized by high morphological variation, often with unclear differences (Gravestock et al. 2001; Skovsted 2004; Wotte and Sundberg 2017). According to Dzik and Mazurek (2013), the discovery of specimens of Pelagiella with bunches of setae, and the presence of the mucronate embryonic shell of its early Cambrian relatives make its molluscan nature questionable. The observed TDS specimens are represented by internal moulds without shell material, which makes specific determination difficult. The morphologies of the TDS specimens are shared by several species of Pelagiella.

The studied specimens are internal moulds and an affiliation with *Costipelagiella* Horný, 1964 cannot be excluded. The specimens are not compressed laterally, which excludes them from *C. nevadense* Skovsted, 2006. The discussed material shows some morphological affinity to *P. subangulata* Tate, 1892, which is characterized by a wide range of variability. Due to the poor preservation (steinkerns), taxonomic assignment only in open nomenclature is possible.

OCCURRENCE: *Pelagiella* has been reported from Terreneuvian to Cambrian Series 3 strata worldwide.

Class Bivalvia Linnæus, 1758 Order uncertain Family Fordillidae Pojeta, 1975 Genus *Pojetaia* Jell, 1980

TYPE SPECIES: *Pojetaia runnegari* Jell, 1980 from the Terreneuvian Parara Limestone, South Australia.

Pojetaia cf. runnegari Jell, 1980 (Pl. 1, Fig. 7)

MATERIAL: One figured phosphatic mould: FG 682-1209/07 from core 1209/78.

REMARKS: One internal mould with a length less than 1.2 mm, compressed laterally, equivalved, irreg-

ularly oval in outline. The ligament is usually straight with an umbo, which is central to subcentral, and each valve holds between one or two teeth (Pl. 1, Fig. 7c). The specimens are preserved as steinkerns, therefore growth lines are not observed.

The taxonomic identification of Cambrian bivalves is often problematic because they are small and usually preserved only as internal moulds (Elicki and Gürsu 2009). *Pojetaia runnegari* was first described by Jell (1980), but later Pojeta (2000) suggested that its origin was roughly contemporaneous with that of *Fordilla* Barrande, 1881. *Pojetaia runnegari* and *Fordilla* spp. are the oldest Cambrian bivalves (Elicki and Gürsu 2009; Vendrasco *et al.* 2011).

Specimens of *Pojetaia* and *Fordilla* can be reliably distinguished from one another by the total size (*Fordilla* is larger) and the ratio of anterior length to total length, which is significantly larger in *Pojetaia* (Elicki and Gürsu 2009). The dimensions of the TDS specimen fit those of *Pojetaia* (Elicki and Gürsu 2009) and are significantly different from those of *Fordilla*. The TDS specimen may be referred to *P. runnegari*, the most widespread of the early Cambrian bivalves (Elicki and Gürsu 2009), but poor preservation makes clear designation on a species level somewhat problematic.

OCCURRENCE: The species is known from many Terreneuvian to Cambrian Series 3 successions: Germany (Elicki 1994), Turkey (Elicki and Gürsu 2009), China (Li and Zhou 1986), Australia (Jell 1980), Newfoundland (Skovsted and Peel 2007), Mongolia (Esakova and Zhegallo 1996), Greenland (Skovsted 2006), and Morocco (Geyer and Streng 1998).

Phylum uncertain Class Cambroclavida Conway Morris and Chen, 1991 Family Zhijinitidae Qian, 1978 Genus *Cambroclavus* Mambetov in Mambetov and Repina, 1979

TYPE SPECIES: *Cambroclavus antis* Mambetov in Mambetov and Repina, 1979, Cambrian Series 2, Stage 3, Kazakhstan.

Cambroclavus sp. (Pl. 1, Fig. 8)

MATERIAL: One phosphatized sclerite: FG 682-1209/08 from core 1209/78.

REMARKS: Elongate sclerite. The anterior shield is circular and separated from the elongated poste-

rior region by variable lateral contraction, giving the sclerite a dumbbell shape. The upper surface of the sclerite is rounded and shows a median ridge that originated in the posterior field. The dorsal spine is vertical, elongate, gently recurved and located on the anterior shield, with a more or less circular cross-section. The posterior shield is elongate. The upper surface of the sclerite, including the base, is smooth.

Cambroclaves are a group of enigmatic, phosphatized, hollow, spine-shaped sclerites, known from widespread Terreneuvian to Cambrian Series 3 fossil assemblages (Wotte 2009). The cambroclave specimens described by Conway Morris in Bengtson et al. (1990) have more affinity to a dumbbell shape, which is a fundamental feature of the sclerite described herein, and point to a possible affiliation with C. absonus Conway Morris, 1990 or C. antis. The sclerite shows no distinct furrow around the anterior spine, as figured in Mambetov and Repina (1979, pl. 14, figs 1, 2). No lateral corrugation of the anterior spine is observed, which rules out affiliation of the described sclerite to Cambroclavus ludwigsdorfensis Elicki, 1994, C. clavus Mambetov in Mambetov and Repina, 1979, and C. undulatus Mambetov in Mambetov and Repina, 1979. Due to poor preservation, the specimen is retained in open nomenclature.

OCCURRENCE: Cambroclavids are known from Cambrian Series 2 (Stage 3) to Series 3 (Stage 5) successions from Germany (Elicki and Schneider 1992; Elicki 1994, 2005; Elicki and Wotte 2003), Sardinia (Elicki and Wotte 2003; Elicki 2005, 2006), Spain (Wotte 2009), France (Devaere *et al.* 2013), Australia (Betts *et al.* 2016), China (Bengtson *et al.* 1990; Li *et al.* 2016), and Canada (Landing 1991).

Phylum Porifera Grant, 1836 Class Calcarea Bowerbank, 1864 Order Heteractinidae Hinde, 1888 Family Eiffeliidae Rigby, 1986 Genus *Eiffelia* Walcott, 1920

TYPE SPECIES: *Eiffelia globosa* Walcott, 1920; Burgess Shale Formation, Cambrian Series 3 (Stage 5), British Columbia, Canada.

Eiffelia araniformis (Pl. 1, Fig. 9)

MATERIAL: One isolated phosphatized spicule: FG 682-1209/09 from core 1209/78.

