Upper Devonian microvertebrates from the Canning Basin, Western Australia

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ABSTRACT:


A diverse microvertebrate fauna is described from the Virgin Hills and Napier formations, Bugle Gap Limestone Canning Basin, Western Australia. Measured sections at Horse Spring and Casey Falls (Virgin Hills Formation) and South Oscar Range (Napier Formation) comprise proximal to distal slope carbonates ranging in age from the Late Devonian Frasnian to middle Famennian. A total of 18 chondrichthyan taxa are identified based on teeth, including the first record of Thrinacodus tranquillus, Cladoides wildungensis, Protacrodus serra and Lissodus lusavorichi from the Canning Basin. A new species, Diademodus dominicus sp. nov. is also described and provides the first record of this genus outside of Laurussia. In addition, the upper range of Australolepis seddoni has been extended to Late Devonian conodont Zone 11, making it the youngest known occurrence for this species. The Virgin Hills and Napier formations microvertebrate faunas show close affinities to faunas recovered from other areas of Gondwana, including eastern Australia, Iran, Morocco and South China, which is consistent with known conodont and trilobite faunas of the same age.

Key words: Vertebrate palaeontology; Chondrichthyes; Thelodonti; Frasnian; Famennian; Lennard Shelf; Gondwana.

INTRODUCTION

The Canning Basin, Western Australia (Text-fig. 1) is well known for the preservation of Devonian reef complexes as well as both invertebrate and vertebrate fossils (Playford et al. 2009; Klapper 2007; Becker et al. 1993), many of which have proved important in studies of biostratigraphy and correlation. Conodont and ammonoid faunas have both been extensively used in biostratigraphy over the past century with conodont zonations established for the Frasnian (Glenister and Klapper 1966; Klapper 2007; Nicoll and Playford 1993; Metzger 1994) as well as ammonoid zonations for the Frasnian and Famennian (Peterson 1975; Becker et al. 1993; Becker 2000) recognised in the Canning Basin. These fossils have proven useful in determining ages of slope and basin strata enabling successful correlation across physically disconnected localities (Glenister and Klapper 1966; Becker et al. 1993; Becker 2000). However, correlation between slope and platform facies
within the Canning Basin remains problematic as many conodont and ammonoid faunas are absent or undiagnostic in shallow water environments. In contrast, microvertebrates have proven useful in correlating such environments (Turner 1997; Hairapetian et al. 2000; Trinajstic and George 2009). To date, globally correlative microvertebrate zonations are best resolved for the Silurian and early to middle Devonian (Turner 1993; Young 1995; Burrow and Simpson 1995; Valiukevicius 1995; Burrow 1996, 1997; Turner 1997; Burrow and Turner 1998, 2000; Basden et al. 2000; Valiukevicius and Kruchek 2000). Furthermore, a Late Devonian zonation scheme based on pheobodont sharks has been established from the Frasnian to the end Famennian (Ginter and Ivanov 1995a; Ginter 2000; Jones and Turner 2000; Young and Turner 2000). Where known, the Devonian shark fauna of Western Australia (Trinajstic and George 2009; Trinajstic et al. 2014) has been correlated to the Frasnian pheobodont zonation of Ginter and Ivanov (1995a). However, to date little work has been published on Famennian shark taxa from Western Australia. Thelodont zonation has also recently been extended into the Famennian (Hairapetian et al., in press), although the known taxa are currently restricted to Iran and north-western Australia because other thelodonts appear to have become extinct at the end of the Givetian in other areas of the world (Turner 1997; Trinajstic 2001; Märs et al. 2007).

Detailed taxonomic studies on vertebrate faunas of Frasnian-age strata have been undertaken in the

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**Text-fig. 1.** Simplified geological map of the Devonian Reef complexes of the Lennard Shelf, northern Canning Basin, showing the South Oscar Range, Horse Spring and Casey Falls measured sections and main facies types (modified after Playford et al. 2009)
Carnarvon (Turner and Dring 1981; Long 1991; Trinajstic and George 2009) and Canning basins of Western Australia (Long and Trinajstic 2000; Trinajstic and George 2009; Long and Trinajstic 2010) with greatest number of studies in Western Australia done on the Frasnian macrovertebrate faunas of the Gogo Formation (see Trinajstic et al. 2014 for a review). Biostratigraphic studies into vertebrate faunas, from both the Canning and Carnarvon basins however, are less common.

The utility of Western Australian microvertebrates to date strata has been, in part, due to the ability to identify isolated scales through comparison with the exceptionally preserved fauna from the Frasnian Gogo Formation. This has been most successful with placoderms (Trinajstic 1999a), acanthodians (Burrow et al. 2010) and palaeoniscoids (Trinajstic 1999b). However, a single articulated shark (Long and Trinajstic 2010), which is not comparable with any of the isolated teeth recovered so far, has been reported from the Gogo Formation. A second, incomplete specimen comprising meckels cartilage, a shoulder girdle and associated teeth has highlighted the high degree of heterodonty present in Frasnian sharks and the recognition of this variation is important when diagnosing species from isolated teeth. Previous microvertebrate studies in the Canning and Carnarvon basins have revealed the presence of the youngest-recorded thelodont scales, first in early Frasnian strata of the Carnarvon Basin (Turner and Dring 1981) and later, younger scales in the middle Frasnian (Trinajstic and George 2009) and middle Famennian of the Canning Basin (Hairapetian et al., in press).

In contrast to Frasnian vertebrate faunas, Famennian macrovertebrates from Western Australia are rare (Trinajstic et al. 2014), with bothriolepid and phyllolepid placoderms reported from the Willaraddie Sandstone in the Carnarvon Basin (Long and Trinajstic 2000) and coccosteid and dinichthyid placoderms described from the Napier and Virgin Hills formations in the Canning Basin (Long 1987; Trinajstic et al. 2014). Microvertebrate taxa from the Famennian in Canning Basin are even more understudied than those of the Frasnian with only a few reports of dipnoan, acanthodian and chondrichthyan remains from the Gumhole Formation and lower parts of the Yellow Drum Formation (Young 1987; Turner 1993; Edwards 1997; Burrow et al. 2010; Trinajstic et al. 2014). Although rare, thelodont scales as well as the teeth of *Stethacanthus cf. thomasi* and *Thrinacodus ferox*

<table>
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<th>Localities</th>
<th>Casey Falls (CL)</th>
<th>South Oscar (SO)</th>
<th>Horse Spring (VHS)</th>
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Table 1. Distribution and abundances of microvertebrate remains from sections measured at Horse Spring, South Oscar and Casey Falls, Canning Basin, Western Australia.
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Turner, 1982, and a possible Late Famennian otolith have been reported from the Middle to Late Famennian in the Napier Formation (Turner 1993; Trinajstic et al. 2014; Hairapetian et al., in press). Even with a small number of Famennian microvertebrate fossils, the presence of *Thrinacodus ferox* and Famennian thelodonts, indicates a faunal connection with other areas in north Gondwana.

This work readdresses a previous lack of study through the description of a microvertebrate fauna recovered from measured sections (Text-fig. 2; Table 1) that encompass distal slope to platform top facies and crop out along the Lennard Shelf in the Canning Basin of Western Australia. The discontinuous nature of the reef complexes has meant the use of microvertebrates adds a significant control to how the Lennard Shelf carbonate system is correlated and reconstructed. Microvertebrates recovered from sampled strata are compared with the known age ranges reported from other sites globally as well as those species previously described from north-western Australia. In addition to providing the first comprehensive study into Late Famennian chondrichthyans within the Canning Basin, this work also determines faunal links between the Canning Basin and other areas along the margins of northern Gondwana and southern Laurussia.

MATERIALS AND METHODS

Carbonate rock samples (~20 kg each) were processed in a 10% buffered acetic acid solution at Macquarie and Curtin universities (following the methodology of Jeppsson et al. 1999), with resulting residues further separated by either heavy liquid fractionation (Macquarie University) or sieving (0.125 mm sieve; Curtin University) before being picked under a Nikon stereomicroscope. Both conodont and microvertebrate remains were well preserved with the conodont elements indicating a Conodont Alteration Index (CAI) of 1. For SEM photography, specimens were mounted on adhesive carbon tape fixed to 10 mm diameter aluminium stubs and coated with 5 µm of platinum. Specimens were imaged using a Zeiss Evo 40XVP SEM at the Centre for Materials Research at Curtin University as well as a Hitachi TM-3030 desktop SEM at Applied Geology at Curtin University with accelerating voltages ranging from 5–15 kV and under variable pressure.

The 13-fold Montagne Noire (MN) conodont Zonation (Klapper 1989, 2007) modified by Girard et al. (2005) was used to determine the age ranges of the associated Frasnian microvertebrates recovered in this study, as it provides greater resolution than that of Zeigler and Sandberg (1990). The standard conodont Zonation (Zeigler and Sandberg 1990) is used in the Famennian.

Microvertebrate specimens are housed at the Western Australian Museum (WAM).

GEOLOGICAL SETTING

Throughout the Devonian, the Canning Basin (Text-fig. 1) occupied an equatorial position, approximately 12–14° south of the equator, along the northern margins of Gondwana (Scotese and McKerrow 1990). Development of the basin was initiated during the Ordovician. Following a period of uplift and erosion in the early Devonian, extension during the middle Devonian to early Carboniferous led to rapid subsidence of the NW trending Fitzroy Trough along the northern margin of the Canning Basin (Drummond et al. 1991). The margin successions include the late Givetian to late Famennian reef complexes, which are well exposed along the inner Lennard Shelf, and include numerous well-preserved fossilised fringing reefs and atolls (Playford 1980). Difficulty in correlating Frasnian and Famennian sections within the Lennard Shelf is partly due to a complex underlying topography on which the reef was established, as well as depositional heterogeneity (Playford et al. 2009). Tectonic activity during the period of reef building (Chow et al. 2013) and deformation following post-depositional exhumation have also added to a fragmentary Frasnian and Famennian record across the basin (Playford et al. 2009).