REMARKS: The spicule has a regular six-rayed disk, with the rays diverging at c. $60-80^{\circ}$. All rays are arranged and set at a low angle to the plane of the central disk. The rays are slender with length at $60-90~\mu m$. The central ray is short and occurs on the convex side, subperpendicularly to the plane of the six rays.

Bengtson *et al.* (1990) provided a full discussion of this species. The spicule with a regular six-rayed disk and perpendicular central rays can be compared with Cambrian (Terreneuvian to Series 2) specimens of *E. araniformis* from South Australia (Bengtson *et al.* 1990, fig. 12). Bengtson in Bengtson *et al.* (1990) argued that the species of *Actinoites* Duan, 1984, *Lenastella* Missarzhevsky and Mambetov, 1981, and *Niphadus* Duan, 1984 are synonyms of *E. araniformis*. The morphology of the spicule from the TDS is similar to the specimens described in Bengtson *et al.* (1990) and therefore the spicule is referred to *E. araniformis*.

OCCURRENCE: *Eiffelia araniformis* has been reported from Terreneuvian to Cambrian Series 3 strata worldwide.

Phylum uncertain Class Hyolitha Marek, 1963 Order Hyolithida Sysoev, 1957 Family Nelegerocornidae Meshkova, 1974 Genus *Microcornus* Mambetov, 1972

TYPE SPECIES: *Microcornus parvulus* Mambetov, 1972, Cambrian Series 2, Stage 3, Kazakhstan.

Microcornus sp. (Pl. 2, Figs 1–4)

MATERIAL: Four internal moulds: FG 682-1209/10, 11, 12, and 13 from core 1209/78.

REMARKS: The conchs are incomplete remains of the more-or-less apical area, a slightly curved, tapering cone with curvature toward the dorsum. The apical angle is small and creates a narrow appearance for the conch. The surface of the internal mould is smooth. The cross-section is sub-triangular in shape (Pl. 2, Figs 1a, 2a, 3a, 4a). The dorsal side is more strongly rounded, grading into tightly rounded lateral edges and then into straight flanks; flanks meet at prominent dorsal ridge; dorsum low, giving the conch a depressed rounded triangular cross-section.

The specimens from the TDS are slightly concave

to nearly round on the ventral side, and the subtriangular cross-section indicates their assignment to *Microcornus*. They are similar to *M. eximius* Duan, 1984. *Microcornus petilus* Bengtson in Bengtson *et al.*, 1990 has a large apical angle and generally a finer ornamentation on the shell. These characteristics are not observed in the TDS specimens. They are more rounded compared to *M. breviligulatus* Qian, 1989 and *M. egregius* Demidenko in Gravestock *et al.*, 2001. Generally, according to Dzik (1978), the affiliation of the *Microcornus* specimens seems to be a problem because *Microcornus* should represent hyolithid larval stages. Due to poor preservation, the specimen is retained in open nomenclature.

OCCURRENCE: The genus *Microcornus* is reported from Cambrian Series 2 to Series 3 of Australia (Bengtson *et al.* 1990; Gravestock *et al.* 2001), Siberia (Kouchinsky *et al.* 2015), Mongolia (Missarzhevsky 1981), Greenland (Malinky and Skovsted 2004), North America (Wotte and Sundberg 2017), South China (Yang 2014), Antarctica (Wrona 2003), Germany (Elicki and Schneider 1992; Elicki 2007), Spain (Wotte 2006) and Turkey (Sarmiento *et al.* 2001).

Hyolithid indet. (Pl. 2, Fig. 5)

MATERIAL: One specimen: FG 682-1673A/01 from core 1673A/81.

REMARKS: The conch is preserved in dorsal view as a composite mould. The operculum is poorly preserved and difficult to delimit from the surface of the conch. The conch exhibits an apical angle of about 15–20°; it is 6 mm long and the average width is 2 mm.

The specimen is poorly preserved, which prevents precise taxonomic assignment. Martí Mus and Bergström (2005) described the best preserved hyolithid conchs with articulated helens and an operculum of *Maxilites* Marek, 1972 from the Cambrian Series 3 of the Czech Republic (Peel 2010; Valent *et al.* 2013). The Czech specimens are preserved as moulds and exquisite details of the musculature and internal structures can be observed. Such details are not present in the TDS specimen. The general shape of the TDS specimen is similar to *Haplophrentis* Babcock and Robison, 1988, but due to very poor preservation, the specimen is only generally assigned as a hyolithid remain.

Order Orthothecida Marek, 1966 Family Gracilithecidae Sysoev, 1972 Genus *Gracilitheca* Sysoev, 1968

TYPE SPECIES: *Gracilitheca ternata* Sysoev, 1968, Terreneuvian to Cambrian Series 2, Siberia.

Gracilitheca sp. (Pl. 2, Figs 6–8)

MATERIAL: Three figured internal moulds: FG 682-1209/15–17 from core 1209/78.

REMARKS: Phosphatic or silicified fragments of internal moulds. They are straight, with a triangular cross-section produced by concave ventral and arched dorsal sides. The latter is composed of flattened lateral portions forming a median longitudinal ridge on the top. The ventral side is also delimited by longitudinal ridges. The cross-section acquires a rounded triangular shape. The aperture and initial part are not preserved. The operculum is not preserved.

This specimen is assigned to *Gracilitheca* based on general morphological features and its specific triangular cross-section. The TDS specimens are preserved as internal moulds and retained in open nomenclature as the specimens do not preserve enough details.

OCCURRENCE: *Gracilitheca* is known from Siberia (Sysoev 1968), North America (Landing 1988), Morocco (Marek *et al.* 1997; Malinky and Skovsted 2004), Czech Republic (Valent *et al.* 2013), and Germany (this report).

Family Cupithecidae Duan, 1984 Genus *Cupitheca* Duan in Xing *et al.*, 1984

TYPE SPECIES: *Paragloborilus mirus* He in Qian, 1977, Terreneuvian (upper Fortunian to lower Stage 2); South China.

?*Cupitheca* sp. (Pl. 2, Fig. 9)

MATERIAL: One internal mould: FG 682-1209/18 from core 1209/78.