**Studied sections**

**Casey Falls**

A section was measured at Casey Falls (18°44′0″ S, 126°05′8″ E; Text-figs 1 and 2) approximately 80 km south east of Fitzroy Crossing (Text-fig. 1). The measured section represents 420 m of toe-of-slope to upper slope carbonate sediments of the Virgin Hills Formation and overlying Bugle Gap Limestone, respectively (Table 2). The lower 100 m of the section mainly comprises reworked silty skeletal to non-skeletal wackestone/packstones derived from platform, margin and slope environments. The upper 320 m consists mostly of stromatoidal microbial boundstones and microbially stabilised packstones and grainstones (Playton et al. 2013). Minor platform derived skeletal-peloidal pack-
stones and grainstones occur in the upper 40 m of the section. The Frasian strata within the section are limited to the first 2 m of the section and range from Conodont zones (CZ) 13a to 13b, with zone 13c not resolvable. The Frasian–Famennian boundary is located between 2.2 and 7.9 m above the section base, however finer biostratigraphic resolution was impossible at this level due to a paucity of conodonts. Overlying Famennian sediments, from 7.9 m, yield conodont zones from Late triangularis to Late marginifera CZ. Addi-

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<td>BUGLE GAP LIMESTONE</td>
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| VIRGIN HILLS FORMATION | marginal-slope and basin facies |
| GOGO FORMATION | basin facies |
| SADLER LIMESTONE | marginal-slope facies |

Table 2. Correlation table of Upper Devonian and Lower Carboniferous units on the Lennard Shelf (after Playford et al. 2009). The discontinuous nature of the reef complexes coupled with the large geographic range of the outcrops resulted in contemporaneous facies receiving different formation names. Dashed lines represent chronostratigraphic boundaries; bold line represents the cessation of reef building in the Canning Basin.
ational samples were taken from poorly outcropping beds approximately 580 m north (samples 1984–95 and 1984–96) and 600 m (samples 1984–97 and 1984–98) north-north-east from the top of the measured section at Casey Falls with conodonts indicating a Late *marginifera* age. These samples represent the shallow water carbonate derived material of the Bugle Gap Limestone.

![Text-fig. 2. Simplified stratigraphic columns of the sections at Horse Spring, Casey Falls and South Oscar Range, showing main facies types and locations of recovered microvertebrate remains (modified after Playton *et al.* 2013)]
South Oscar Range

The section measured at the southern end of South Oscar Range (17°54′53″ S, 125°17′56″ E, Text-figs 1 and 2) spans 585 m of the Napier Formation and represents the seaward side of an offshore island and fringing reef complex (Playford et al. 2009). The preserved sequence comprises transported slope facies dominated by platform derived packstones and grainstones/rudstones with commonly occurring peloids, coated grains and skeletal fragments. Debris deposits consisting of allochthonous blocks and megabreccias of reefal margin material can occur locally and are concentrated in particular parts of the section, reflecting brittle failure of the early-lithified bound margin. Bioclasts are abundant throughout the entire section and dominated by branching and laminar stromatoporoids (in the Frasnian; George 1999; Stephens and Sumner 2003), crinoids, corals and brachiopods. The Frasnian beds range from MN 6 to 13b CZ and in addition to conodonts, yielded chondrichthyan teeth and scales as well as acanthodian scales. The Frasnian–Famennian boundary was located between 228.7 and 233 m above the section base. A diverse Frasnian microvertebrate fauna has also been recorded by Trinajstic and George (2009) with scales of the thelodont Australolepis seddoni recorded as occurring with conodont elements and phoebodont teeth for the first time, thus allowing the age of A. seddoni to be constrained to MN 4–10 CZ. The phoebodont teeth were also correlated to the known phoebodont based zonation of Ginter and Ivanov (1995a).

SYSTEMATIC PALAEONTOLOGY

Class Thelodonti Jaekel, 1911
Order Thelodontiformes Kiaer, 1932
Family Turiniidae Obruchev, 1964
Genus Australolepis Turner and Dring, 1981


Australolepis seddoni (Turner and Dring, 1981) (Text-fig. 3A)

1969. Fish tooth type b; Seddon, p 30, fig. 2a–b.
1981. Nikoliiviid gen. et sp. indet. Turner and Dring: 46, fig. 6A–C.
2013. Australolepis seddoni Turner and Dring, 1981; Chow et al., pl. 1C.

MATERIAL: One broken scale from the Virgin Hills Formation, Horse Spring, sample VHS-315.

DESCRIPTION: A damaged scale, less than 1 mm in length, with part of the base and neck not preserved. Unornamented, conically shaped crown possessing six primary ribs radiating from the apex, with the two anterior ribs bifurcating towards the crown base (Text-fig. 3A). The posterior of the scale is complete and preserves a shallow neck separating the crown from the elliptical base that has a well-developed pulp canal surrounded by tubercular swellings.
REMARKS: Scales attributed to *A. seddoni* are found along the northern margins of Gondwana (Yazdi and Turner 2000; Turner et al. 2002; Märss et al. 2007; Hairapetian et al. 2006; Trinajstic and George 2009) and their presence in mainly shallow-water facies indicate that they inhabited a near-shore, marine to marginal marine environment (Burrow 1997; Turner 1999; Märss et al. 2007). Within the Canning Basin, the majority of scales (46) have been recovered from the Virgin Hills Formation at Horse Spring (Trinajstic and George 2009; this work). This locality represents a distal slope environment formed in depths in excess of 200 m (Playford et al. 2009). This is in contrast to the high numbers (730) recovered from the shallow marine environment of the contemporaneous, Gneudna Formation, Carnarvon Basin, Western Australia (Turner and Dring 1981; Trinajstic 2001) and suggests that the scales were transported downslope to deeper water prior to fossilization. This transport and likely consequential abrasion might have contributed to the lack of fine ornament, diagnostic for the taxon. However, as the co-occurring phoebodont teeth preserve the delicate lateral carinae and striations on the cusps, this explanation no longer seems likely for all scales. Another possibility for the absence of ornament is intraspecific variation, with both ornamented and non-ornamented scales present in shallow water facies at Hull Range (Chow et al. 2013).

DISTRIBUTION AND STRATIGRAPHIC RANGE: Associated conodont elements extend the upper-known age range of *A. seddoni* in the Canning Basin from MN 10 CZ to MN 11 CZ. As the upper age of the Gneudna Formation type section is constrained by the remains of *A. seddoni*, the youngest age for the Gneudna may now be extended to MN 11 CZ. Outside Australia scales of *A. seddoni* have been reported in the Shishtu Formation, Shotori Range, eastern Iran and Chahriseh section, Esfahan, Central Iran from *falsiovalis* to *rhenana* CZ (= MN 1–10 CZ) (Yazdi and Turner 2000; Turner et al. 2002; Hairapetian et al. 2006).

**Class** Chondrichthyes Huxley, 1880  
**Subclass** Elasmobranchii Bonaparte, 1838  
**Order** Phoebodontiformes Ginter, Hairapetian and Klug, 2002  
**Family** Phoebodontidae Williams, 1985  
**Genus** Phoebodus St. John and Worthen, 1875

**Type Species:** *Phoebodus sophiae* St. John and Worthen, 1875

*Phoebodus bifurcatus* Ginter and Ivanov, 1992  
(Text-fig. 3B–E)

1995. *Phoebodus bifurcatus* Ginter and Ivanov, 1992; Ginter, p. 61, fig. 1F–G.
2009. *Phoebodus bifurcatus* Ginter and Ivanov, 1992; Trinajstic and George, p. 648, pl. 1, fig. 9.

MATERIAL: Six teeth from the Virgin Hills Formation, Horse Spring, samples VHS-308, VHS-310, VHS-311, VHS-315.

DESCRIPTION: Teeth with three linguually inclined cusps bearing a thin, lateral carinae extending from the bases of the cusps to the apices and lacking a distinct neck between the base and crown (Text-fig. 3B–E). The cusps are generally of nearly equal size however the size of the central cusp is significantly reduced in one specimen (WAM 14.8.2, Text-fig. 3D). The base is characterised by a labially directed semicircular arch which defines the lingual border and varies between specimens from strongly (Text-fig. 3C) to weakly bifurcated (Text-fig. 3E). A distinct button is lacking on most specimens, although a rounded thickening, close to the lingual rim is sometimes present (Text-fig. 3E). A large foramina is located in the centre of the lingual arch and in some specimens there are smaller adjacent canal openings.