REMARKS: The specimen is an elongate tube, straight or slightly curved, with the proximal part terminating in a thinner and convex inflated septum-like transverse wall. The cross-section is circu-

lar (Pl. 2, Fig. 9b). The surface of the tube is smooth. The operculum is not observed.

Bengtson in Bengtson *et al.* (1990) used *Actinotheca* Xiao and Zhou, 1984 as a generic name for decollating tubular fossils from the Terreneuvian to Cambrian Series 2. Because the name turned out to be occupied by a genus of tabulate corals *Actinotheca* French, 1889, Demidenko (2001) suggested that it should be replaced by the next valid junior synonym, *Cupitheca* Duan in Xing *et al.*, 1984 (*Cupittheca* of Demidenko 2001). The generic name was spelled "*Cupittheca*" by Duan (in Xing *et al.* 1984, p. 152) but this appears to be a print error (Wrona 2003; Malinky and Skovsted 2004; see Demidenko in Gravestock *et al.* 2001 for a detailed discussion).

The TDS specimen is assigned to Cupitheca because it has features that fit the diagnosis, such as a straight conical tube with an apical termination created by a septum, and a circular cross-section. The conch is similar to Actinotheca mira He in Qian, 1977 from South China (see discussion of the genera Actinotheca and Cupitheca by Bengtson et al. 1990 and Gravestock et al. 2001). According to Bengtson in Bengtson et al. (1990), the surface sculpture is a distinguishing feature of species belonging to Cupitheca, which makes specific differentiation of internal moulds difficult. The TDS specimen represents an internal mould with no external ornamentation, which makes assignment of the specimen to some of the taxa defined by Bengtson in Bengtson et al. (1990) difficult. Nevertheless, the specimen is assigned to *Cupitheca* and left in open nomenclature.

OCCURRENCE: *Cupitheca* is known from Cambrian strata in many localities worldwide: Siberia (Kouchinsky *et al.* 2017), China (Yang 2014), Antarctica (Wrona 2003), Australia (Bengtson *et al.* 1990; Skovsted *et al.* 2016), Greenland (Malinky and Skovsted 2004), Western Newfoundland (Skovsted and Peel 2007), Himalaya (Gilbert *et al.* 2016), Korea (Lee 2008), Spain (Jensen *et al.* 2010), and Germany (this report).

Order Orthothecida Marek, 1966 Family Circothecidae Missarzhevsky in Rozanov *et al.*, 1969

Genus *Conotheca* Missarzhevsky in Rozanov *et al.*, 1969

TYPE SPECIES: *Conotheca mammilata* Missarzhevsky in Rozanov *et al.*, 1969, Terreneuvian to Cambrian Series 2, Siberia.

Conotheca sp. (Pl. 2, Fig. 10)

MATERIAL: One conch: FG 682-1209/19 from core 1209/78.

REMARKS: The specimen is gently curved with an oval to round cross-section (Pl. 2, Fig. 10b). In dorsal view, the apical angle is c. 15–20°. The shell is covered with closely though irregularly spaced transverse lines of varying intensity in the upper part of the conch.

Conotheca is not identifiable to species level when the operculum is not preserved, because the diagnostic morphologic features of the tubular conch are relatively few; they are limited mostly to the nature of ornamentation, apical angle and cross-section. In the TDS specimen, the absence of the operculum makes assignment of the conch to a species difficult. The TDS material is similar to C. petasotheca Landing and Bartowski, 1996, although assignement to C. australiensis Bengtson in Bengtson et al., 1990, and C. laurentiensis Landing and Bartowski, 1996 cannot be excluded. The TDS conch is assigned to Conotheca and left in open nomenclature.

OCCURRENCE: *Conotheca* known from numerous Terreneuvian to Cambrian Series 3 successions: Siberia (Rozanov *et al.* 1969; Kouchinsky *et al.* 2017), Australia (Gravestock *et al.* 2001), China (Yang 2014), Spain (Wotte 2006), Germany (Elicki 2007), France (Devaere *et al.* 2013), Turkey (Sarmiento *et al.* 2001), North America (Landing 1988; Skovsted and Peel 2010), and Greenland (Malinky and Skovsted 2004).

Family and genus uncertain Orthothecid conch indet. (Pl. 2, Fig. 11)

MATERIAL: One internal mould: FG 682-1209/20 from core 1209/78.

REMARKS: The specimen is a gently curved tubular conch with an oval, almost circular cross-section, symmetrical about the plane, in which the tube is curved. The outer surface is smooth and no surface ornamentation is visible. The specimen is preserved as a phosphatic or calcareous internal mould. The conch is provisionally regarded as an orthothecid because of the circular cross-section. Due to poor preservation no further taxonomic details are determinable.

Phylum Brachiopoda Duméril, 1806 Class Lingulata Gorjansky and Popov, 1985 Order Lingulida Waagen, 1885 Superfamily Linguloidea Menke, 1828 Family Eoobolidae Holmer *et al.*, 1996 Genus *Eoobolus* Matthew, 1902

TYPE SPECIES: *Obolus triparilis* Matthew, 1902, Cambrian Series 3, Canada.

?*Eoobolus* sp. (Pl. 3, Fig. 1)

MATERIAL: Fragment of ventral valve: FG 682-1209/23 from core 1209/78.

REMARKS: Incomplete specimen, only the upper shelf is preserved. The fragmentary ventral valve is convex. The pseudointerarea is close to an orthocline, triangular with a prominent deep pedicle groove, widening anteriorly. The ventral visceral area is indistinguishable to faintly discernible.

The brachiopod *Eoobolus* is one of the first representatives of the superfamily Linguloidea in the Terreneuvian (Balthasar 2009). The fragmentary specimen from the TDS resembles *Eoobolus* in some aspects of its general morphology, but not all the features of the species referred to that genus were observed on the specimen due to its incompleteness. The valve fragment described herein is very similar to the material described from Canada (Balthasar 2009), Siberia (Ushatinskaya and Korovnikov 2014), and China (Percival *et al.* 2016). However, the specimens from the TDS are only tentatively assigned to *Eoobolus*? and left in open nomenclature.