REMARKS: The teeth typical of *Phoebodus bifurcatus* Ginter and Ivanov, 1992 are characterised by having five ornamented cusps and a bifurcating base with a distinct button surrounded by foramina (Ginter and Ivanov 1992). However, tricuspid forms, lacking intermediate cusplets, are known from the Confusion Range, Utah, USA and were attributed to juveniles (Turner and Youngquist 1995). Within the Canning Basin only tricuspid forms of *Ph. bifurcatus* have been identified but were previously attributed to juvenile sharks based on deeply bifurcated bases, an indistinct button and lack of ornament on the cusps (Trinajstic and George 2009,
pl. 1, fig. 9). The teeth attributed to juvenile individuals have been found in association with adult forms in the South Urals, Holy Cross Mountains and central Iran (Ginter and Ivanov 1992; Hairapetian and Ginter 2010) whereas in the Confusion Range, Utah, USA and the Canning Basin only juvenile teeth have been reported. Here we question the assignment of these tricuspid teeth to juvenile sharks as the size of some deeply bifurcated, tricuspid teeth (Text-fig. 3C–D, measuring up to 1.2 mm from furthest point on the lingual margin of the base to the labial edge), recovered from Horse Spring, are of comparable or larger size than other teeth attributed to adult forms previously reported (Ginter and Ivanov 1992, fig. 4B–F)). It is also unlikely that these larger teeth are symphyseal due to the asymmetry of the base and crown on one specimen (WAM 14.7.4, Text-fig. 3C). This would leave the possibility that the Canning Basin teeth, and potentially other smaller tricuspid teeth, represent sexual dimorphism (e.g. Peyer 1968; Hairapetian and Ginter 2010) to 1.2 mm from furthest point on the lingual margin of the base to the labial edge, recovered from Horse Spring, are of comparable or larger size than other teeth attributed to adult forms previously reported (Ginter and Ivanov 1992, fig. 4B–F)). It is also unlikely that these larger teeth are symphyseal due to the asymmetry of the base and crown on one specimen (WAM 14.7.4, Text-fig. 3C). This would leave the possibility that the Canning Basin teeth, and potentially other smaller tricuspid teeth, represent sexual dimorphism (e.g. Peyer 1968; Hairapetian and Ginter 2010), a sub species of Ph. bifurcatus, intraspecific variation or even a separate, closely related species.

DISTRIBUTION AND STRATIGRAPHIC RANGE: In the Canning Basin, Australia Ph. bifurcatus has only been recorded from MN 11 CZ; within the Holy Cross Mountains, Poland, South Urals, Russia, southern China and Utah, USA Ph. bifurcatus ranges from MN 11–12 CZ; in Kale Sardar, eastern Iran the range extends from the rhenana–lower linguiformis Zones (MN 11–13b CZ).

*Phoebodus fastigatus* Ginter and Ivanov, 1992

(Text-fig. 3D–F)


1990. *Phoebodus limpidus* sp. nov.; Ginter, p. 75–76, pl. 4, fig. 6a–b.

1992. *Phoebodus fastigatus* sp. nov.: Ginter and Ivanov, p. 66–67, fig. 3A–B, G.

1993. *Phoebodus* sp.: Liszkowski and Racki, fig. 3F, H, K.

1995. *Phoebodus fastigatus* Ginter and Ivanov, 1992; Ginter, p. 59, fig. 1C–D.


2000. *Phoebodus fastigatus* Ginter and Ivanov, 1992; Ginter and Ivanov, p. 327, pl. 1, fig. E.

2003. *Phoebodus fastigatus* Ginter and Ivanov, 1992; Abousalam, pl. 27, figs 13–16.

2004. *Phoebodus fastigatus* Ginter and Ivanov, 1992; Ginter, fig. 2H–J.


2007. *Phoebodus fastigatus* Ginter and Ivanov, 1992; Liao et al., p. 173, fig. 3A–M.

2008. *Phoebodus fastigatus* Ginter and Ivanov, 1992; Ginter et al., p. 170, text-fig. 2E–H.


MATERIAL: Fourteen teeth from Virgin Hills Formation, Horse Spring, samples VHS-310, VHS-311, VHS-312 and VHS-315.

DESCRIPTION: Teeth with three main cusps. The medial cusp is approximately a third of the length of the lateral cusps and one half to one third the size at the base of the lateral cusps (Text-fig. 3F–H). When present, the accessory cusps are reduced in size, approximately half the size of the central cusp. The lateral cusps are long and slender with some specimens bearing slight torsion (Text-fig. 3F). The central and lateral cusps are all rounded in cross section (Text-fig. 3G), however, the labial faces of the cusps vary between teeth from smooth (Text-fig. 3H) to ornamented with faint to strongly developed striations (Text-fig. 3F). The base is rectangular to trapezoidal in outline, extending further mesio-distally than labio-lingually and bearing a straight to slightly convex thin lingual edge. The lingual face of the base is perforated by numerous small foramina with a round to slightly ovoid centrally located button (Text-fig. 3F).

REMARKS: The teeth are attributed to *Phoebodus fastigatus* Ginter and Ivanov, 1992 based on the thin elongate lingual base with a straight lingual margin and centrally located ovoid button (Ginter and Ivanov 1992). The tooth crowns vary from the diagnostic thin elongate recurved lateral cusps (Text-fig. 3F–G) to shorter more robust cusps with less distal divergence (Text-fig. 3H). A smooth labial cusp face is described as a diagnostic feature of *Ph. fastigatus* (Ginter and Ivanov 1992, 2000) and one of the features distinguishing it from *Ph. bifurcatus*. This feature is not present in all teeth from the Canning Basin with faint to coarse cristae observed here and on teeth previously described by Trinajstic and George (2009, pl. 1, figs 11–16). As there is no correlation between tooth size and the presence of or lack of ornament, these differences are attributed to intraspecific variation rather than ontogeny.
DISTRIBUTION AND STRATIGRAPHIC RANGE: This species ranges from MN 6–11 CZ in the Canning Basin, Western Australia. Worldwide the range is greater, from the Givetian to Frasnian rhenana CZ in the USA, Morocco, Mauritania, Spain, Poland, China, Russia and eastern Australia.

Phoebodus cf. fastigatus (Ginter and Ivanov 1992) (Text-fig. 3I–J)

MATERIAL: Three teeth from the Virgin Hills Formation, Horse Spring, sample VHS-312.

DESCRIPTION: Three small teeth, measuring less than 0.5 mm mesiodistally across the base, with damage to both the cusps and bases (Text-fig. 3I–J). Crowns with three to five lingually inclined cusps, comprising a main central cusp, two lateral cusps and, when present, smaller intermediate cusplets (Text-fig. 3J). The lateral cusps, when preserved, are slightly larger in size than the medial cusp and in one specimen show torsion towards the apex of one lateral cusp (Text-fig. 3I). When present, the intermediate cusps are small and fused to the base of the mesial margin of the lateral cusps. Prominent striations are present on the labial faces of the cusps, whereas the lingual faces are smooth. The base forms a roughly rectangular to trapezoid outline and is perforated by numerous small foramina. A rounded oval button is often difficult to determine on most specimens, however a faint outline can be seen, positioned centrally on one specimen (Text-fig. 3I).

REMARKS: The teeth described here are too poorly preserved for a definitive diagnosis, however they share similarities to the teeth attributed to the ornamented forms of Ph. fastigatus in this work. When preserved, the lingual margin of the base in Ph. cf. fastigatus is thin; a diagnostic feature of Ph. fastigatus (Ginter and Ivanov 1992), however, the cusps of Ph. cf. fastigatus are ornamented with strong striations, ovoid in cross section and lack a distinct rounded button. The bases are not well enough preserved for an outline to be determined. Despite poor preservation, in one specimen the base appears lingually narrow (Text-fig. 3I). The teeth described here are consistently smaller than other teeth attributed to Ph. fastigatus and this may be evidence of ontogenetic variation. Furthermore the teeth commonly lack a distinct button, a feature attributed to juvenile forms in both Phoebodus gothicus Ginter, 1990 and Phoebodus bifurcatus (Ginter and Ivanov 1992). One tooth (WAM 14.7.10, Text-fig. 3G) may represent an intermediate form between teeth designated Ph. fastigatus and the smaller teeth of Ph. cf. fastigatus as it possesses an elongate lateral cusp and more prominent ovoid button similar to Ph. fastigatus, but exhibits the coarse cristae and smaller size of Ph. cf. fastigatus. Therefore an ontogenetic series is supported by the presence of this transitional form.

Phoebodus latus Ginter and Ivanov, 1995a (Text-fig. 3K–O)

1993. Phoebodus sp. Liszkowski and Racki, p. fig. 5L–M.
1995a. Phoebodus latus sp. nov.; Ginter and Ivanov, p. 355, pl. 1, figs 3, 4.
1995b. Phoebodus latus Ginter and Ivanov, 1995b; Ginter and Ivanov, p. 19, fig. 1.
1995. Phoebodus latus Ginter and Ivanov, 1995b; Ginter, fig. 1E.
2009. Phoebodus latus Ginter and Ivanov, 1995b; Trinajstic and George, pl. 1, figs 18, 19.
2011. Phoebodus latus Ginter and Ivanov, 1995b; Ivanov and Lucas, fig. 2A, B.

MATERIAL: Eight teeth from the Virgin Hills Formation, Horse Spring, samples VHS-310 VHS-311 and VHS-312.

DESCRIPTION: The teeth here attributed to Phoebodus latus Ginter and Ivanov, 1995a are highly variable with three different morphotypes identified here. The first morphotype (Text-fig. 3K–L) comprises five to seven smooth and almost straight conically shaped cusps with well-developed lateral carinas. The cusps diverge from the centre of the crown, which is defined by a central cusp with a base approximately 25% smaller than the base of the two lateral cusps (Text-fig. 3K–L). A pair of intermediate cusplets, slightly smaller than the medial cusp, are present on all specimens. One specimen (WAM 14.7.7, Text-fig. 3L) exhibits a second pair of small broken cusplets fused at the base of the lateral cusps. The base is roughly trapezoidal in outline, thickened along the lingual edge and slightly arched. A faint outline of an ovoid button is preserved.

A single, well preserved tooth represents the second morphotype (Text-fig. 3M). The crown comprises three main cusps of almost equal size with a pair of intermediate cusplets approximately a third of the length of the lateral cusps. The cusps are ovoid in cross section and all lingually directed, with the medial and lateral cusps extending beyond the lingual margin of the base (Text-fig. 3M). The cusps exhibit a well-defined lateral
carina as well as striations on the labial faces. The lingual faces of the cusps are smooth. The base forms a roughly rectangular outline with the very faint outline of a centrally located ovoid button. Morphotype three (Text-fig. 3N–O) is characterised by three main cusps and two pairs of intermediate cusplets, all of which are smooth with well-developed lateral carinas and a slight lingual inclination. The size of the central cusp ranges from small (Text-fig. 3O), with the basal width approximately half of the size of the first intermediate cusplets, to significantly larger (Text-fig. 3N), almost the same size as the lateral cusps. The second pair are approximately half the size of the intermediate cusplets. The base is arched and slightly bifurcated (Text-fig. 3N) along the lingual rim which is also perforated by a horizontal row of foramina. The mesio–distally elongate button is positioned toward the edge of the lingual rim of the base (Text-fig. 3N).

REMARKS: The crown morphology of *Phoebodus latus* Ginter and Ivanov, 1995a is highly variable with the presence of intermediate cusplets greater than other known species attributed to *Phoebodus* (Ginter 2008). The third morphotype (Text-fig. 3O–O) resembles other *Ph. latus* teeth recovered from the South Urals, Russia and the Holy Cross Mountains, Poland (Ginter and Ivanov 1992, fig. 7A). The tooth also bears strong resemblance to another tooth from Horse Spring, previously described by Trinajstic and George (2009, pl. 1. figs 18, 19). The first and third morphotypes identified here lack the ornament on the labial faces of the cusps (Text-figs 3K, O) seen in other examples of *Ph. latus* (Ginter and Ivanov 1995a). The crown of the second morphotype more closely resembles other teeth from *Ph. latus* (Text-fig. 3M) with the presence of three almost equally sized main cusps ornamented with ridges on the labial faces. However, the base does not show the diagnostically distinct ovoid button positioned close to the lingual rim (Text-fig. 3M). There appears to be a great deal of diversity attributed to *Ph. latus* that would suggest a high degree of heterodonty. Until more specimens are found, it is difficult to determine whether the teeth attributed to *Ph. latus* in the Canning Basin are products of heterodonty or if they represent intra-specific variation.

DISTRIBUTION AND STRATIGRAPHIC RANGE: Within the Canning Basin, the range of this species is confined to MN 9–11 CZ. In other regions, the species is longer ranging, from the falsiovalis to linguiformis conodont Zones (MN 1–13b) in Poland, the Middle and South Urals, and Timan of Russia.

**Phoebodus sp. 1**  
(Text-fig. 3P)

MATERIAL: One tooth from the Virgin Hills Formation, Horse Spring, sample VHS-312.

DESCRIPTION: A robust tooth comprising two lingually inclined, almost equally sized laterally diverging cusps, which are rounded in cross section and bear a single faint lateral carina on the inner face (Text-fig. 3P1–2). There is no evidence of a third cusp forming. The base is asymmetric and perforated by numerous pores concentrated along the basal rim (3P2). The base is thickest along the lingual rim, forming a rounded bulbous edge. A well-defined circular button is located centrally on the base and reaches almost to the lingual edge. A rounded, striated labiobasal thickening extends from the base (Text-fig. 3P1).

REMARKS: The thin, smooth cusps, which are rounded in cross section, and the centrally located spherical button are similar to the morphological features that diagnose *Ph. fastigatus* (Ginter and Ivanov 1992, fig. 3A–B). In addition, the tooth shares a similar cusp morphology to teeth from the same horizon, attributed to *Ph. fastigatus* (Text-fig. 3H). However, the bicuspid crown, asymmetric outline of the base and thickened lingual rim makes an accurate diagnosis problematic. It is suggested that these differences are pathological, and that the tooth is most likely from *Ph. fastigatus*. However, as less than 1% of chondrichthyan teeth, both fossil and extant (Becker et al. 2000) are known to exhibit pathology the diagnosis of *Ph. fastigatus* cannot be confirmed and until further teeth are recovered this tooth is placed in open nomenclature.

**Genus Thrinacodus** St. John and Worthen, 1875

**TYPE SPECIES:** *Thrinacodus nanus* St. John and Worthen, 1975

**Thrinacodus tranquillus** Ginter, 2000  
(Text-fig. 4A–C)

2000. *Thrinacodus tranquillus* sp. nov.; Ginter, p. 374–377, figs 2a–c, 3a–f, 4a–c, 5h–k.


2002. *Thrinacodus tranquillus* Ginter, 2000; Ginter et al., p. 186–188, text-fig. 9f–h, pl. 2, fig. h, pl. 3, fig. h, pl. 11, figs h–I.
REFERENCES: These teeth largely conform to the diagnosis of *Thrinacodus tranquillus* Ginter, 2000, and the differences between the tooth forms are attributed to heterodonty. Previous work (Ginter 2000; Ginter et al. 2002; Duncan 2003; Ginter and Turner 2010) has suggested the presence of two tooth types; the first with a flattened base positioned asymmetrically in relation to the crown and a smaller almost completely symmetrical form (Duncan 2003). The first morphotype (Text-fig 4A) is the most common tooth form attributed to *Th. tranquillus* and has been reported from various locations across northern Gondwana (Ginter and Turner 2010). The second, smaller morphotype (Text-fig. 4B), which has had its assignment to *Th. tranquillus* recently questioned (Ginter and Turner 2010), is less common and with few examples recorded from the Montagne Noire (Ginter 2000). Based on reconstructions of tooth placement within thrinacodont jaws (Turner 1982; Ginter et al. 2002; Duncan 2003), the supposed symmetrical symphyseal teeth only comprise a small proportion of teeth thus likely lower yields of these teeth are to be expected.

A third morphotype (Text-fig. 4C) has been attributed to *Th. tranquillus* with similar teeth previously recorded from Montagne Noire and Oum El Jerane and Tizi Nersas, Morocco, conversely these specimens were designated *Thrinacodus cf. ferox* (Ginter 2000, fig. 2; Ginter et al. 2002, fig. 9A–E). It was suggested by Ginter et al. (2002) that the teeth of *Th. cf. ferox* may represent the lateral most teeth of *Th. tranquillus*, which was supported by the presence of intermediate forms between the typical teeth of *Th. tranquillus* and *Th. cf. ferox*. A similar variety of tooth forms is present in the Canning Basin, however the teeth are recovered from Late *marginifera* CZ, dated deposits making their attribution to the *Th. ferox* unlikely as it is regarded as a late Famennian to late Tournaisian species (Ginter and Turner 2010).

DISTRIBUTION AND STRATIGRAPHIC RANGE: In the Canning Basin this species is recorded from the Late *marginifera* CZ. A similar age *rhomboidea* or *Late marginifera* CZ is recorded for Chahriseh, Iran (Long and Hairapetian 2000; Ginter et al. 2002; Hairapetian and Ginter 2009) and also a *marginifera* CZ for Hunan, China (Lelièvre and Derycke 1998). In Oum El Jerane and Tizi Nersas, Morocco this species ranges from *rhomboidea* or *Late marginifera* CZ to Middle to Late *expansa* (Ginter et al. 2002; Derycke et al. 2008). Younger known occurrences have been recorded from the Late *trachytera* CZ, Holy Cross Mountains, Poland, Lower to upper *expansa* CZ, Thuringia, Germany and *expansa* CZ in Montagne Noire. Within the South Urals the range of this species is recorded from the Late *expansa* to Early *praesusulcata* CZ (Ivanov 1996).

**TYPE SPECIES:** *Diademodus hydei* Harris, 1951
Diademodus dominicus sp. nov.
(Text-fig. 4D–E)

ETYMOLOGY: In honour of Mr Dominicus ‘Tim’ Mueller, M.Sc., who guided the careers of many aspiring geoscientists.

HOLOTYPE: Specimen WAM 14.7.1 from the Virgin Hills Formation, Horse Spring, Canning Basin, Western Australia; sample VHS-310; Frasnian, MN 11 CZ; Text-fig. 4D.

MATERIAL: Two teeth from the Virgin Hills Formation, Horse Spring, sample VHS-310.

DIAGNOSIS: Teeth with a multicuspid phoebodont type crown comprising three cusps and three to four pairs of intermediate cusplets (Text-fig. 4D). The cusps are thin and ovoid in cross section, becoming circular towards the cusp apex and bearing prominent lateral carinae. A slight lingual inclination is present on all cusps with the larger lateral cusps recurving occlusally. The base is highly arched, trapezoid in outline and becoming narrower labially (Text-fig. 4E). The labial margin is rounded and perforated by a row of foramina. A mesiodistally elongate hump occupies the majority of the basolabial shelf with the labial margin abutting the base of the cusps. There is a slight decrease in height of the shelf towards the lingual margin.

DESCRIPTION: Distinctive teeth with symmetrical crowns that fan out mesiodistally and comprise three main cusps and three to four pairs of lateral cusps (Text-fig. 4D–E). The medial cusps are broken in both specimens, however the bases are approximately half the width of the lateral cusps. A pair of small, first lateral cusplets are independent of the medial cusp but are fused at the base with the second lateral cusplets which are slightly smaller in size than the lateral cusps. In one holotype (WAM 14.7.1, Text-fig. 4E) a small fourth cusplet is present and erupts from the base of the lateral cusp. The preserved cusps appear ovoid in cross section (elongate labio-lingually), at the base and exhibit smooth lingual and labial faces with well-developed lateral carinae. No discernible crown-base interface is present on the labial faces of the teeth. The base is highly arched (Text-fig. 4E), approximately twice as long mesiodistally as it is labioliungually (Text-fig. 4E). A roughly trapezoid to almost rectangular shape of the base can be determined, with rounded lingual corners. A row of foramina is present running along the lingual rim. An ovoid button is positioned centrally on the base extending between the third lateral cusplets and gradually dissipates towards the lingual margin of the base. The labial face of the button preserves the openings to a row of canals and terminates abruptly, close to the base of the crown.