OCCURRENCE: *Eoobolus* is known from Cambrian Series 3 successions of Siberia (Ushatinskaya and Korovnikov 2014), Antarctica (Holmer *et al.* 1996), Australia (Betts *et al.* 2016), Greenland (Skovsted and Holmer 2005), Canada (Balthasar 2009), Canada (Matthew 1902), and also known from the Drumian strata of China (Percival *et al.* 2016).

Superfamily Linguloidea Menke, 1828 Family Obolidae King, 1846 Subfamily Obolinae King, 1846 Genus *Lingulella* Salter, 1866

TYPE SPECIES: *Lingula davisii* McCoy, 1851 from the Furongian of Wales, UK.

?Lingulella sp. (Pl. 3, Fig. 2)

MATERIAL: One ventral valve fragment: FG 682-1691/01 from core 1691/81.

REMARKS: The specimen has a rounded elliptical outline and a wide rounded warble. The valve is gently convex and the shell is small (length 5.7 mm and width 3.6 mm), with the maximum width occurring at or slightly posterior to the mid-length of the valve. A growing strip is clear, other sculpture is not recognizable. Exterior ornament on the valve consists of fine, closely spaced growth lines with frequent small growth disruptions. Growth disruptions aligned laterally give the appearance of a faint radial ornament.

Species of *Lingulella* are geographically widespread, morphologically conservative, and stratigraphically long-ranging, which make this genus a catch for all elongate subacuminate oboliids with prominent external concentric growth lines (Brock and Cooper 1993). The general morphology of the studied specimen is as in *Lingulella*. The shell is not sufficiently characterized for a firm generic assignment, and therefore it is left in open nomenclature.

OCCURRENCE: *Lingulella* has been reported from Cambrian to Ordovician deposits worldwide.

Superfamily Acrotheloidea Walcott and Schuchert in Walcott, 1908
Family Acrothelidae Walcott and Schuchert in Walcott, 1908
Genus Acrothele Linnarsson, 1876

TYPE SPECIES: *Acrothele coriacea* Linnarsson, 1876, Cambrian Series 3, Sweden.

?Acrothele sp. (Pl. 3, Fig. 3)

MATERIAL: One dorsal valve: FG 682-1673A/02 from core 1673A/81.

REMARKS: The shell is small and round in outline and the dorsal valve is slightly convex. The brephic shell is rounded and the valve bears two well-pronounced tubercles.

The specimen is poorly preserved, which make taxonomic determination difficult. The specimen from the TDS is similar to *Acrothele yurii* Ushatinskaya and Korovnikov, 2016, because the

umbo of the dorsal valve is near the posterior margin, which distinguishes *A. yurii* from other species of *Acrothele*. In addition, it differs from *A. olenekensis* Pelman and Pereladov, 1986 in having a rounded brephic shell, a small foramen confined to the brephic shell, and the presence of two pairs of tubercles on the brephic dorsal valve. Acrothelids are often determined and described in open nomenclature due to their strong intraspecific variability and incomplete preservation (Ushatinskaya and Korovnikov 2016). Accordingly, the described specimen is assigned to *Acrothele* and left in open nomenclature.

OCCURRENCE: Acrothele is reported from many Cambrian Series 3 successions around the world: Germany (Elicki 2007), Spain (Wotte and Mergl 2007), Czech Republic (Mergl and Šlehoverova 1990), North America (Robison 1964), Australia (Engelbretsen 1996), and Siberia (Ushatinskaya and Korovnikov 2016).

Class Obolellata Williams *et al.*, 1996 Order Obolellida Rowell, 1965 Family Trematobolidae Popov and Holmer, 2000 Genus *Trematobolus* Matthew, 1893

TYPE SPECIES: *Trematobolus insignis* Matthew, 1893, Terreneuvian to Cambrian Series 2, New Brunswick, Canada.

?*Trematobolus* sp. (Pl. 3, Fig. 4)

MATERIAL: One dorsal valve: FG 682-1209/24 from core 1209/78.

REMARKS: The shell is large, dorsi-biconvex, subcircular in outline, and 14 mm wide. The dorsal valve is slightly convex (deformed specimen), widest at mid-length. Lateral margins are slightly curved (Pl. 3, Fig. 4a). The ornamentation consists of fine growth lamellae. The ventral view of the dorsal valve is subcircular in outline, slightly wider than long, with the maximum width occurring at the mid-length (Pl. 3, Fig. 4b). The posterior slope of the valve is steep, with a small, narrowly triangular pseudointerarea. Growth lines are often distinct over the entire surface of the pseudointerarea. Those growth lines are parallel to the anterior border of the pseudointerarea, which bears a pair of tooth-like, anterodorsally directed small nodes. The visceral area is not defined well.

The TDS specimen is similar to *Trematobolus* sp. described from limestones (3D preservation) by Geyer *et al.* (2014). Both the specimen described by Geyer *et al.* (2014) and the specimen described herein were collected from core 1209/78, at the depth of 305–340 m and 400–420 m, respectively. The specimen described herein is strongly deformed (preservation in siltstone), which make precise determination difficult. However, due to the mentioned characteristics and the similarity between the two specimens, I follow the discussion given by Geyer *et al.* (2014) and tentatively assign the studied specimen to *Trematobolus* sp.

OCCURRENCE: *Trematobolus* is known from Spain (Liñán and Mergl 2001), Germany (Geyer *et al.* 2014), Morocco (Geyer and Mergl 1995), and Jordan (Geyer and Mergl 1995).

Phylum Hemichordata Bateson, 1885 Class Pterobranchia Lankester, 1877 Subclass Graptolithina? Bronn, 1849 Family Rhabdopleuridae? Fowler, 1892 Genus *Sphenoecium* Chapman and Thomas, 1936

TYPE SPECIES: *Sphenothallus filicoides* Chapman, 1917, Cambrian Series 3, Heathcote and Monegeeta districts of Victoria, Australia; by subsequent designation of Bulman (1970, p. *V*57).

?Sphenoecium sp. (Pl. 3, Fig. 5)

MATERIAL: One specimen: FG 682-1691/02 from core 1691/81.

REMARKS: The tubarium of the specimen is composed of numerous elongate, slender and parallel-sided tubes. Tubes are preserved as thin flat films that lack obvious evidence of folding or wrinkling, and are distinguished from the surrounding shale by their smoother texture and darker brown colour. The tube form is slightly flexuous; definitive branching is not evident. Individual tubes are c. 4–10 mm long. The tube width gradually increases from approximately 0.1 mm in the proximal area to 0.4 mm at midlength and 0.5 mm at the aperture (Pl. 3, Fig. 5). The apertures are simple, without an obvious lip or ornamentation. Fusellar banding and occasional oblique sutures are not well defined due to poor preservation.

The TDS specimen shares general characteristics with the genus *Sphenoecium*. The genus name

was introduced by Chapman and Thomas (1936) as a replacement for the name *Sphenothallus* Chapman, 1917, recognized as a junior homonym of *Sphenothallus* Hall, 1847. *Fascillitubus* Obut and Sobolevskaya, 1967 can be regarded as a synonym of *Sphenoecium*.

Sphenoecium is a tubarium construction of colonial pterobranchs from short creeping and branching, slowly widening organic tubes; the tubes are formed from fusellar half-rings or full-rings with irregularly developed oblique sutures (Maletz and Steiner 2015). The construction of the tubarium of Sphenoecium is known from chemically isolated fragments of S. johanssoni Bengston and Urbanek, 1986 (see Maletz and Steiner 2015). Individual thecae in the TDS specimen are smaller than in S. mesocambricus Öpik, 1933 and S. discoidalis Chapman and Thomas, 1936. The erect thecal tubes are not widely spaced, which excludes the TDS specimen from S. obuti Durman and Sennikov, 1993. The specimen described herein is most similar to S. wheelerensis Maletz and Steiner, 2015. Due to the poor preservation, the specimen is retained in open nomenclature.

OCCURRENCE: Pterobranch remains are known from all continents, except for Antarctica, and commonly found in the Cambrian Series 3, Stage 5 strata (Maletz and Steiner 2015). They are known from the Cambrian Series 3 to Furongian of Sweden (Öpik 1933), Siberia (Obut 1964, 1974), Australia (Rickards et al. 1990), China (Zhao et al. 2011; Maletz and Steiner 2015), North America (Walcott 1919; LoDuca et al. 2015), Spain (Sdzuy 1974), Czech Republic (Maletz et al. 2005), and Germany (this report).

Phylum uncertain
Class Coeloscleritophora Bengtson and
Missarzhevsky, 1981
Order Chancelloriida Walcott, 1920
Family Chancelloriidae Walcott, 1920
Genus *Chancelloria* Walcott, 1920

TYPE SPECIES: *Chancelloria eros* Walcott, 1920, Burgess Shale Formation, Cambrian Series 3, Canada.

Chancelloria sp. (Pl. 4, Figs 1–4)

MATERIAL: Four figured internal moulds of isolated sclerites: FG 682-1209/25, 26, 27, and 28 from core 1209/78, plus many additional isolated single rays.

REMARKS: Due to the presence of isolated sclerites, poor preservation and broken nature of the specimens, no detailed systematic affiliation can be given. The sclerites have three to five lateral rays (3–5+0). Such forms are relatively rare in *Chancelloria*. Sclerites of 4–5+0 type are similar to those of *Chancelloria eros* Walcott, 1920 (Bengtson *et al.* 1990). The sclerite of 3+0 type (Pl. 4, Fig. 3) may derive from *Allonnia* sp. Generally, the morphology of the TDS material may suggest affinity to the genus *Chancelloria* and is retained in open nomenclature.

OCCURRENCE: *Chancelloria* is extremely common among Cambrian microfaunas worldwide.

Genus Archiasterella Sdzuy, 1969

TYPE SPECIES: *Archiasterella pentactina* Sdzuy, 1969, Terreneuvian to Cambrian Series 2, Cazalla de la Sierra, Ossa Morena, Spain.

Archiasterella cf. hirundo (Bengtson in Bengtson et al., 1990)
(Pl. 4, Fig. 5)

MATERIAL: One single spicule: FG 682-1209/29 from core 1209/78, plus numerous probably related single rays.

REMARKS: The sclerite has 4+0 rays and is distinctly bilaterally symmetrical. The lateral rays typically meet in a sagittal suture. The two lateral rays are slightly recurved upwards away from the basal plane. The vertical ray is strongly curved upward. The posterior ray is long and straight.

The TDS sclerite resembles the well-characterized species A. hirundo (Bengtson et al. 1990). Many Archiasterella species have the ray formulae other than the 4+0 condition. The TDS sclerite has 4+0 rays, therefore comparison with other species is restricted only to species with this condition. Archiasterella tetractina Duan, 1984 has 4+0 rays, but lacks a recurved adaptcal ray suggesting that it may not actually belong within Archiasterella (Randell et al. 2005). Two specimens of Vasil'eva and Sayutina (1988) were illustrated as the new species A. tetractina (non Duan, 1984) that was later renamed as A. tetraspina (Vasil'eva in Vasil'eva and Sayutina, 1993). According to Moore et al. (2014), this material may also not represent Archiasterella. Lee and Lee (1988) described a new species based upon a single incomplete sclerite with a 4+0 condition. This sclerite is too incomplete to determine whether any of the rays is oriented differently from the others, and thus its assignment to *Archiasterella* remains uncertain. However, the configuration and arrangement of rays of the TDS spicule suggest a most probable affiliation with *A. hirundo*. Due to the poor preservation (no surface structures), the TDS spicule is referred to as *Archiasterella* cf. *hirundo*.

OCCURRENCE: *Archiasterella hirundo* is known from Cambrian Series 2 of South Australia, Germany, Spain, Turkey, and southern China (Bengtson *et al.* 1990; Elicki 1994, 2011; Fernández-Remolar 2001; Sarmiento *et al.* 2001).

Phylum Echinodermata Klein, 1734 Class, order and family uncertain Echinoderm ossicles indet. (Pl. 4, Figs 6–15)

MATERIAL: Nine figured specimens: FG 682-1209/30–38 from core 1209/78, plus several indeterminate echinoderm ossicles.