REMARKS: The teeth of the type specimen Diademodus hydei Harris, 1951 are partially obscured in matrix thus similarities to the genus are based on the visible labial tooth face of Diademodus hydei as well as comparison to the other member of this genus, Diademodus utahensis Ginter, 2008. The teeth of Diademodus all share the typical phoebodont tooth crown but can be differentiated on the number of intermediate cusplets (Ginter 2008). The cusp arrangement and variation in cusplet size present in D. dominicus sp. nov. (Text-fig. 4E) is similar to D. utahensis (Ginter 2008, fig. 1B) and differs from the original description of D. hydei which was figured having two prominent central cusps and intermediate cusplets of almost equal size (Harris 1951, fig. 3b), however, reinterpreted by Ginter (2008) as possessing a single prominent central cusp. The teeth of D. dominicus sp. nov. are further distinguished as some cusplets are fused at the base with the adjacent cusp or cusplets, a feature not apparent in D. hydei or D. utahensis. The lingual face of teeth from D. dominicus sp. nov. most closely resemble those of D. hydei as these teeth lack both the mesiodistal extension of the base beyond the crown as well as the presence of a basolabial shelf, features both present in D. utahensis. The tooth is distinguished from both D. hydei and D. utahensis by a shorter mesiodistal length of the base which is also significantly arched. The basal outline of D. dominicus sp. nov. narrows lingually and is substantially different from that of D. utahensis which forms a trapezoid shape becoming wider lingually and slightly compressed centrally on the lingual margin. The orolingual hump of D. dominicus sp. nov. is also far more prominent than that of D. utahensis. We believe these differences are significant enough for the establishment of the species Diademodus dominicus sp. nov.

Despite the low number of teeth recovered from D. utahensis and the obscured nature of teeth from D. hydei, heterodony within each species is so far limited to variation in the number of intermediate cusplets, reported in D. utahensis (Ginter 2008) and D. dominicus sp. nov. Ginter (2008) suggests that anterior teeth of D. utahensis may be narrower and speculates they would be similar in form to teeth attributed to Ph. fastigatus (in Ginter 2008, fig. 3B). This would appear more likely than the teeth of D. dominicus sp. nov. representing anterior D. utahensis teeth as the base and cusp morphology is far more similar between D. utahensis and the Ph. fastigatus tooth examples.
Order Symmoriida Zangerl, 1981
Family Stethacanthidae Lund, 1974
Genus Stethacanthus Newberry, 1889

TYPE SPECIES: Physonemus altonensis St. John and Worthen, 1875.

Stethacanthus sp. 1
(Text-fig. 4F–G)

MATERIAL: One tooth from the Virgin Hills Formation, Casey Falls, sample CL-471; one tooth from the Bugle Gap Limestone, Casey Falls, sample 1984-97.

DESCRIPTION: Small teeth, less than 1 mm mesiodistally and consisting of a slender prominent central cusp and two smaller, slightly diverging lateral cusps approximately one third of the size of the medial cusp (Text-fig. F–G1). A pair of small intermediate cusplets is present with one specimen possessing a second pair of cusplets on the distal side of the lateral cusps (Text-fig. 4F1–2). Thin vertical striations extend from the base to tip on the lateral and intermediate cusps on the well-preserved cusps. The base forms a distinct triangular outline (Text-fig. G2), extending beyond the crown mesio-distally with rounded edges and a low profile. The apex of the lingual projection on the base is slightly thickened and possesses multiple foramina along the lingual rim.

REMARKS: The teeth described here possess characters found in other stethacanthid tooth types including upright, unconnected cusps ornamented in sub-parallel cristae and a base lacking a labial depression. The triangular shaped lingual extension of the base (Text-fig. 4G2) is a common feature of Moroccan stethacanthid teeth (Ginter et al. 2002, pl. 10C) and also seen in other Famennian teeth from the Montagne Noire (Ginter 2000, fig 7C) and Dalme, Iran (Long and Hairapetian 2000, fig 4i). The teeth from the Canning Basin also appear to exhibit a thickening around the lingual apex of the base, which is similar to a tooth from Dalme, Iran (Long and Hairapetian 2000, fig. 4e) but lack a well-formed labial boss. One tooth variant (WAM 14.7.26, Text-fig. 4F) exhibits a sixth cusp, a feature not usually found in Famennian stethacanthid teeth from Morocco and Iran (Ginter et al. 2002). Further specimens are required to determine if this tooth represents a non-typical variant of Stethacanthus sp. 1 or if there is a degree of heterodonty within this species.

Order Ctenacanthiformes Glikman, 1964
Family Ctenacanthidae Dean, 1909
Genus Cladoides Maisey, 2001

**Cladodoides wildungensis** Jaekel, 1921  
(Text-fig. 5A)

1921. *Cladodoides wildungensis* sp. nov.: Jaekel, 1921.  
1991. *Stethacanthus* sp.; Ginter, p. 75, pl. 8: 4, pl. 9: 2, 3.  
1995b, *Stethacanthus thomasi* Turner; Ginter, fig. 2A.  
1996. “*symmoriid with button partially divided*”; Ginter and Ivanov, fig. 4C.  
1996. “*stethacanthid?*”; Ginter and Ivanov, fig. 5C, D.  
2000. *Stethacanthus* cf. *thomasi* (Turner); Ginter and Ivanov, pl. 1: J.  
2002. *Stethacanthus resistens* sp. nov.; Ginter et al., figs 2, 3, 4C–I, 5C.  
2010. *Cladodoides wildungensis* Jaekel; Ginter et al., fig. 66A–J.

**MATERIAL:** Two teeth from the Virgin Hills Formation, Horse Spring, sample VHS-310; Casey Falls, sample Cl-9.

**DESCRIPTION:** Symmetrical teeth with five lingually inclined and strongly diverging cusps. The central cusp is large, almost twice the width at the base of the lateral cusps (Text-fig. 5A1). The intermediate cusps are significantly smaller, approximately half the height of the lateral cusps. Both the lingual and labial faces of the cusps are ornamented with sub parallel ridges. The base is ovoid in outline, slightly arched and elongated mesiodistally, extending beyond the crown. A small basirolabial projection is preserved on one tooth (Text-fig. 5A1), between the distal edges of the medial cusp. Due to abrasion, only the remnants of a mesiodistal elongate button are preserved on both teeth. One tooth (Text-fig. 5A1) preserves an outline of a button, extending the distance between the two intermediate cusps, as well as a series of grooves, presumably once pore canals running labiolingually.

**REMARKS:** Ginter et al. (2010) recognises two morphotypes of *Cladodoides wildungensis* Jaekel (1921). The first morphotype, characterised by its larger size and long, slender medial cusp; and the second smaller morphotype, to which the Canning Basin teeth are here assigned, that are less than two mm across the length of the base with diverging cusps ornamented in coarse sub-parallel cristae. The teeth described here are markedly similar to the holotype of *C. wildungensis* (figured in Ginter et al. 2010, fig. 66l) as well as resembling teeth from Poland (Ginter et al. 2002, fig. 2C–F) designated *Stethacanthus resistens*. Ginter et al. (2010) suggests that it is possible that some, if not all teeth previously attributed to *Stethacanthus resistens* belong to the *C. wildungensis*. If so, it appears *C. wildungensis* is considerably cosmopolitan in nature.

**DISTRIBUTION AND STRATIGRAPHIC RANGE:** This species ranges from MN 11 CZ– lower crepida in the Canning Basin, Western Australia. Due to a degree of synonymy with *S. cf. thomasi* a definitive range is difficult to ascertain, however a range from the Upper Frasnian – middle Famennian is likely in Poland, Moravia, Germany, Morocco and Russia.

Ctenacanthiforme gen. et sp. indet 1  
(Text-fig. 5B)

**MATERIAL:** One tooth from the Napier Formation, South Oscar Range, sample SO-200.

**DESCRIPTION:** This highly symmetrical tooth is well preserved, by comparison to other teeth in the sample, and possesses five slender, lingually inclined cusps that fan out mesiodistally (Text-fig. 5B1 –2). A large prominent central cusp is flanked by two smaller intermediate cusps. The lateral cusps are approximately two thirds the length of the medial cusp and diverge at approximately 45 degrees to the medial cusp. In labial view, the cusps appear more triangular and connected at the bases (Text-fig. 5B3). All cusps are ovoid in cross section and ornamented on both the lingual and labial faces with strong vertical striations. The base is trapezoid in shape extending further linguallly than labially and broadest along the lingual margin (Text-fig. 5B3). The base is thickest in a trapezoid area between the lingual edge of the lingual button and lateral cusps. A thin baso-labial projection is present with a row of foramina. A well developed, ovoid shaped button is positioned close to the lingual border and perforated by a row of four large foramina (Text-fig. 5B1).

**REMARKS:** The tooth shares commonalities between many cladodont-like teeth outlined by Duffin and Ginter (2006). The connection of the tooth cusps by an enameloid or orthodentine layer (Text-fig. 5B3) excludes this tooth from belonging to the symmoriids and stethcanthids. The tooth lacks a basirolabial depression as seen in *Cladodus* and *Ctenacanthus* and is thereby more similar to *Cladoides* which may also lack this feature. The baso-labial shelf is straight which is characteristic of *Cladoides* but is far more developed in Ctenacanthiforme gen. et sp. indet 1 and this pro-
jection may obscure the presence of any basolabial depression. The tooth is differentiated from other Ctenacanthiformes based on the strongly rounded or bi-convex central cusp and lateral cusps, which are almost two thirds the size of the central cusps. The tooth is referred to the Ctenacanthiformes, perhaps temporally, based on the orthodentine or enameloid connective tissue present between the cusps and similarities with the genus *Cladoides*.