REMARKS: The specimens recovered from the TDS samples show a wide range of morphology of disarticulated plates with a microporous stereome microstructure. The morphology of the thecal plates ranges from platy, irregular brick-like segments to rounded or polygonal in outline. This morphology is typical for eocrinoids and edrioasteroids. However, no assignment to a particular taxon is possible. It is generally difficult to determine the systematic affinity of such disarticulated thecal plates. Also, the value of echinoderm residues preserved as isolated ossicles is generally limited. Only in some cases, it is possible to suggest some wider systematic affinity. The specimens are illustrated here for a complete record of the microfaunal assemblage.

OCCURRENCE: Reported from Cambrian successions worldwide.

Class Polyplacophora De Blainville, 1816 Order Sachitida He, 1980 Family Siphogonuchitidae Qian, 1977 Genus *Siphogonuchites* Qian, 1977

?Siphogonuchitid indet. (Pl. 2, Fig. 13)

MATERIAL: One phosphatic internal mould: FG 682-1209/22 from core 1209/78.

REMARKS: The phosphatic specimen is flattened, elongated, and curved. The upper side carries one prominent longitudinal ridge that continues all the way from the base to the top. The lower side is concave and has a flattened surface. The cross-section is angled to irregularly subtriangular with a rounded corner. The apex is unknown due to poor preservation.

The Family Siphogonuchitidae was introduced by Qian (1977) with the original generic content of *Siphogonuchites*, *Palaeosulcachites* Qian, 1977, *Carinachites* Qian, 1977, *Lopochites* Qian, 1977, with many genera later added to this family. The main feature which excludes the affiliation of the studied specimen to *Protohertzina* and *Amphigeisina* is the shape of the upper side. One prominent longitudinal ridge is present in the studied specimen, whereas in *Protohertzina* the median ridge is bordered by two usually distinct lateral ridges and in *Amphigeisina* it is flat with two lateral ridges from the base to the top.

The specimen shows similarity to the material reported from Cambrian Series 2 of eastern Greenland (Skovsted 2006), and from southern China and northern Iran (Qian and Bengtson 1989; Steiner *et al.* 2004). The preservation of this specimen is too poor for precise determination. Siphogonuchitids are typical elements of the Cambrian Terreneuvian skeletal fauna and are mostly grouped within the class Coeloscleritomorpha Bengtson and Missarzhevsky, 1981.

OCCURRENCE: Siphogonuchitids are known mainly from Asia: China (Steiner *et al.* 2004), Mongolia (Esakova and Zhegallo 1996), Siberia (Khomentovsky *et al.* 1990), but have been also reported from other Cambrian successions: Australia (Bengtson *et al.* 1990), Greenland (Skovsted 2006), France (Devaere *et al.* 2013), Iran (Hamdi 1995), Jordan (Elicki 2011), and for the first time from Germany (this report).

Phylum Chaetognatha? Leuckart, 1854 Class, order, and family uncertain Genus *Amphigeisina* Bengtson, 1976

TYPE SPECIES: *Hertzina? danica* Poulsen, 1966; Terreneuvian to Cambrian Series 3, Bornholm, Denmark.

Amphigeisina sp. (Pl. 2, Fig. 12)

MATERIAL: One sclerite: FG 682-1209/21 from core 1209/78.

REMARKS: The specimen is a long, slender and gently curved, simple, bilaterally symmetrical sclerite. A relatively narrow concave side extends from the tip towards the base. The upper side is broader, flat with two lateral ridges from the base to the top, and opens in the slightly flared basal part. The internal cavity extends to the tip. The wall is probably composed of two layers, but this is not clear because of poor preservation. The surface is smooth. The two sides meet to produce two longitudinal carinae flanked by two thin flanges formed by the outer layer. The flanges reduce gradually towards the tip, where they merge with the carinae.

The specimen described here from the TDS is the first report on the presence of *Amphigeisina* from Germany. The specimen is very similar to *Amphigeisina danica* (Poulsen, 1966) described from Cambrian Series 3 strata in Siberia (Demidenko 2006; Kouchinsky *et al.* 2015). A similar taxon is *Protohertzina* Missarzhevsky, 1973, but the upper side of the latter has a median ridge bordered by two usually distinct lateral ridges, which exclude the affiliation of the TDS specimen to this genus. The TDS specimen is similar to *Amphigeisina danica* but due to the poor preservation it is retained in open nomenclature.

OCCURRENCE: *Amphigeisina* is known from the UK (Hinz 1987), North America (Landing 1974), Denmark (Poulsen 1966; Bengtson 1976), and China (Steiner *et al.* 2007).

DISCUSSION AND CONCLUSIONS

The SSF specimens examined in this contribution were mostly collected from drill core 1209/78 (Text-fig. 4), from very rare and thin carbonate coquina layers that occur occasionally in the lower part of the generally siliciclastic Cambrian Series 3, Stage 5 succession of the TDS. Only a few other taxa are described here from siliciclastic strata of other drill cores (nos. 1673A/81 and 1691/81) that were so far not investigated palaeontologically. The SSF fauna consists of pelagiellid molluscs, bivalves, hyoliths, chancelloriids, poriferids, protoconodonts, cambroclavids, brachiopods, and problematic fossils (siphogonuchitids?), as well as undeterminable echinoderm remains (Table 1; Text-fig. 4). The preservation of the biota is rather poor. Diagenetic overprint led largely

Region	West Gondwana								
Tegion .	Germany			nia	cco		e	ı blic	, X
Таха	This study	TDS	GS	Sardinia	Могоссо	Spain	Franc	Czech Republic	Turkey
Molluscs									
Pelagiella sp.	×		×	×	×	×			
Pojetaia cf. runnegari	×		×		×				×
Cambroclaves									
Cambroclavus sp.	×	×	×	×		×	×		
Porifers									
Eiffelia araniformis	×		×						×
Hyoliths									
Microcornus sp.	×		×			×			×
Hyolithid indet.	×								
Gracilitheca sp.	×				×			×	
?Cupitheca sp.	×					×			
Conotheca sp.	×					×	×		×
Orthothecid conch indet.	×								
Brachiopods									
?Eoobolus sp.	×								
?Lingulella sp.	×	×				×		×	
?Acrothele sp.	×	×				×	×		
?Trematobolus sp.	×	×			×	×			
Hemichordates									
?Sphenoecium sp.	×					×		×	
Chancelloriids									
Chancelloria sp.	×	×	×			×	×		×
Archiasterella cf. hirundo	×	×	×			×	×		×
Miscellaneous taxa									
Echinoderm ossicles indet.	×	×	×	×	×	×	×	×	×
?Siphogonuchitid indet.	×						×		
Amphigeisina sp.	×								