Ctenacanthiform gen. et sp. indet 2  
(Text-fig. 5C)

MATERIAL: One tooth from the Virgin Hills Formation, Casey Falls, sample 1984-97; one tooth from the Napier Formation, South Oscar Range, sample SO-200.

DESCRIPTION: Heavily abraded teeth with a large medial cusp flanked by four smaller triangular shaped cusps consisting of a pair of small lateral cusps and a pair of intermediate cusplets approximately half the size of the lateral cusps (Text-fig. 5C1-2). The medial cusp is inclined slightly and convex along the lingual edge in cross section. Where the outer enameloid is preserved on the central cusp, faint vertical striations can be observed. The intermediate and lateral cusps do not show any ornamentation with the lateral cusps diverging at a 45 degree angle. A distinct rim is preserved on the labial face of the teeth and marks the crown base interface (Text-fig. 5C1-2). The base extends lingually almost three times the width of the crown and forms a large dome shape that extends beyond the crown mesiodistally (Text-fig. 5A1). A very faint ovoid button, approximately twice as long as it is wide, is preserved on one specimen and positioned closer to the lingual margin than the crown (Text-fig. 5A2). A series of foramina perforate the lingual face of the base with a single large nutritive canal positioned centrally along the lingual margin. The labial face of the base in both specimens is highly abraded (Text-fig. 5A3), however remnants of a thickened baso-labial shelf and slight depression are present.
REMARKS: These teeth have been attributed to the Ctenacanthiformes based on features shared with other members of this group. The prominent central cusp and diminished lateral and intermediate cusps are similar to *Cladodus conicus* Agassiz, 1843 although the base on the tooth from the Canning Basin specimens is much thicker and the striations preserved on the cusps are coarser. The central cusp is flattened along the labial face and exhibits a convex lingual face, and in this respect closely resembles *Cladodus*. The significant abrasion to the labial faces of the teeth does not allow further taxonomic refinement. The teeth do feature a unique rim running across the crown base interface that is not typically seen in other cladodont type teeth.

DISTRIBUTION AND STRATIGRAPHIC RANGE: Frasnian conodont Zone 13b to Late *marginifera* conodont Zone in the Canning Basin, Western Australia.

Cohort Euselachii Hay, 1902
Order indet.
Family Protacrontidae Zangerl, 1981
Genus Protacrodus Jaekel, 1925

**TYPE SPECIES:** *Protacrodus vetustus* Jaekel, 1925

*Protacrodus serra* Ginter, Hairapetian and Klug, 2002

(Text-fig. 6A–B)

1990. “*Cladodus*” sp.; Ginter, p. 77, pl. 4, fig. 9.
2002. *Protacrodus serra*, sp. nov.; Ginter, Hairapetian and Klug, p. 195, text-fig. 11; pl. 2, figs L–N; pl. 11, figs A–C.
2010. *Protacrodus serra* Ginter, Hairapetian and Klug, 2002; Ginter et al., p. 87, fig. 80A–C.
2011. *Protacrodus serra* Ginter, Hairapetian and Klug, 2002; Ginter et al., p. 168, fig. 10J–J.


DESCRIPTION: The teeth described here can be assigned to morphotype one of *Protacrodus serra* (Ginter et al. 2002, text-fig. 11 C–E) and comprise asymmetrical crowns with three cusps directed distally to one side (Text-fig. 6A). The medial cusp is large, over twice the size of the lateral cusps with a narrow labiolingual profile. Both the lingual and labial faces of the cusps are ornamented in strong cristae. In one specimen (WAM 14.7.23, Text-fig. 6B), a row of small pointed cusplets is present along the crown base interface on the labial face. The base is narrow lingually, with a slightly extended and straight lingual rim (Text-fig. 6B) that is perforated by a large canal located centrally. The labial face of the base does not protrude and is perforated by a row of small foramina (Text-fig. 6A2).

REMARKS: Evaluating variation present in the teeth of different shark taxa is difficult as odontological studies remain incomplete (Straube et al. 2008). The presence of labial cusplets (Text-fig. 6B) on *P. serra* may be evidence of intraspecific variation or the presence of sexually based heterodonty which is seen in the dentition of some fossil (Parmley and Cicimurri 2003) and extant shark species (Raschi et al. 1982). The same labial cusplets are apparent in other teeth attributed to *P. serra* from Iran (Ginter et al. 2002, pl. 2L–M; Ginter et al. 2011, text-fig 10J–J), which rules out regional influences for variation. Cusplets also appear on teeth independent of size (Text-fig. 6A–B) and morphotype (Ginter et al. 2002, pl. 2L–M) which makes ontogenetic variation unlikely.

DISTRIBUTION AND STRATIGRAPHIC RANGE: This species has an older occurrence in the Canning Basin, Western Australia (Lower *crepida* to Upper *marginifera* Zone) as compared to other areas of the world where it occurs in the Early *expansa* Zone in Dalmeh, Iran, Late *expansa* Zone, Tizi and Oum El Jerane, Morocco and probable *expansa* Zone Khor Virap, Armenia. In China this species has been recorded from the middle Tournaisian. Another protacodont species, *D. masureae* (Burrow et al. 2010) is known to range from the Late Devonian into the Carboniferous.

Genus Deihim Ginter, Hairapetian and Klug, 2002

**TYPE SPECIES:** *Deihim mansureae* Ginter, Hairapetian and Klug, 2002

*Deihim mansureae* Ginter, Hairapetian and Klug, 2002 (Text-fig. 6C–D)

2000. *?Protacrodus* sp. Long and Hairapetian, p. 217–218, fig. 4O.

2002. Deihim mansureae gen. et sp. nov.; Ginter, Hairapetian and Klug, p. 191–193, text-fig. 10; pl. 1, fig. r; pl. 2, fig. k; pl. 4, figs F–G, J–M; pl. 5, figs A–M.


2009. Deihim mansureae Ginter, Hairapetian and Klug, 2002; Hairapetian and Ginter, p. 176, 179, fig. 2D, 4H.

2010. Deihim mansureae Ginter, Hairapetian and Klug, 2002; Hairapetian and Ginter, p. 362, fig. 3A.

2010. Deihim mansureae Ginter, Hairapetian and Klug, 2002; Ginter et al., p. 88, fig. 81A–J.

2011. Deihim mansureae Ginter, Hairapetian and Klug, 2002; Ginter et al., p. 166, 169, 8A–E, 11C.


2013. Deihim mansureae Ginter, Hairapetian and Klug; Habibi et al., p. 30, fig. 4.


DESCRIPTION: Four morphotypes were originally assigned to the species Deihim mansureae Ginter et al. 2002, of which, two are represented here (Text-fig. C–D). A single tooth attributed to morphotype one (Text-fig. D1–2) appears to possess a significantly worn central and lateral cusps with only the base of the crown preserved.

The base of a large main central cusp can be distinguished, flanked by two smaller highly fused lateral cusps. Cristae is present on the base of the lateral cusps on the lingual face. The labial face of the crown possesses three large, robust cusplets (Text-fig. 6D3), immediately above the base. The crown-base interface on the lingual face is highly arched and defined by a shallow groove. The base is narrower, mesiodistally, than the crown and extends lingually with a series of deep canals.

Two teeth (Text-fig. 6C1–2), assigned to morphotype four were recovered and are characterised by an elongate central cusp and two well-spaced diverging lateral cusps approximately one third of the size of the medial cusp. Sub-parallel cristae are present on both the labial and lingual faces of each cusp. Four cusplets are present above the crown-base interface on the labial face. The base is roughly ovoid in outline (Text-fig. 6C2) and does not extend beyond the crown mesiodistally. The lingual face of the base forms an arch, protruding lingually and is perforated by a few foramina.

REMARKS: The teeth described here conform to the diagnosis of Deihim mansureae with the presence of a prominent central cusp and lateral cusps that bear strong cristae, on a convex crown-base interface. A row of labial cusplets present above the crown base interface is also characteristic of this species.

The teeth assigned to morphotype four resemble those of Ginter et al. (2010, pl. 5 D–F) with the presence of a slender central cusp with two lateral cusps, approximately half the size of the central cusp and diverging at 45 degree angles. These teeth also appear less perforated foramina on the lingual face of the base and lack the fusing of the central and lateral cusps that is seen in the other morphotypes of this species.

DISTRIBUTION AND STRATIGRAPHIC RANGE: Deihim mansureae ranges from the Late marginifera CZ to possibly the Tournaisian in the Canning Basin. An older occurrence of this species is known from the rhenana (MN 11 CZ) in Kale Sardar area of Iran (Hairapetian and Ginter 2010). In the Famennian of Iran and Armenia Deihim mansureae ranges into the Upper crepida conodont Zone. Teeth attributed to this species have also been found in the Sly Gap Formation, New Mexico, USA (Ivanov and Lucas 2011), however no concise age range beyond the Late Devonian and Early Carboniferous can be ascertained.

Deihim cf. mansureae Ginter, Hairapetian and Klug, 2002 (Text-fig. 6E)

MATERIAL: Two teeth from the Bugle Gap Limestone, Casey Falls, sample 1984-98.

DESCRIPTION: Teeth with prominent central cusps and two smaller lateral cusps approximately one third the size of the medial cusp, diverging at approximate 45 degree angles (Text-fig. 6E1). Each cusp is rounded with sub parallel striations on both the labial and lingual faces. A row of seven rounded cusplets, ornamented in strong cristae are present on the base of the crown along the labial face (Text-fig. 6E2). The base is ovoid in outline, extending lingually and possessing a single small pore located centrally on the lingual rim. On the labial face the base is thin and forms an undulated contact with the crown.

REMARKS: These teeth are similar to the morphotype 1 of Deihim mansureae as they comprise a crown with a large central cusp with two strongly diverging lateral cusps, all ornamented in a strong cristae as well.
as a row of lingual cusplets along the labial face of the tooth. The cusplets present on the teeth described here are unusual but bear a strong resemblance to the cusplets on a tooth from *D. mansureae* in central Iran (Hairapetian and Ginter 2009, figs 5H2–3). Despite the similarities with *D. mansureae*, the tooth lacks both the arched lingual groove associated with the crown base interface as well as the series of deep pore canals on the lingual face diagnostic of this species (Ginter et al. 2002). These features are also lacking on a tooth attributed to *Protacodus cf. vetustus* from the Montagne Noire (Ginter 2000, fig. 8C). The Montagne Noire tooth shares the similar sub-parallel cristae to the teeth from the Canning Basin, however it lacks labial cusplets.

*Deihim* sp. 1
(Text-fig. 6F)
DESCRIPTION: A well preserved, almost symmetrical tooth with a prominent central cusp and two pairs of highly fused lateral cusplets (Text-fig. 6F1). Cusps are ornamented on the lingual face in strong cristae along the base of the cusps with a single ridge running up the centre of the central cusp. A crown-base interface is prominent, marked by a groove on both the lingual and labial faces of the tooth (Text-fig. 6F1). A triangular labial peg is positioned at the base of the central cusp with accompanying smaller projections on either side (Text-fig. 6F2–3). The base is approximately the same height as the central cusp with a straight lingual edge and rounded distal margins. A series of foramina are present on the lingual face of the base in almost two rows, one slightly below the crown base interface and another close to the lingual rim (Text-fig. 6F1). A large articulation socket is present on the basal-labial face of the base (Text-fig. 6F2).

REMARKS: The tooth described here comprises a large central cusp with diverging cusps ornamented in strong cristae which are diagnostic of the genus Deihim (Ginter et al. 2002). Further similarities include the presence of a crown base interface separating a lingually extended base perforated by a series of canals. Despite the similarities, Deihim sp. 1 comprises a central cusp which is more prominent than that found on morphotype 1 of D. mansureae and it also lacks the characteristic labial cusplets. Instead the tooth from the Canning Basin bears small tubercle-like projections above the crown base interface on the extended labial face of the crown. The general morphology of Deihim sp. 1 also bears a slight resemblance to Protacrodus orientalis Li, 1988 with the high central cusp, fused laterals and a labial projection of the central cusp (Li 1988, pl. 1, fig. 2), however the size difference in the cusps excludes this tooth from P. orientalis (Ginter et al. 2002). There is also a very strong resemblance to another tooth attributed to P. orientalis from the Menggongao Formation, South China (Lelièvre and Derycke 1998, fig. 7B) which is narrower mesiodistally with the same cusp morphology and deep base. Recently, the inclusion of P. orientalis as a species of Protacrodus has been questioned (Ginter et al. 2010). It is possible that P. orientalis belongs to the genus Deihim and the tooth from the Canning Basin represents a less abraded form of this species.

Protacrodonididae gen. sp. indet. 1
(Text-fig. 7A)

MATERIAL: One tooth from Bugle Gap Limestone, Casey Falls, sample 1984–94.

DESCRIPTION: Highly elongate and slightly arched tooth comprising a prominent pyramidal central cusp accompanied by four to five fused lateral cusps approximately a quarter of the size of the central cusp (Text-fig. 7A1). Both the labial and lingual faces of the cusps are ornamented in coarse wavy cristae directed towards the cusp apices. Approximately 80% of the lingual face of the base is broken, however the preserved section is the same height as the lateral cusps and extends both lingually and further mesiodistally than the crown on one side (Text-fig. 7A2). A small number of foramina are present on the lingual face of the base. The labial margin of the base is thin and perforated by small foramina and lacks a distinct crown base interface (Text-fig. 7A2).

REMARKS: The tooth is placed under Protacrodonididae based on the coarse cristae and lack of articulation devices also found in other Devonian Protacrodonts as well as its occurrence within the middle Devonian. The tooth is also similar to a partial tooth of a ?Protacrodon (Ginter and Sun 2007, fig. 4E) from Muhua, southern China in that it shares the highly elongate shape, with extensively fused lateral cusps flanking a prominent central cusp and a well-developed cristae. However, the tooth described here lacks the festoon-like sculpture on the lower aspect of the crown present on the Chinese tooth (Ginter and Sun 2007).

?Protacrodonididae fam. gen. sp. indet.
(Text-fig. 7B)

MATERIAL: One tooth from the Bugle Gap Limestone, Casey Falls, sample CL-471.

DESCRIPTION: A mesio-distally elongate tooth with a crown consisting of highly fused cusps that are almost indistinguishable. Strong cristae are present on the labial and lingual faces as well as the occlusal surface (Text-fig. 7B₁₋₃). The labial face of the crown is higher and slopes lingually. The distal margins of the crown taper inward to the crown-base interface, whereas the labial and lingual faces are almost straight. The base is roughly rectangular in outline with a slightly thickened, protruding lingual margin. A row of pore canals are present along the lingual face of the base (Text-fig. 7B₂), where as possible remnants of canals are present on the labial face (Text-fig. 7B₃).
REMARKS: Teeth with coarse cristae, a reduced base lacking both a lingual extension and articulation devices are common among Devonian protacrodonts. The crown of the tooth described here, retains more highly fused cusps than other protacrodont teeth and in this respect it more closely resembles an orodont. The reduced base and lack of distinct margin between the crown and base along the labial margin appear similar to the lingual face of the teeth attributed to *P. serra* and *Protacrodus aequalis* Ivanov 1996. The samples taken above the section measured at Casey Falls are dominated by protacrodontid-type teeth (*Protacrodus serra*, *Deihim mansureae*, *D. cf. mansureae*, *Deihim* sp. 1, *Protacrodontidae* gen. et sp. indet.; Table 1), which could suggest it belongs to the family Protacrodontidae and is highly modified, possibly representing a lateral crushing tooth.

**Order Hybodontiformes** Cappetta, Duffin and Zidek, 1993

**Superfamily Hybodontoidea** Owen, 1846

**Family Lonchidiidae** Herman, 1977

**Genus Lissodus** Brough, 1935

**TYPE SPECIES:** *Hybodus africanus* Bloom, 1909.

*Lissodus lusavorichi* Ginter, Hairapetian and Grigoryan, 2011

(Text-fig. 7C)

2011. *Lissodus lusavorichi* sp. nov. Ginter, Hairapetian and Grigoryan, p. 160, fig. 10E–F.

**DESCRIPTION:** The tooth crown is broadly triangular with a well-developed occlusal crest (Text-fig. 7C1). The crown extends laterally beyond the base terminating in triangular shoulders. The crown base interface is slightly arched on the lingual face and prominent on the labial face (Text-fig. 7C2). A well-developed labial peg is present and accompanied by small cusplets or tubercle like projections. The base is typically euselachian, projecting lingually with a rounded margin. A series of deep, elongated canals perforate the lingual face of the base (Text-fig. 7C).

**REMARKS:** The teeth from Casey Falls are attributed to *Lissodus lusavorichi* Ginter *et al.* 2011, as they possess the same smooth broad triangular crown, well-developed occlusal crest and horizontal longitudinal crown shoulders as well as the tubercles associated with the labial peg. The presence of this species represents the oldest record of the family Lonchididae in the Canning Basin.

**DISTRIBUTION AND STRATIGRAPHIC RANGE:**

The species has been previously described from the *expansa* zone in Dalmeh, Iran and probable *expansa* zone, Khor Virap Armenia. Its presence in the Upper *marginifera* conodont Zone in the Canning Basin, refines the lower age range of this species.
DISCUSSION

Biostratigraphy

Thelodonts have been successfully used in biostratigraphy since the 1960’s (Gross 1967) with current global zonation schemes (in Gradstein et al. 2012) developed for the Silurian (Turner 1973; Bassett et al. 1982; Märs et al. 1995; Blom 1999; Taliamaa 2000) and Devonian, (Turner 1995a, 1999; Blom and Goujet 2002; Hairapetian et al., in press). Subsequent to the Givetian–Frasnian boundary, thelodonts are restricted to the margins of northern Gondwana (Turner 1997; Trinajstic 2001; Märs et al. 2007) and there is a marked reduction in diversity with the Turinidae Family represented by the species Australolepis seddoni (Seddon 1969; Turner and Dring 1981; Trinajstic and George 2009), Turinia hutkensis (Bliick and Goujet 1978; Hairapetian et al. 2006) and a new genus and species of turinid (Hairapetian et al., in press). Of these, A. seddoni has been particularly useful in defining shallow water Frasnian strata in northern Gondwana (Turner et al. 2002). The species is known from the falsiovalis to rhenana conodont Zones (MN 1–10 CZ) in Iran and provided a lower Frasnian age constraint for the Charishe section (Turner et al. 2002). Within Australia this species is now recognised as ranging from MN 6–11 CZ and has been previously used to constrain a Frasnian age for the base of the Gneudna Formation type section in the Carnarvon Basin (Turner and Dring 1981; Trinajstic and George 2009). Within the Canning Basin A. seddoni has been used to confirm a Frasnian age for the Hull platform (Chow et al. 2013). Although the increased range of A. seddoni, decreases the ability to obtain more finely resolved ages for strata, its confinement to the Frasnian still makes this species an important zone fossil.