Table 1. Faunal comparisons and distribution of western Gondwana taxa within the described Cambrian fauna from Germany. Sources: see the text for references. Abbreviations: GS – Görlitz Syncline; TDS – Torgau-Doberlug Syncline

to replacement of the original mineralogy (pseudomorphs) and to the subsequent dissolution of the shells. Additionally, there was loss of pseudomorphs in some cases due to the nature of the applied chemical preparation. However, in the cases when the original mineralogy was transformed into phosphate or when steinkerns were formed, micropalaeontological investigation was possible to some extent.

The microfauna collected from drill cores 1673A/81 and 1691/81 are new for the TDS (Text-fig. 4); some of the SSFs reported from drill core 1209/78 have previously been known from this region (Geyer *et al.* 2014), but others represent the first evidence for these forms in the TDS and in Central Europe in general (Text-fig. 4). The trilobite remains described by Geyer *et al.* (2014) from drill core 1209/78 and the SSFs described here were collected from the same interval. The trilobites such as *Protolenus* (*Hupeolenus*)

bergstroemi, Cambrunicornia saxonica, and Ornamentaspis? aff. todraensis Geyer, 1990 correspond to the earliest middle Cambrian (approximately Cambrian Stage 5). The brachiopod Trematobolus with its short stratigraphic range fits very well to this interpretation. As already stated by Geyer et al. (2014), based mainly on the trilobite content, the metazoan fauna from drill core 1209/78 represents the oldest middle Cambrian Series 3 (Stage 5) fauna of the whole Saxothuringian Zone. The SSF assemblage reported here is rather typical for the carbonate environments of Cambrian Series 2 and Series 3 (Stage 5) successions from Sardinia, Morocco, Spain, Czech Republic, and Turkey, which means they are similar over the whole of West Gondwana (Elicki 1994, 2007; Fernández Remolar 1999; Sarmiento et al. 2001).

Palaeoecologically, the shelly fauna is clearly dominated by epifaunal suspension feeders. The

limited grain size of the sediments and the related sedimentological features point to a rather distal depositional environment (mid- to outer shelf). It can be assumed that areas with a higher carbonate content and an enriched shelly fauna existed more proximally. From there material was redeposited or transported to a distal area by occasional storms producing thin coquina layers.

The described SSF fauna is similar to assemblages reported from more-or-less all West Gondwanan successions (Sardinia, Morocco, Spain, Czech Republic and Turkey) and shows a distinct relation to the Mediterranean shelf of West Gondwana (Table 1). Nevertheless, some of the biota have a rather widespread palaeogeographic distribution. Common molluscs (Pelagiella and Pojetaia) occurred along the Gondwanan margins (Elicki 1994, 1996, 2007). Cambroclaves are reported from Sardinia, Australia, Kazakhstan, China, Spain, and Germany (Mambetov and Repina 1979; Bengtson et al. 1990; Elicki and Wotte 2003; Elicki 2007; Wotte 2009). The graptolite Sphenoecium sp. is common from Cambrian Series 3 and known from all continents, except for Antarctica. It has been reported from Central Europe (Czech Republic, Spain; Table 1), and for the first time from Germany (Sdzuy 1974; Maletz et al. 2005; Maletz and Steiner 2015). The most robust palaeogeographic indications come from the trilobites as already published by Geyer et al. (2014). These data point to close connections with other West Gondwana regions (Morocco, Spain).

According to Elicki (2007), the palaeontological data from Central Europe contradicts the hypothesis of deposition of Cambrian strata within separate pullapart basins as formerly assumed. The Cambrian fauna known previously from Germany and the new SSF assemblage reported herein indicate a rather wide geographical distribution in other parts of West Gondwana (Table 1). This wide distribution needs not only open-marine conditions but also long-duration migration paths for faunal exchange between these parts of West Gondwana (Sdzuy 1972; Elicki and Debrenne 1993; Elicki 1994, 2007; Geyer and Elicki 1995; Liñán *et al.* 2004).

It is generally concluded that the fauna described herein from the carbonate layers of the TDS is well known from many Cambrian Series 2 to Series 3 successions deposited in low-latitude environments. Nevertheless, some of the taxa may support a somewhat closer relation to West Gondwana (Morocco, Spain, Sardinia, Czech Republic, Turkey; Table 1). The similarity in the biota and the sedimentary facies of Cambrian successions in West Gondwana indi-

cates that these successions were located quite close to each other during the entire Cambrian at least until the beginning of the Early Ordovician (Sdzuy 1972; Elicki and Debrenne 1993; Geyer and Elicki 1995; Álvaro *et al.* 1998; Sarmiento *et al.* 2001; Liñán *et al.* 2004; Elicki 2007).

General progress and more detailed studies of small shelly fossils and also of palynomorphs (see Atnisha *et al.* 2017b) will allow for a more consistent stratigraphic correlation and palaeogeographic reconstruction of the TDS and other West Gondwana regions.