The phoebodont taxa recovered in this study conform to previously established age ranges obtained for these species in the Canning Basin (Trinajstic and George 2009) with the exception of Phoebodus latus where the range has been extended from MN 10 CZ to MN 11 CZ. This newly extended range of Ph. latus in the Canning Basin brings it closer in line with the known upper age range for this species (MN 13 CZ) in the Southern Urals and South Timan in Russia as well as the Holy Cross Mountains in Poland (Ginter and Ivanov 1995b; Ivanov 1999; Ivanov and Lucas 2011). The extension of the known age ranges for A. seddoni and Ph. latus in the Canning Basin indicates further refinement of the age ranges for other phoebodont taxa such as Ph. bifurcatus and Ph. fastigatus is likely with further collecting. Work so far undertaken describing the phoebodont fauna from the Canning Basin (Trinajstic and George 2009), has proven their utility in biostratigraphy as well as increasing the geographic distribution in which these species can be used.

A disappearance in phoebodont fauna is seen in the section at Horse Spring from MN 11 CZ and occurs too early to be attributed to the global Kellwasser Events. Despite the distinct anoxic black limestones and shales associated with the Kellwasser horizons not being recognised in the Canning Basin (Becker et al. 1993), perturbations are recorded in the faunas. Reductions in the species diversity and abundances are recorded within conodonts (Klapper 2007), ammonoids (Becker et al. 1993) and trilobites (Feist and McNamara 2013) in the Canning Basin which are also seen in the rest of the world. The absence of microvertebrates at Horse Spring above MN CZ 11 was previously noted by Trinajstic and George (2009) who suggested post depositional dolomitisation as the main factor contributing to the lack of microvertebrates. However, this does not seem to be the cause as microvertebrates are often recovered from dolomitised beds. In addition, conodonts were found in the dolomitised beds from 13–19 m (Text-fig. 2) in the section at Horse Spring where microvertebrate fossils were lacking. A global non-occurrence of the Phoebodus bifurcatus group from the linguiformis Zone (MN 13b CZ) in the Frasnian to the Famennian late crepida conodont Zone has been identified in the Northern Hemisphere (Ginter and Ivanov 1995b; Ginter et al. 2002). Ginter et al. (2002) suggests the disappearance mainly resulted from global cooling (Copper 1998) rather than fluctuations in sea level or anoxia. Given the much earlier disappearance of phoebodonts in the section at Horse Spring, and no evidence of major temperature excursions in sea surface temperatures during the MN 11 CZ (Joachimski et al. 2004), it is not likely temperature fluctuations are the reason for the disappearance of a phoebodont non-occurrence at Horse Spring.

Another notable global change recorded towards the end of the Frasnian is the semichatovae Transgression (Sandberg et al. 1997). This event occurs at the base of the MN 11 Zone and coincides with the absence of phoebodont taxa in the Canning Basin. Despite reported localised tectonic overprinting (Southgate et al. 1993; George and Chow 2002; Chow et al. 2013), evidence of a longer-term 2nd order transgressive event is observable in these sections. The effect of these localised and global events may be highly influential in the faunal changes and extinctions seen in the basin (George and Chow 2002). The phoebodont group has been previously noted as niche sensitive (Ginter...
Biogeography and chondrichthyan assemblages

Chondrichthyan biofacies present in the depauperate Famennian intervals of the Virgin Hills and Napier formations are often difficult to determine. These intervals often comprise largely shallow water and slope derived material. This results in the microvertebrate remains often not representing an original environment of habitation. In contrast to the described sections, samples taken from the shallow fore-reef deposits of the Bugle Gap Limestone (Guppy et al. 1958; Playford and Lowry 1966; Playford et al. 2009) preserve a more original chondrichthyan biofacies. A diverse shark genera including Thrinacodus, Stehacanthus, Protacrodus, Deihim and Lissodus were present in these shallow water facies. Despite the low numbers of teeth recovered from the Bugle Gap Limestone, the proportionally higher number of protacrodont and hybodont type sharks as well as the absence of Phoebodus and Jaloconodontus, is indicative of a shallow water Protacrodus biofacies (Ginter 2000, 2001). Similar assemblages have also been recorded in Famennian strata from both Dalmeh and Chahriseh, Iran (Ginter et al. 2002; Hairapetian and Ginter 2010).

An increase in cosmopolitanism is seen in many Late Devonian marine groups, including brachiopods (Copper 1998), ammonoids (House 1973), trilobites (McNamara and Feist 2006), crustaceans (Rode and Lieberman 2005), conodonts (Girard et al. 2010) and fish (Ginter and Turner 1999; Burrow et al. 2010). This faunal exchange between areas was likely driven by the closure of the Rheic Ocean (Nance et al. 2012; Domeier and Torsvik 2014) during the Late Devonian, culminating in an increased proximity of Laurussia and Gondwana (Young 2003; Burrow et al. 2010; Young et al. 2010). This period also saw the rifting of the Palaeotethys resulting in a separation of the north
UPPER DEVONIAN MICROVERTEBRATES FROM WESTERN AUSTRALIA

China, Indochina, Tarim and South China blocks from Gondwana (Metcalfe 1984, 2011). Connections between South China and north-west Australia were maintained in the Late Devonian by the occurrence of an extensive, shallow water shelf (Golonka 2007; Metcalfe 2011; Golonka and Gaweda 2012). Previous work on Late Devonian conodonts (Burrett et al. 1990; Girard et al. 2010) and fish (Long 1993) support a close relationship between the two areas. In addition, Trinajstic and George (2009) had previously shown close affinities, in regards to the temporal ranges, of the Frasnian phoebodonts *Ph. fastigatus* and *Ph. bifurcatus* in the Canning Basin and South China. The presence of a likely shallow water taxon, *Deihim* sp.1 (Text-fig. 6G1–3) in both the Canning Basin and Menggongo Formation of South China (Lelièvre and Derycke 1998, fig. 7B), would support the presence of a shallow water platform between the two areas into the middle part of the Famennian. Determining the extent of faunal exchange further into the Famennian or even Carboniferous is difficult, as studies into shark faunas of these ages are limited for the north west of Australia. However *Th. ferax* has been reported in the late Famennian in both South China and Australia (Long and Burrett 1989). The similarities between the tooth identified here as Protacrodonidae sp. 1 (Text-fig. 7A1–2) found in a middle Famennian of the Virgin Hills Formation and a tooth from the middle Tournaisian in Muhua, South China (Ginter and Sun 2007, fig. E1–2) provides further support for a close connection between China and Australia (Metcalfe 2001) however, further data from microvertebrates are required.

Clear links between the shark taxa in the north-west of Australia and other parts of Laurussia, South China and Northern areas of Gondwana are demonstrated here. Despite the similarities, there is also a degree of endemicism amongst some of the shark species present in the Canning Basin. This is not unusual as endemicism in some conodont (Klapper 1995, 2007) and many placoderm (Long 2006) species has also been noted. Of the Frasnian taxa *Diademodus dominicus* sp. nov. had not been previously recorded outside of Laurussia. The presence of this species along the eastern margins of North Gondwana would also suggest a greater distribution of this genus than previously recognised. Of the other phoebodontids, *Phoebodus* tooth forms found in the Horse Spring section often possess unusual features. These include the tricuspid form of *Ph. bifurcatus* which appear relatively common in the Canning Basin but rare in other areas of the world. The teeth of *Ph. fastigatus* also differ from the diagnostic smooth cusps, often possessing cusps with heavy ornament on the labial face. In addition to the Frasnian taxa, there are two types of ctenacanthiform and a species of *Stethacanthus* that have not been found anywhere else. Despite an increased cosmopolitanism in Devonian shark faunas, a considerable endemic component is present in the Canning Basin. The mechanisms for this are difficult to determine and may be the result of lack of sampling, although unique palaeoenvironmental conditions or environmental niches might have existed in this part of East Gondwana (north-western Australia) at this time.

CONCLUSIONS

A largely undescribed chondrichthyan fauna ranging in age from MN 11 CZ in the late Frasnian to the late *marginifera* Zone in the middle Famennian was recorded from the Virgin Hills Formation, and the Bugle Gap Limestone, and Napier Formation in the Canning Basin. The section at Horse Springs reveals a similar diversity of phoebodonts to that previously described by Trinajstic and George (2009). A new species of diademodont, *Diademodus dominicus* sp. nov., a genus which until now had only previously been described from Laurussia was recovered from the phoebodont-rich beds. This species, in addition to variation in *Phoebodus fastigatus*, as well as newly described cladodont and stethacanthid teeth, reveals a potentially endemic Frasnian component to the Canning Basin.

The Famennian Bugle Gap limestone and uppermost beds of the section at Casey Falls reveals a diverse shark fauna that had not been previously described from the Canning Basin. Many of these, including *Th. tranquillus*, *C. wildungensis*, *P. serra*, *D. mansureae* and *L. lusavorichi*, represent ages similar to or the same as those previously described for these species. The presence of this shark fauna in the Canning Basin also allows for a more comprehensive understanding of species distribution in the Famennian. A number of genera including *Phoebodus* and *Protocodus* occur in both eastern Australia and Western Australia (Turner and Young 2000), however, these new reports on chondrichthyans from the Canning Basin indicate stronger faunal links between the Canning Basin and other areas of Gondwana such as Iran and Morocco as well as Poland and the United States of America in Laurussia. In addition, faunal similarities between South China and the Canning Basin are present, with some shark species indicating potential faunal exchange extending into the Late Famennian and possible Early Carboniferous.
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