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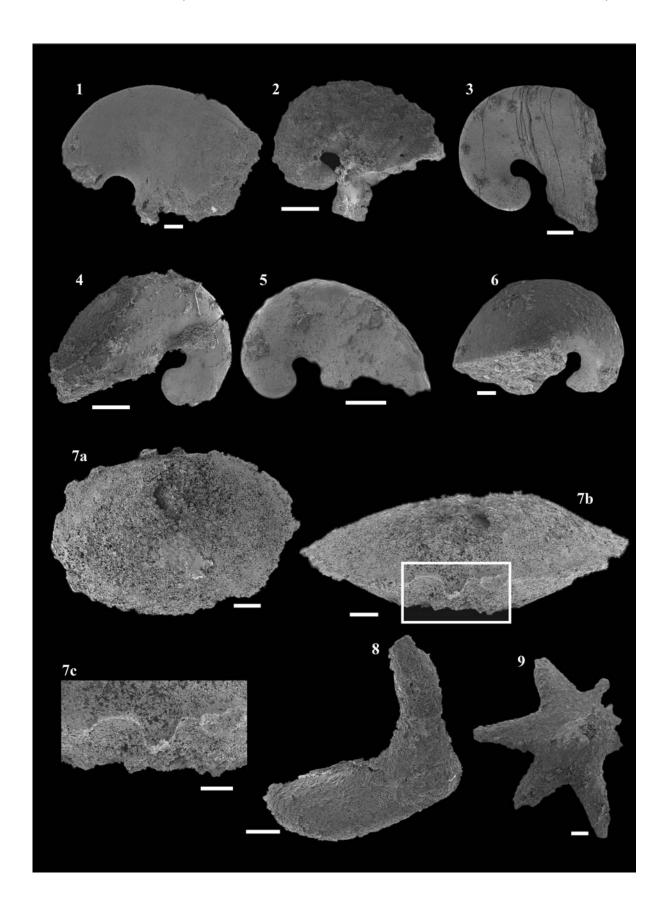
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Molluscan fauna and other small shelly fossils from the Torgau-Doberlug Syncline

1-6 – *Pelagiella* sp. 1 – FG 682-1209/01, 2 – FG 682-1209/02, 3 – FG 682-1209/03, 4 – FG 682-1209/04, 5 – FG 682-1209/05, 6 – FG 682-1209/06; lateral views. **7** – *Pojetaia runnegari* Jell, 1980. FG 682-1209/07. 7a – lateral view, 7b – apical view, 7c – details few showing close up of teeth (White square in 7b). **8** – *Cambroclavus* sp. FG 682-1209/08, upper view. **9** – *Eiffelia araniformis*. FG 682-1209/09. Scale bars 100 μm.



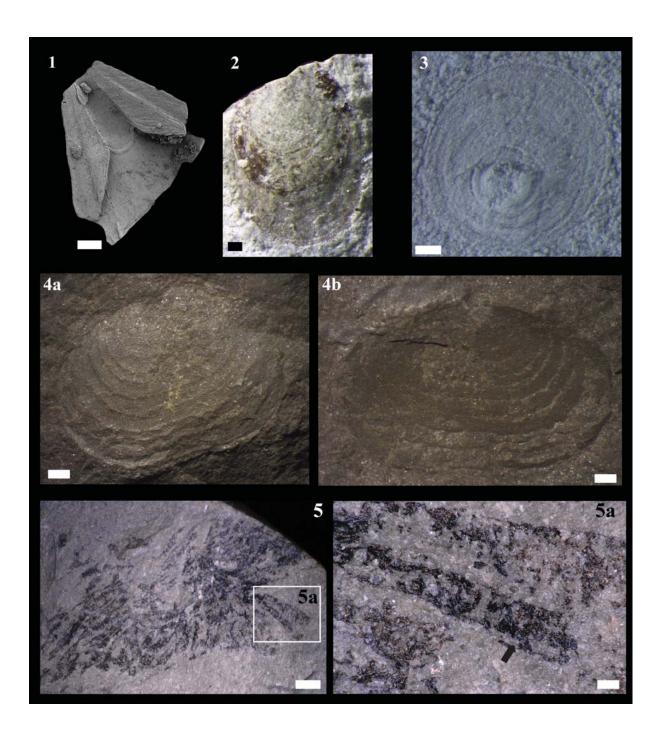
Hyoliths and some problematic taxa from the Torgau-Doberlug Syncline

1-4 – *Microcornus* sp. 1 – FG 682-1209/10, 2 – FG 682-1209/11, 3 – FG 682-1209/12, 4 – FG 682-1209/13. 1a–4a – Cross-sections. **5** – Hyolithid indet. FG 682-1673A/01. Scale bar 500 μm. **6-8** – *Gracilitheca* sp. 6 – FG 682-1209/14, 7 – FG 682-1209/15, 8 – FG 682-1209/16. **9** – *?Cupitheca* sp. FG 682-1209/17; 9a – Lateral view; 9b – Apical view showing the cross-section. **10** – *Conotheca* sp. FG 682-1209/18; 10a – Lateral view; 10b – Apical view showing the cross-section. **11** – Orthothecid conch indet. FG 682-1209/19. **12** – *Amphigeisina* sp. FG 682-1209/20. **13** – ?Siphogonuchitid indet. FG 682-1209/21. Scale bars 100 μm.



Brachiopods and hemichordate from the Torgau-Doberlug Syncline

1-?Eoobolus sp. FG 682-1209/22, ventral view. Scale bar 100 μm. 2-?Lingulella sp. FG 682-1691/01, dorsal view. Scale bar 200 μm. 3-?Acrothele sp. FG 682-1673A/02, dorsal view. Scale bar 200 μm. 4-?Trematobolus sp. FG 682-1209/78-23; 4a – Dorsal view, 4b – Interior view of dorsal valve. Scale bar 1 mm. 5-?Sphenoecium sp. FG 682-1691/02. Scale bar 1 mm; 5a-Close-up of tubes, black arrow indicates fusellar banding. Scale bar 100 μm.



Chancelloriid and indeterminate echinoderm ossicles from the Torgau-Doberlug Syncline

1–4 *– Chancelloria* sp. 1 – FG682-1209/24, 2 – FG 682-1209/25, 3 – FG 682-1209/26, 4 – FG 682-1209/27. **5** *– Archiasterella* cf. *hirundo* (Bengtson in Bengston *et al.*, 1990). FG 682-1209/28. **6-15** – Echinoderm ossicles indet. 6 – FG682-1209/29, 7 – FG 682-1209/30, 8 – FG 682-1209/31, 9 – FG 682-1209/32, 10 – FG 682-1209/33, 11 – FG 682-1209/34, 12 – FG 683–1209/35, 13 – FG 682-1209/36, 14 – FG 682-1209/37, 15 – FG 683–1209/38. Scale bars 100 μm.

